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The linkage between hillslope vegetation changes and late-Quaternary fluvial-system aggradation in the Mojave Desert revisited

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Abstract

Valley-floor-channel and alluvial-fan deposits and terraces in the southwestern US record multiple episodes of late Quaternary fluvial aggradation and incision. Perhaps the most well constrained of these episodes took place from the latest Pleistocene to the present in the Mojave Desert. One hypothesis for this episode, i.e. the paleo-vegetation change hypothesis (PVCH), posits that a reduction in hillslope vegetation cover associated with the transition from Pleistocene woodlands to Holocene desert scrub generated a pulse of sediment that triggered a primary phase of aggradation downstream, followed by channel incision, terrace abandonment, and initiation of a secondary phase of aggradation further downstream. A second hypothesis, i.e. the extreme-storm hypothesis, attributes episodes of aggradation and incision to changes in the frequency and/or intensity of extreme storms. In the past decade a growing number of studies has advocated the extreme-storm hypothesis and challenged the PVCH on the basis of inconsistencies in both timing and process. Here I show that in eight out of nine sites where the timing of fluvial-system aggradation in the Mojave Desert is reasonably well constrained, measured ages of primary aggradation and/or incision are consistent with the predictions of the PVCH if the time-transgressive nature of paleo-vegetation changes with elevation is fully taken into account. I also present an alternative process model for PVCH that is more consistent with available data and produces sediment pulses primarily via an increase in drainage density (i.e. a transformation of hillslopes into low-order channels) rather than solely via an increase in sediment yield from hillslopes. This paper further documents the likely important role of changes in upland vegetation cover and drainage density in driving fluvial-system response during semiarid-to-arid climatic changes.

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1 Introduction and motivation

Quaternary deposits of the southwestern US are dominated by valley-floor-channel and alluvial-fan deposits and their associated terraces that record multiple regionally correlative episodes of aggradation, channel incision, and terrace abandonment (Christensen and Purcell, 1985; Bull, 1991; Harvey et al., 1999; Menges et al., 2001; McDonald et al., 2003; Anders et al., 2005). What drives these aggradation and incision episodes has been a fundamental question in the geomorphology and Quaternary geology of the southwestern US for decades. Given the approximate correlation between the timing of fluvial-system aggradation events and semiarid-to-arid transitions recorded in paleoclimatic proxies, together with the correlative nature of Quaternary deposits and terraces across tectonically active and inactive regions, climate change has most often been invoked as the primary trigger for these episodes. How climate change drives episodes of aggradation and incision is debated, however. In this paper I focus on the portion of the geologic record covering the latest Pleistocene to the present because the constraints on both fluvial-system aggradation/incision and its potential driving mechanisms (e.g. paleo-vegetation changes) are most well constrained for deposits of this time period.

In his work on the paleo-vegetation-change hypothesis (PVCH) for fluvial-system response to climatic changes in the southwestern US, Bull (1991) argued that a reduction in vegetation cover during late Quaternary semiarid-to-arid transitions led to sediment pulses characterized by an initial increase in sediment yield (resulting in aggradation in valley-floor channels and/or alluvial fans downstream) followed by a decrease in sediment yield as the reservoir of colluvium stored on hillslopes was depleted (resulting in channel incision and terrace abandonment). As Bull (1991) wrote, "*when the climate changes from semiarid to arid, the concurrent decrease in vegetation cover results in a rapid increase in sediment yield. The sediment-yield maximum is attained quickly, after which the yield progressively decreases as the area of hillslope colluvium decreases and outcrop area increases*." (p. 113). Bull (1991) also invoked changes in

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the frequency and/or intensity of extreme storms, i.e. "*hillslope sediment yields were greatly increased – partly because of increased rainfall intensities associated with the return of monsoon thunderstorms*" (p. 114). As such, Bull (1991) did not envision the PVCH working in isolation. Which of these two driving mechanisms (paleo-vegetation changes or an increase in the frequency and/or intensity of extreme storms) plays a greater role in driving fluvial-system aggradation and incision has been a subject of controversy ever since Bull's seminal work.

As the dating of fluvial-system deposits in the Mojave Desert has improved over the past two decades, numerous studies have used the apparently poor correlation between the timing of paleo-vegetation changes and alluvial-fan aggradation to argue against the PVCH. For example, in the central Mojave subregion of their study, Antinao and McDonald (2013a) tested the PVCH by comparing the timing of fan aggradation at a site (southern Death Valley) sourced by an eroding catchment with lowest elevations of ≈ 400 m a.s.l. against the timing of paleo-vegetation changes constrained by a group of nearby packrat midden sites (which record the vegetation types within a 100 m range) with a lowest elevation of 1060 m a.s.l. (Granite Mountains) (Fig. 1; Supplement of Antinao and McDonald, 2013a). This approximately 600 m gap in elevation between the lowest elevations of the source catchment and the lowest elevation where the timing of paleo-vegetation changes was constrained is problematic because, as I demonstrate in Sect. 2, the transition from woodlands to desert scrub measured at 1060 m a.s.l. occurred 6 ka after vegetation changes in the lowest elevations of the source catchment (i.e. the portion of the catchment responsible for triggering the initiation of fan aggradation according to the PVCH). Antinao and McDonald (2013a) concluded that the onset of fluvial-system aggradation "*began well before changes in catchment vegetation cover*" and, as a result, that "*the ambiguous relation between vegetation change and alluvial fan aggradation indicates that vegetation had a reduced role in LPH aggradation*". To test the PVCH most comprehensively, however, it is necessary to constrain the timing of the woodland-to-desert-scrub transition at the lowest elevations of each source catchment because vegetation changes are time-transgressive

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with elevation (occurring first at low elevations and later at high elevations) and hence it is the vegetation changes at the lowest elevations of source catchments that are responsible for triggering aggradation according to the PVCH.

Miller et al. (2010) concluded that alluvial-fan aggradation in the Mojave Desert was principally due to more frequent and/or more intense storms based on an approximate correlation between the timing of alluvial-fan aggradation and elevated sea-surface temperatures in the Gulf of California (a proxy for monsoon activity). Specifically, both records exhibit two peaks from the latest Pleistocene to the present, i.e. one at ca. 15–7 ka cal BP and another at ca. 6–3 ka cal BP. The dual-pulsed nature of alluvial-fan aggradation from the latest Pleistocene to present was also emphasized by Bull (1991), who noted "*the aggradation event that was associated with the Pleistocene–Holocene climate change consisted of two main pulses (Q3a and Q3b) with an intervening period of stream-channel incision*" (p. 114). Miller et al. (2010) considered the dual-pulsed nature of aggradation to be inconsistent with PVCH. However, two pulses of aggradation separated by approximately 4–8 ka and resulting from a single pulse of sediment yield is precisely what the PVCH predicts (documented in Sect. 2). A sediment pulse from upstream catchments can overwhelm the ability of a downstream valley-floor channel or alluvial fan to convey that increase in sediment, leading to a "primary" phase of aggradation. When sediment supply declines during the waning phase of the sediment pulse, fluvial channels incise into the sediments just deposited and an abandoned terrace is formed. Sediments reworked by channel incision, along with sediments still being supplied by the catchment in the waning phase of the sediment pulse, are then deposited in a "secondary" deposit further downstream (Fig. 2a). In this way, a single pulse of sediment can generate two deposits separated by channel incision, as Bull (1991) stated. In this paper I quantify the timing of both the initiation of primary aggradation and channel-incision/secondary-aggradation as predicted by the PVCH and compare the predictions to available geochronologic data from the Mojave Desert.

The PVCH has also been criticized on the basis of process. The conceptual process model that underpins the PVCH as proposed by Bull (1991) requires that hillslopes be

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stripped nearly to bedrock in some locations, increasing the area of outcrops in the catchment and hence reducing the availability of soil for transport, triggering channel incision and terrace abandonment downstream (Bull, 1991). McDonald et al. (2003) correctly argued that Bull's (1991) conceptual model is inconsistent with the observed abundance of soils on, and the high infiltration capacity of, many Holocene desert hillslopes. Modeling studies (discussed in Sect. 3), however, together with the correlation of drainage density with both aridity and percent bare area in catchments of the southwestern US (Melton, 1957), suggest that the principal effect of a reduction in vegetation cover is not an increase in sediment yield from hillslopes but rather an increase in drainage density as hillslopes are converted into low-order channels capable of rapidly entraining colluvium into the fluvial system (Fig. 2b; Strahler, 1958; Tucker and Slingerland, 1997; Pelletier et al., 2011). More recently, Antinao and McDonald (2013a) challenged the PVCH by noting that Pierson et al. (2007) documented a minimal increase in erosion following the manual removal of *Juniperus* from hillslopes. The Pierson et al. (2007) study is not a proper analog for the woodland-to-desert-scrub transition from the latest Pleistocene to the present in the Mojave Desert, however, because that study was performed in Oregon, erosion was measured in plots less than 10 m wide, and juniper root systems and fallen trees were left intact at the site after cutting (i.e. vegetation cover near the ground surface was higher after cutting than before). In contrast, vegetation cover would be lower and percent bare area would be higher following a woodland-to-desert-scrub transition that occurred over time scales of millennia.

In this paper I demonstrate that the timing of fluvial-system aggradation and incision from the latest Pleistocene to the present in the Mojave Desert is consistent with the PVCH in eight out of nine sites in the Mojave Desert where the timing is reasonably well constrained. I also present an improved model of how the PVCH works that includes transient changes in drainage density associated with semiarid-to-arid climate transitions.

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2 A test of the PVCH for the Mojave Desert

In this section I reexamine the conclusions of Antinao and McDonald (2013a) using the same data but a different methodology. Rather than grouping fan aggradation and paleo-vegetation sites according to spatial proximity and elevation ranges that are similar but nevertheless are associated with differences in the timing of assumed vs. actual paleovegetation changes of up to 6 ka, my methodology honors the dominant role of elevation in controlling plant distributions in the Mojave Desert by first quantifying the relationship between the elevational lower limit of woodland plants vs. time in the Mojave Desert and then applying that relationship to a Geographic Information System (GIS) analysis that predicts the timing of both primary aggradation and channel-incision/secondary-aggradation based on the timing of the woodland-to-desert-scrub transition in the source catchments upstream from every point on the landscape. In both Antinao and McDonald (2013a) and this study, the commonness/abundance of *Juniperus* is used as a proxy for woodland vs. desert scrub (i.e. *Juniperus* rare/absent) vegetation types. Where biomass data are available in the southwestern US, the woodland-to-desert-scrub transition is associated with a step change in biomass (e.g. Whitaker and Niering, 1975). Since woody biomass "*reduces runoff and overland-flow erosion by improving water infiltration, reducing impacts by water droplets, intercepting rain and snow, and physically stabilizing soil by their roots and leaf litter*" (Kort et al., 1998), a transition from commonness/abundance to rarity/absence of *Juniperus* is likely to also be associated with a step-change increase in sediment supply to fluvial systems downstream.

I analyzed the North American Midden Database (Strickland et al., 2005) to identify the elevational lower limit of *Juniperus* as a function of time from the latest Pleistocene to the present in the Mojave Desert (Fig. 3). The solid curve in Fig. 3, which correctly differentiates all 87 (within 2σ uncertainty) available localities in terms of common/abundant vs. rare/absent *Juniperus*, illustrates the dominant role of elevation in controlling woodland vs. desert scrub vegetation types from 17–0 ka cal BP. The consistency of the data, i.e. within age uncertainty all common/abundant records reside on one side of the solid curve in Fig. 3 and all rare/absent records reside on the other side, provides confidence in using the curve to predict the timing of the woodland-to-desert-scrub transition within source catchments upstream from aggradation sites in the Mojave Desert using elevation data.

Primary aggradation is assumed to be initiated downstream when a small but significant portion (5 % is used here) of the source catchment area (defined as upstream areas with slopes > 20 %) changes from commonness/abundance to rarity/absence of *Juniperus* (Figs. 4 and 6). Primary aggradation is assumed to terminate when 50 % of the source catchment area has transitioned from commonness/abundance to rarity/absence of *Juniperus*, at which point the source area for sediment derived from woodlands-to-desert-scrub transition is assumed to decline over time, triggering channel incision, terrace abandonment, and secondary aggradation further downstream. The GIS analysis associates each pixel in the source catchment with the age of the woodland-to-desert scrub transition using the solid curve in Fig. 3. For each pixel, data from the source catchment (defined as all pixels that route water/sediment to the downstream pixel using the multiple flow direction routing (MFD) algorithm of Freeman, 1991) are used to determine the timing of the initiation of primary aggradation and channel-incision/secondary-aggradation based on the 5 % and 50 % thresholds respectively. For each pixel, the percent of the source catchment area that has undergone a change from commonness/abundance to rarity/absence of *Juniperus* is computed in 1 ka intervals starting at 15 ka cal BP. When 5 % of the source catchment area has undergone a change from commonness/abundance to rarity/absence of *Juniperus*, primary aggradation is assumed to be initiated. Similarly, when 50 % of the source catchment area has undergone a change from commonness/abundance to rarity/absence of *Juniperus*, channel-incision/secondary-aggradation is assumed to be initiated. The 5 % and 50 % threshold values are not unique, but the results are not highly sensitive to these values within reasonable ranges, i.e. the predicted ages of aggradation and incision differ by 2 ka or less for 96.5 % and 98.6 % of the study area

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where aggradation is predicted to occur as these values are varied over reasonable ranges from 5 % to 15 % and 40 % to 60 %, respectively. The systematic uncertainty of the GIS analysis is taken to be 2 ka (Table 1) based primarily on the uncertainty associated with the 5 % and 50 % thresholds. I used the MFD method of Freeman (1991) to delineate source areas because it more faithfully represents flow routing pathways in distributary channel systems compared with alternatives such as the $D\infty$ method (Pelletier, 2008).

The results of the GIS analysis indicate that the PVCH predicts that the initiation of primary aggradation associated with vegetation changes from the latest Pleistocene to the present was highly variable in space and occurred over a wide range of time from ca. 15–6 ka cal BP (but possibly earlier given uncertainty in the timing of woodland-to-desert scrub transition at the lowest elevations ca. 17–15 ka cal BP; see dashed line in Fig. 3). According to the PVCH, aggradation began earlier at sites fed by catchments of lower elevation and later at sites fed by catchments of higher elevation (Figs. 4, 6, and 7; Table 1). The predictions of the PVCH are consistent with eight out of nine fan aggradation sites in the Mojave Desert (Table 1). The PVCH predicts that the duration of primary aggradation (which is also equivalent to the time lag between the initiation of primary aggradation and channel-incision/secondary-aggradation) was also highly variable, ranging from 4–8 ka at sites fed by catchments with a wide range of elevations in the 750–1800 m range, to 1–4 ka at sites fed by catchments with a smaller range of such elevations (Fig. 5 and Table 1).

It should be noted that the GIS analysis predicts the timing of aggradation in some upland catchments where fill terraces may not form or may not be preserved. It is nevertheless appropriate to include such areas in the analysis because aggradation can occur locally in such areas. Fill terraces are preserved, for example, in many high-elevation, rapidly eroding catchments of the Transverse Ranges (located in the SW corner of the study area) (Bull, 1991). A more sophisticated analysis would predict the timing of both aggradation/incision and the likelihood of preservation of a fill deposit, but such an approach would introduce new parameters into the analysis and is unnec-

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essary for addressing the key questions of this paper (i.e. the timing of aggradation and incision). Also, flow-routing pathways are identified using the modern DEM (ASTER v2 GDEM), yet channels have been modified by late Quaternary fluvial-system aggradation and incision. As such, some portions of alluvial fans and lowland basins comprised of late Pleistocene and/or Holocene deposits are now disconnected from source catchments (via channel incision and terrace abandonment) and hence appear as black in Figs. 4–6. In such areas one can estimate the predictions of the PVCH approximately using the predicted ages from nearby areas of similar elevation.

The study region considered in this paper is large, i.e. the entire Mojave Desert and some portions of adjacent regions. Adopting such a large study area has the advantage that a lot of data can be brought to bear in order to precisely constrain the relationship between vegetation changes, elevation, and time/age (Fig. 3). One could argue, however, that the study area (Fig. 4) is too large for a single paleo-vegetation change vs. age relationship (i.e. Fig. 3) to be applicable throughout the region (as assumed in the above analysis). We can address this issue, at least for modern vegetation (Fig. 8). I used gridded 30 m pixel^{-1} estimates of shrub and tree cover from the US Geological Survey's (2013) LANDFIRE database to quantify the relationship between percent shrub/tree cover and elevation using 100 m bins of elevation from 500 m to 2000 m a.s.l. I divided the study region illustrated in Fig. 4 (i.e. 34–36° N and 115.0–117.5° W) into 5 equally-spaced bins of latitude and computed the relationship between average shrub/tree cover and elevation for each bin separately (Fig. 8). The results demonstrate that there is no systematic change in the relationship between modern vegetation cover and elevation with latitude from 34–36° N in the Mojave Desert. That is not to say that elevation is a perfect predictor of vegetation cover. Clearly slope gradient and aspect, spatial variations in soil properties, etc. influence the elevational zonation of plants in the southwestern US locally. The scatter of the data plotted in Fig. 8 reflects the significant spatial variations in plant cover that exist within areas of equal elevation. Rather, this analysis demonstrates that any systematic differences in

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average vegetation cover across latitudes is much smaller than the differences across elevation.

3 An improved process model for PVCH

Melton (1957) documented a strong positive correlation between drainage density and both aridity and percent bare area in the southwestern US. Melton's findings provide a basis for predicting an increase in drainage density during semiarid-to-arid transitions. Similarly, Pelletier et al. (2013) quantified drainage density across the elevation/precipitation/vegetation gradient of the Santa Catalina Mountains in Arizona (from Sonoran desert scrub at low elevations that have a mean annual precipitation of $0.2\,\mathrm{m\,a^{-1}}$ through pinyon/juniper woodland to mixed conifer forest at high elevations that have a mean annual precipitation of $0.8\,\mathrm{m\,a^{-1}}$) and obtained a similar inverse relationship between drainage density and water-availability/vegetation-cover. Unfortunately, we lack studies that demonstrate an increase in sediment yield from hillslopes and/or drainage density resulting from a transition from woodland to desert scrub vegetation in the Mojave Desert specifically. However, anthropogenic disturbances in the Mojave Desert that are associated with reductions in vegetation cover consistently result in higher erosion rates from hillslopes and low-order fluvial valleys (e.g. Lovitch and Bainbridge, 1999; Iverson, 2006; and references therein).

Theoretical models for the controls on drainage density further suggest that vegetation cover is the most important climatically related variable controlling drainage density. Perron et al. (2008, 2009), for example, demonstrated that the spacing of first-order valleys (closely related to drainage density) is a function of the relative rates of colluvial sediment transport (dominant on hillslopes) and slope-wash/fluvial erosion (dominant in valley bottoms). A reduction in vegetation cover decreases the rate of colluvial sediment transport because fewer plants are present to drive bioturbation, while simultaneously increasing slope-wash/fluvial erosion rates via a reduction in protective cover. Both of these effects combine to promote higher drainage density. Multiple lines

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of evidence also indicate that sediment yield is far more sensitive to changes in vegetation cover (via a threshold shear stress for detachment or entrainment) than changes in runoff intensity (Tucker and Slingerland, 1997). The critical shear stress criterion for detachment is (Tucker and Slingerland, 1997)

$$P_s A S^2 > \tau_c^3 \quad (1)$$

where P_s is runoff ($\mathrm{L\,T^{-1}}$), A is contributing area ($\mathrm{L^2}$), S is slope ($\mathrm{L\,L^{-1}}$), and τ_c is the shear stress threshold that depends on vegetation cover ($\mathrm{L\,T^{-1/3}}$). Equation (1) shows that a 2-fold decrease in τ_c is equivalent to an 8-fold increase in in runoff intensity. I cannot quantify τ_c for desert scrub vs. woodland vegetation types, but Prosser and Dietrich (1995), working in central coastal California, documented a strong sensitivity of τ_c to percent bare area, and Melton (1957) documented a strong correlation between drainage density and both aridity and percent bare area in the southwestern US. According to Eq. (1), a 2-fold decrease in τ_c can increase the drainage density (which goes as the square root of contributing area) by almost 3-fold.

Figure 9 presents a schematic diagram of an alternative conceptual model for the PVCH that includes elevation and is qualitatively based on the results of numerical modeling studies (e.g. Tucker and Slingerland, 1997). Numerical models predict that sediment pulses produced via an increase in drainage density are temporary (i.e. they are sediment *pulses* with both waxing and waning phases) because first-order valley heads eventually stabilize (following a transition from lower to higher drainage density) as they become adjusted to a smaller contributing area associated with the lower vegetation cover and hence lower shear stress threshold for detachment of arid climates. The stabilization of valley heads following drainage network expansion causes a reduction in sediment supply that triggers channel incision, terrace abandonment, and secondary aggradation further downstream (Tucker and Slingerland, 1997). The relationships among vegetation cover, drainage density, and sediment yield according to this conceptual model are shown for three hypothetical drainage basins with the same relief (200 m) but different mean elevations. Figure 9 assumes only that drainage density varies inversely with vegetation cover, that sediment supply is driven by changes in drainage density over time (via the conversion of hillslopes into low-order channels), and that vegetation cover varies with time according to Fig. 3.

4 Discussion

The predictions of the PVCH are consistent with eight out of nine sites of aggradation and incision in the Mojave Desert (Figs. 4–7, Table 1). Aggradation begins at a time consistent with the predictions of the PVCH in all cases except one (Chambless) in which the measured initiation of aggradation occurs significantly later than the prediction (Table 1). There is no discrepancy, however, if the dated deposit at Chambless corresponds to a site of secondary aggradation (i.e. if sediments with a depositional age ca. 15 ka cal BP are located upstream) because the predicted age of channel-incision/secondary-aggradation according to the GIS analysis is consistent with the measured age (Table 1). The Chambless site is not included in Fig. 7 because of this ambiguity. The likelihood that two pulses of aggradation will occur from one pulse of sediment yield complicates testing of any model for late Quaternary fluvial-system aggradation, but this complexity can be addressed in future studies with detailed geologic mapping and simultaneous dating of both primary and nearby/inset secondary late Quaternary deposits.

The results of this paper are consistent with the time-transgressive nature of aggradation and incision with elevation documented by Weldon (1986) in his study of the late Quaternary history of Cajon Pass (located near the southwestern corner of the study area shown in Fig. 4) despite significant differences in vegetation types between the Mojave Desert and the Transverse Ranges (Fig. 10). Weldon (1986) documented a wave of aggradation followed by incision that moved up Cajon Pass from elevations of 500 m a.s.l. (i.e. the Freeway Crossing) to 1500 m a.s.l. (the summit) during the time interval from 15 to 6 ka cal BP (Fig. 10a). Weldon's (1986) study is particularly valuable because it demonstrates the time-transgressive nature of aggradation with elevation at

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a single site rather than by combining many study sites within a large region. Weldon (1986) also demonstrated that the hillslope sediment yield from 15–6 ka cal BP was an order of magnitude higher than sediment yields during either the late Pleistocene or the mid-late Holocene. The PVCH accurately predicts the time-transgressive wave of aggradation documented by Weldon (1986) (Fig. 10b).

The PVCH purportedly fails in two sites considered by Antinao and McDonald (2013a), i.e. the western side of the Providence Mountains (dated by Clarke, 1994) and the Sierra El Mayor piedmont of Baja California deposits (dated by Spelz et al., 2008 and Armstrong et al., 2010). The Baja California sites are not part of the Mojave Desert but are important to address here because, according to Antinao and McDonald (2013a), they provide a basis for a reduced relevance of PVCH in the late Quaternary evolution of the southwestern US. Fan aggradation is not always triggered by climatic variations, however, so it is important to consider other possible triggers for aggradation when evaluating the PVCH in specific cases. The Kelso Dunes have migrated across the distal portion of the western piedmont of the Providence Mtns. in late Quaternary time. The site location map in Fig. 2 of Clarke (1994) and the stratigraphic columns in Fig. 3 of Clarke (1994) clearly show that the deposits dated by Clarke (1994) include aeolian sand deposits at least 1 m thick at each sample locality and grade to aeolian sand deposits at least 8 m thick at the distal end of the fan. Seven out of the eight sediment samples dated by Clarke (1994) were aeolian sediments and the sole fluvial sediment sample dated was located directly atop aeolian sediments. Therefore, despite the common interpretation that the Clarke (1994) ages record climatically driven fluvial-system aggradation, it is at least possible that the ages measured by Clarke (1994) more strongly reflect the history of the Kelso Dunes in addition to fluvial aggradation triggered by the local base-level rise associated with the migration of the Kelso Dunes across the distal portion of the fan. Given this possibility, the Clarke (1994) data may not be the most reliable data to use when testing alternative models for the climatic triggering of fluvial-system aggradation.

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Similarly, fan deposits on the eastern piedmont of Sierra El Mayor dated by Armstrong et al. (2010) are potentially problematic for the purposes of testing the PVCH because they were deposited atop, and shortly following deposition of, fluvial and deltaic sediments of the Colorado River (Armstrong et al., 2010), which today is located less than 1 km from the dated fan deposits. As such, aggradation of the Colorado River may have resulted in a local increase in base level, triggering aggradation of fan sediments without a climatically driven change in upstream sediment supply. Although the response of fluvial channels to base-level rise depends on a number of factors including the magnitude of the base level change and the slope of the channel affected by base-level changes (Schumm, 1993), Leopold and Bull (1979) presented one specific example in which a modest increase in base level (i.e. ~ 3 m) triggered aggradation at a distance of ~ 1 km. As such, it is possible that deposition on the western piedmont of Sierra El Mayor is principally influenced by the base-level control exerted by the Colorado River rather than by a climatically driven increase in sediment supply.

Spelz et al. (2008) dated fan aggradation on the western piedmont of the Sierra El Mayor that is unaffected by the Colorado River. However, the cosmogenic ages of Spelz et al. on boulders of the terrace associated with latest Pleistocene aggradation varied by 300 % (700 % if all of the data are considered – one date of 76 ka cal BP was excluded in the average). As such, the age control at this site is not ideal. Spelz et al. (2008) used a weighted average of boulder ages to obtain an estimated surface age of 15.5±2.2 ka cal BP (i.e. well before the retreat of late-Pleistocene plants at 10.7±0.5 ka cal BP as constrained by the presence of boojum tree (*Fouquieria columnaris*), Anderson and Van Devender, 1995). However, a more accurate surface-exposure age is, in many cases, obtained by using the youngest boulder age that is not an obvious outlier, especially when inheritance is a potentially important factor (e.g. Applegate et al., 2010; Heyman et al., 2011). In this case that age would be 11±1.3 ka cal BP, i.e. broadly consistent with the timing of paleo-vegetation changes. A significant amount of text in this discussion has been devoted to the Providence Mountains and Sierra El Mayor sites despite the fact that they are not central to this paper. I feel compelled to

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address these sites because they have been used by Antinao and McDonald (2013a) to challenge the PVCH. However, I don't want the discussion of these sites to distract from the larger point: with or without these sites, the vast majority of data from the Mojave Desert are consistent with the PVCH if the time-transgressive nature of paleo-vegetation changes with elevation is fully taken into account.

Many authors, including Bull (1991), Harvey et al. (1999), McDonald et al. (2003), Miller et al. (2010), and Antinao and McDonald (2013a, b) have invoked changes in the frequency and/or intensity of extreme storms to drive fluvial-system aggradation in the southwestern US. Miller et al. (2010), for example, documented a correlation between sea-surface temperatures in the Gulf of California (a proxy for monsoon activity) and alluvial-fan aggradation in the Mojave Desert. The role of monsoon thunderstorms specifically in driving aggradation in the Mojave Desert is likely minimal given the limited impact of the North American Monsoon (NAM) system on the Mojave Desert (Higgins et al., 1997). Recent investigations of the spatial extent of NAM show no significant impact on the Mojave Desert (e.g. Dominguez et al., 2009). Antinao and McDonald (2013b) also argued against a significant role for NAM in driving fan aggradation in the southwestern US, invoking changes in the frequency and/or intensity of tropical storms instead.

It is likely that changes in extreme storms played a role in changing Mojave Desert landscapes in the late Quaternary, but there is no evidence based on either timing or process that changes in the frequency and/or intensity of extreme storms was the dominant cause of fluvial-system aggradation and incision. The correlation between sea-surface temperatures in the Gulf of California and alluvial-fan aggradation in the Mojave Desert is excellent for the 6–3 ka cal BP time period but significantly poorer for the latest-Pleistocene-to-early-Holocene time period, i.e. primary aggradation occurred ca. 14–7 ka cal BP while elevated sea-surface temperatures occurred 15–11 ka cal BP. Antinao and McDonald (2013b) argued that more frequent and/or intense El-Niño-like conditions in the tropical Pacific increased moisture delivery to the southwestern US ca. 14.5–8 ka cal BP, triggering fan aggradation. Neither of these mechanisms

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(increased monsoon activity or increased El-Niño-like conditions in the tropical Pacific), however, honors the time-transgressive nature of fluvial-system aggradation in the Mojave Desert with elevation, in which the initiation of primary aggradation occurred earlier, i.e. ca. 15–12 ka cal BP at lower-elevation sites, i.e. 400–900 m a.s.l., and later, i.e. ca. 12–6 ka cal BP, at higher-elevation sites, i.e. 900–1500 m. If changes in extreme storms were the dominant cause of fluvial-system aggradation, it seems likely that the timing of aggradation to be relatively elevation-independent. The results presented here, however, further demonstrate that the initiation of late-Quaternary fluvial-system aggradation was time-transgressive with elevation in a manner that closely tracks the retreat of the elevational lower limit of woodlands to higher elevations from ca. 15–6 ka cal BP.

There is also little process-based evidence that changes in extreme storms led to fluvial-system aggradation where we observe it to have occurred. While it is likely that more frequent and/or more intense storms would have led to an increase in sediment yield from hillslopes during the latest Pleistocene to early Holocene in the Mojave Desert, fluvial-system aggradation requires an increase in sediment yield from hillslopes *without a comparable increase in the ability of fluvial channels to transport that increase in sediment*. That is, any climate change that simultaneously increases hillslope sediment yield *and* the ability of fluvial channels to convey that increase in sediment yield would not be expected to result in aggradation except in the lowest elevations of closed basins (e.g. playas). In contrast, the conceptual model for the PVCH presented here predicts valley-floor and alluvial-fan aggradation specifically because the sediment pulse from hillslopes was not accompanied by a significant increase in the ability of fluvial channels to convey that increase in sediment yield.

5 Conclusions

This paper builds upon previous studies (e.g. Weldon, 1986) as well as comprehensive new geochronologic studies (e.g. Miller et al., 2010) to further demonstrate that fluvial-

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system aggradation from the latest Pleistocene to the present in the southwestern US was time-transgressive with elevation. As such, the timing of the initiation of primary fluvial-system aggradation and channel-incision/secondary-aggradation exhibit spatial variability that depends sensitively on the elevation ranges of source catchments. In order to predict these timings for specific locations using the PVCH, I tightly constrained the relationship between paleo-vegetation changes and elevation and then used a GIS analysis that incorporated the elevation ranges of source catchments. The PVCH predicts the correct timing of the initiation of aggradation in eight out of nine cases in the Mojave Desert where reasonable age control exists. It is likely that changes in the frequency or intensity of extreme storms have contributed to cycles of fluvial-system aggradation and incision in the southwestern US, but the results of this paper suggest that the recent trend of discounting the importance of hillslope vegetation changes deserves reexamination.

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Table 1. Comparison of available data and model predictions for the timing of fluvial-system aggradation and incision from the latest Pleistocene to the present. Sites are listed in order of increasing elevation to highlight the relationship between elevation and the age of initiation of primary aggradation in the data and model.

Locality	UTM Easting (m)[5]	UTM Northing (m)	Elevation[6] (m a.s.l.)	Predicted initiation primary aggrad.[3] (ka BP)	Predicted incision[3] (ka BP)	Measured initiation primary aggrad.[2] (ka BP)	Measured incision[1] (ka BP)
Southern DV[7]	525 510[6]	3 978 760[6]	0	15 ± 2	8 ± 2	17.4 ± 2.4	10.7 ± 1.54
Chambless[8]	639 525	3 818 044	250	15 ± 2	8 ± 2	< 9.4 ± 0.67[4]	N/A
Kelso Wash[8]	594 631	3 876 604	360	13 ± 2	5 ± 2	< 13.3 ± 0.7	N/A
Red Pass[9]	564 267	3 909 254	500	13 ± 2	6 ± 2	12.1 ± 0.6 12.8 ± 0.5 11.5 ± 0.2	N/A
Fenner Wash[8]	658 533	3 843 945	520	12 ± 2	6 ± 2	12.1 ± 0.77 11.1 ± 0.3 10.6 ± 0.26	N/A
Coyote Wash[8]	522 313	3 875 008	540	11 ± 2	9 ± 2	< 13.6 ± 0.3	N/A
Sheep Creek[8]	451 338	3 829 441	870	10 ± 2	5 ± 2	10.6 ± 0.56 11.8 ± 0.56	7.74 ± 0.54 8.31 ± 0.51 9.44 ± 0.49
Johnson V.[10]	550 100[6]	3 796 840[6]	980	8 ± 2	6 ± 2	< 12.3 ± 0.9 10.3 ± 1.1	N/A
Grassy V.[8]	477 828	3 902 734	1040	6 ± 2	5 ± 2	< 10.6 ± 0.6	N/A

[1] Channel incision/terrace abandonment ages are given in the two cases where the report makes clear that the sample came from near the top of the deposit.
[2] Ages indicated by < are maximum values for the initiation of alluvial aggradation because the underlying unit (groundwater discharge or fluvial) was dated or because partial bleaching of TL samples was noted.
[3] Uncertainty associated with model predictions is estimated to be 2 ka based on limitations of the paleovegetation record and the magnitude of typical variations in the model predictions with reasonable variations in the model.
[4] Aggradation is initiated much later than predicted at this site. Deposit could be associated with secondary aggradation.
[5] UTM coordinates are in NAD83 datum, zone 11.
[6] Estimated from aerial photographs and Digital Elevation Models.
[7] Sohn et al. (2007).
[8] Miller et al. (2010).
[9] Mahan et al. (2007).
[10] Rockwell et al. (2000).

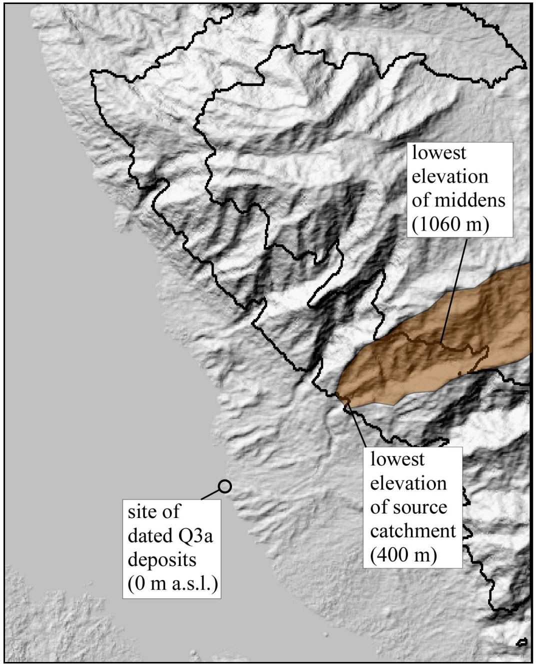

Fig. 1. Illustration of the key elevations associated with the southern Death Valley fan aggradation site. Location is shown in Fig. 4. Samples were collected by Sohn et al. (2007) at approximately 0 m a.s.l. The timing of the initiation of fan aggradation is controlled, according to the PVCH, by the timing of the retreat of the lower elevational limit of woodland vegetation through the lowest elevations of the source catchment (shown in brown), i.e. 400 m a.s.l. In their central Mojave subregion, Antinao and McDonald (2013a) used midden records with lowest elevations of 1060 m a.s.l. (Supplement of Antinao and McDonald, 2013a), resulting in an elevational gap of approximately 600 m. This gap produces a 6 ka time lag between assumed and actual paleo-vegetation changes in the lowest elevations of the source catchment.

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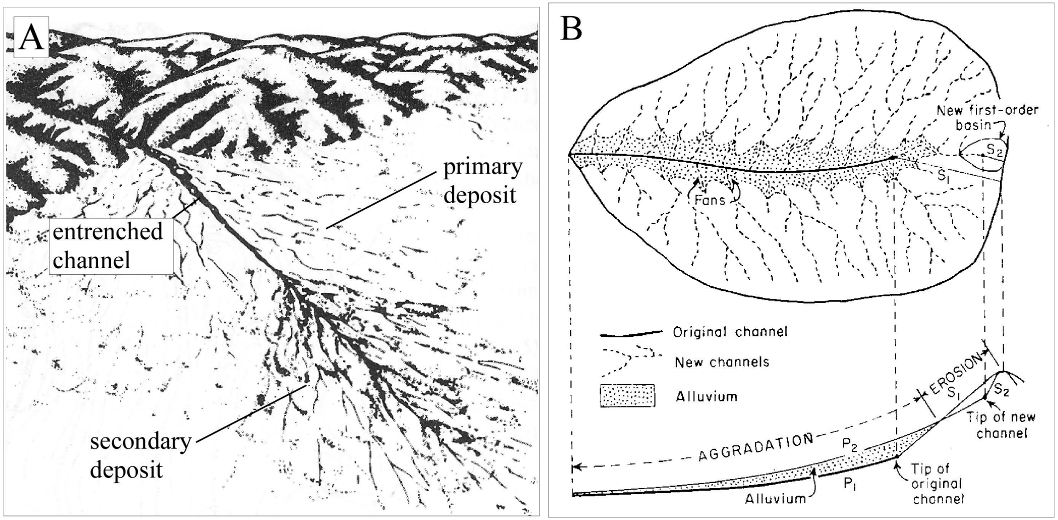

Fig. 2. Schematic diagrams drawn from the literature illustrating **(A)** fan and **(B)** source-catchment responses to semiarid-to-arid transitions. **(A)** Illustration of paired deposits (i.e. a secondary deposit downstream and inset into a primary deposit) resulting from a single pulse of sediment (from Bull, 1968). In **(A)**, a pulse of sediment from upstream catchments leads to a "primary" phase of aggradation. When sediment supply declines during the waning phase of the sediment pulse, fluvial channels incise into the sediments just deposited and an abandoned fan terrace is formed. Sediments reworked by channel incision, along with sediments still being supplied by the catchment in the waning phase of the sediment pulse, are then deposited in a "secondary" deposit farther downslope. **(B)** Schematic diagram illustrating the response of a catchment to an increase in erosivity (from Strahler, 1968). Low-order drainages grow headward, converting hillslopes and into low-order fluvial channels capable of rapidly transporting material stored as colluvium during the previous period. When channels cease growing headward (i.e. when the catchment stabilizes to a new, higher drainage density), the pulse of sediment wanes and channel-incision/secondary-aggradation is triggered (not shown).

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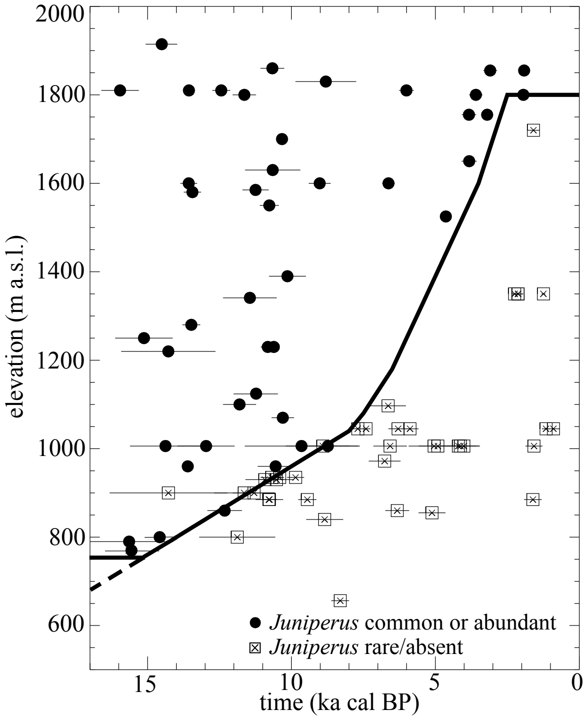

Fig. 3. Plot of every midden in the North American Midden Database (Strickland et al., 2005) between 34° and 37° N latitude, 117.5° and 115° W longitude, 500 m and 2000 m a.m.s.l., and 1 and 17 ka in age (cal BP; ^{14}C years were converted using IntCal09 of Reimer et al. (2009); 95 % confidence intervals are shown) with *Juniperus* common/abundant (closed circles, total of 46 samples) and *Juniperus* rare/absent (open squares; total of 41 samples). In the latter cases, *Juniperus* is specifically noted as absent or a complete taxa list is available that does not include *Juniperus*. The solid curve is the lower elevational limit of woodlands adopted in the model. The dashed line indicates that the lower limit of common/abundant *Juniperus* is poorly constrained from 17–15 ka.

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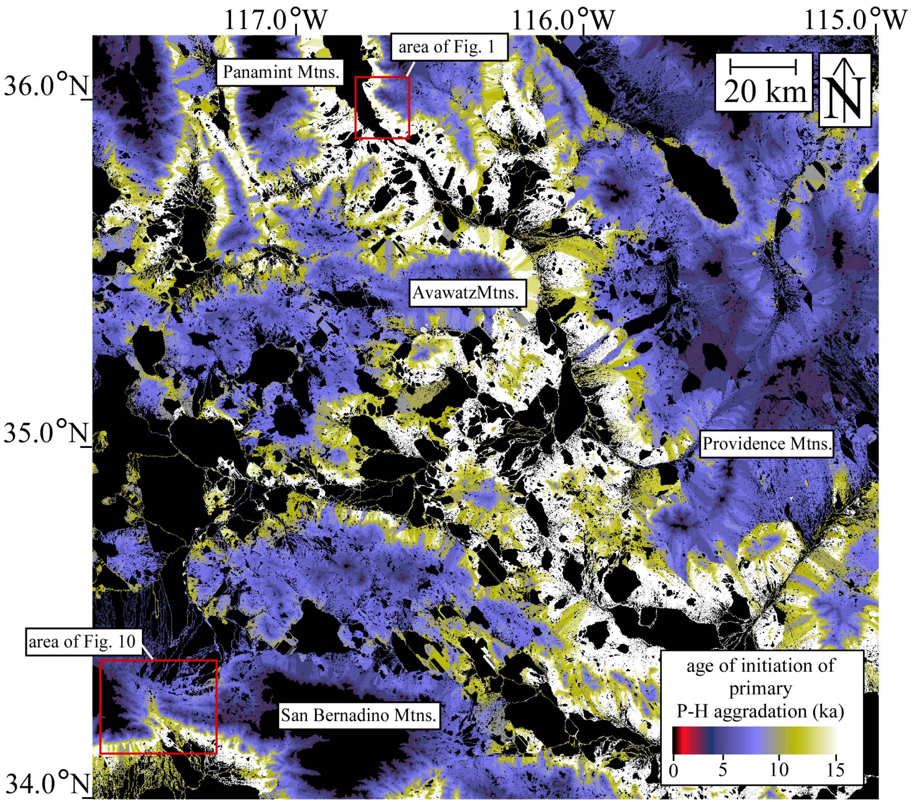

Fig. 4. Color map of the model-predicted age of initiation of aggradation associated with vegetation changes of the latest Pleistocene to mid Holocene in the Mojave Desert (where deposits of such ages are present). Black areas are locations not downstream (based on flow pathways defined by the modern DEM) from areas that have undergone P–H changes in *Juniperus* cover.

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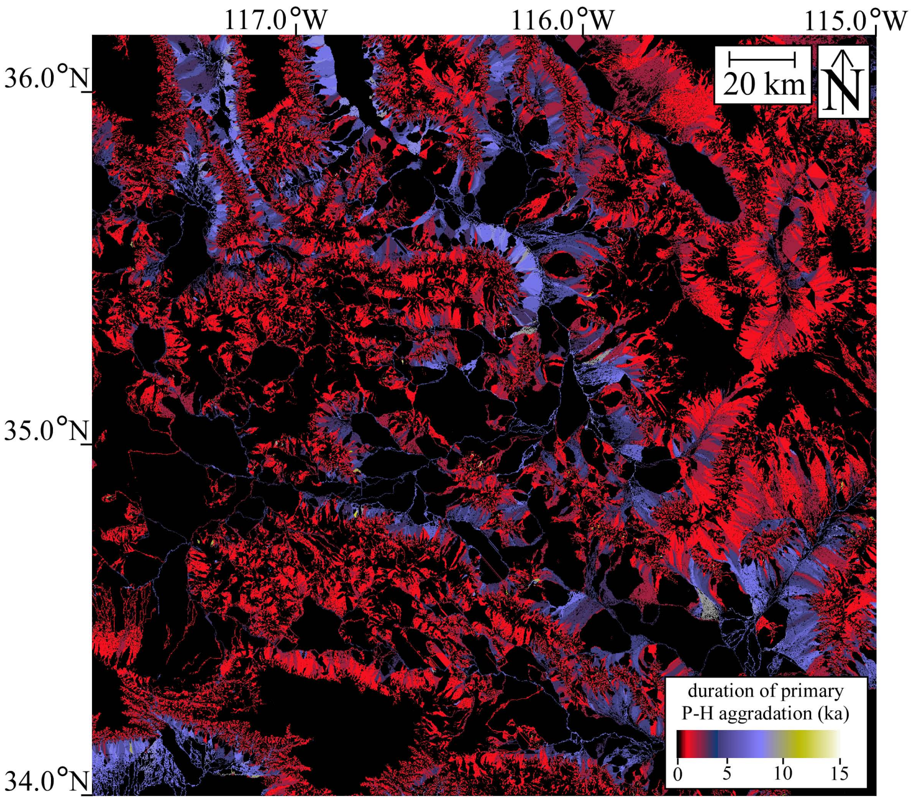

Fig. 5. Color map of the duration of primary aggradation associated with vegetation changes of the latest Pleistocene to mid Holocene in the Mojave Desert. Values range from 4–8 ka at alluvial sites fed by catchments with a wide range of elevations in the 750–1800 m range (typically lower-elevation sites) to 1–4 ka at sites fed by catchments with a smaller range of such elevations (typically higher-elevation sites).

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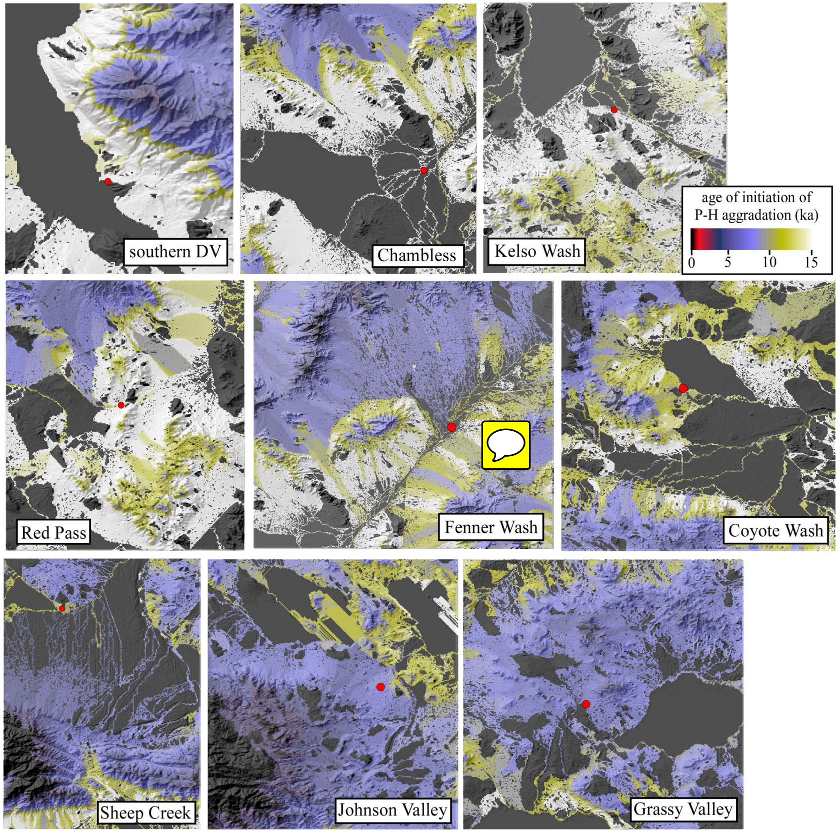

Fig. 6. Subsets of Fig. 4 shown in greater detail and overlain on a shaded-relief image for each of the nine areas where the timing of fan aggradation is reasonably well constrained. Study areas are shown as lowest to highest elevation (of the dated sample locations) from upper left to lower right. The predicted age of initiation occurs systematically later at higher elevations, as also documented in Table 1 and Fig. 7.

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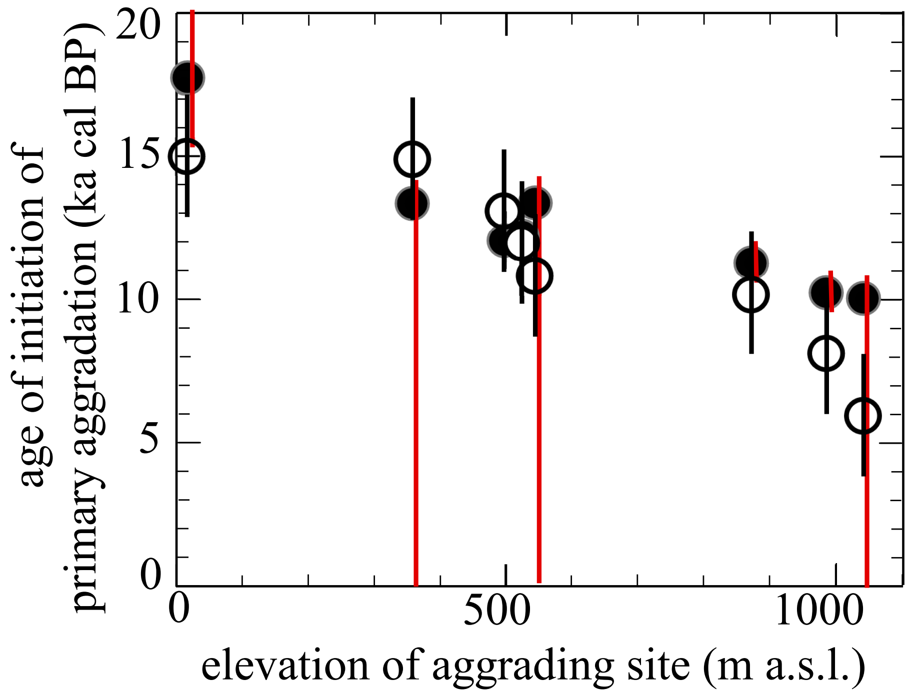

Fig. 7. Plot of measured (closed circles) and predicted (open circles) ages of the initiation of primary aggradation according to the PVCH. All data from Table 1 are plotted except for Chambless site, which was not included because of ambiguity in whether the site represents primary or secondary aggradation (see Sect. 4 for discussion). Uncertainties of the measured ages are shown in red. Ages increase with decreasing elevation of the site of aggradation (which correlates with the lowest elevations of the source catchments). Uncertainty values are provided and explained in Table 1.

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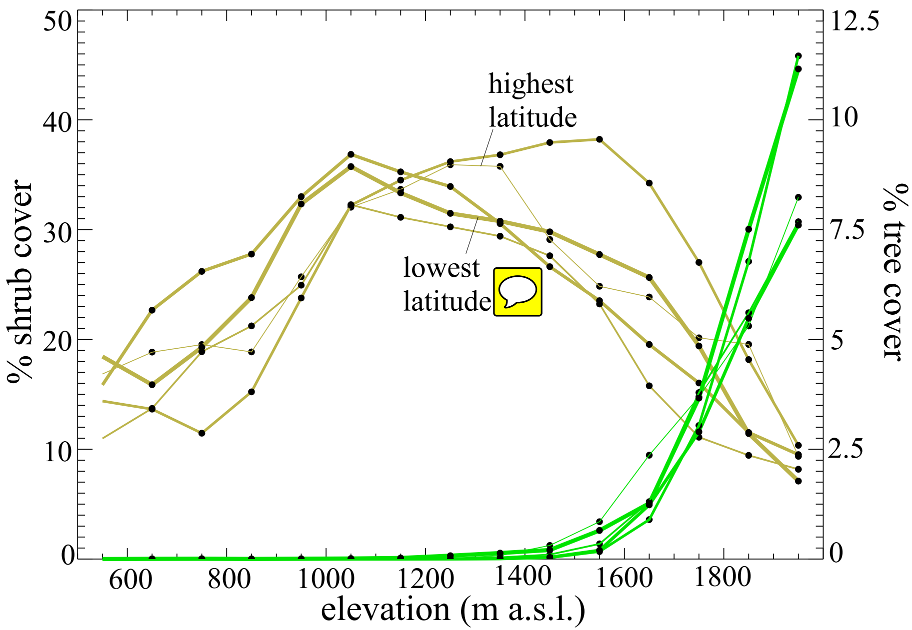

Fig. 8. Plot of the average percent shrub (in brown) and tree (in green) cover in the Mojave Desert as a function of elevation for 5 bins of latitude (thinner lines indicate higher latitudes) using the US Geological Survey's (2013) LANDFIRE database. The data exhibit some spatial variation in the vegetation-cover–elevation relationship, but there is no systematic difference across latitudes.

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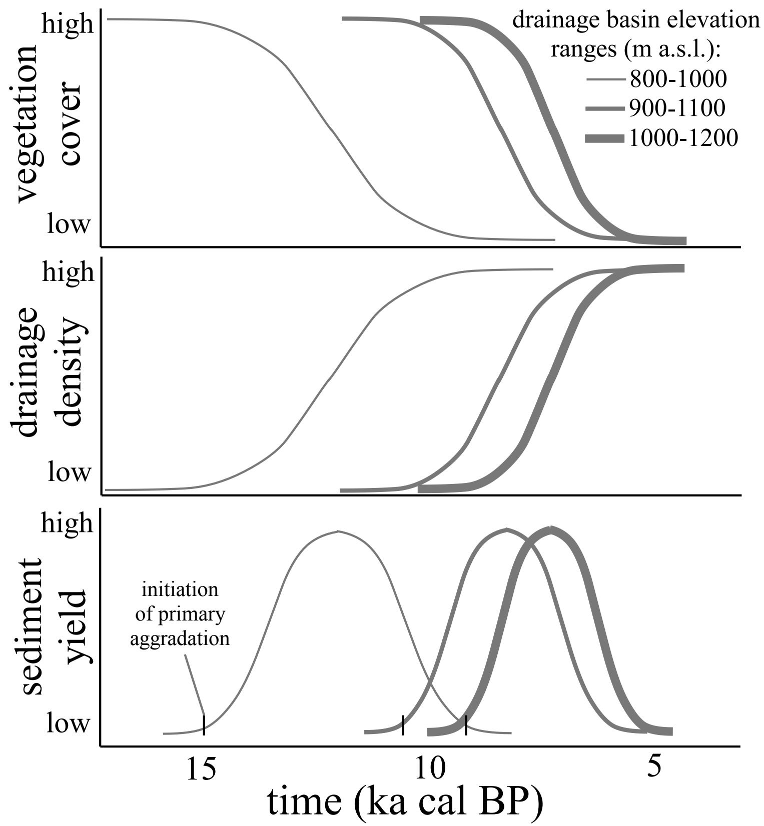

Fig. 9. Conceptual model illustrating relationships among vegetation cover, drainage density, and sediment yield from source catchments for three hypothetical catchments (each with 200 m of relief) of different mean elevation.

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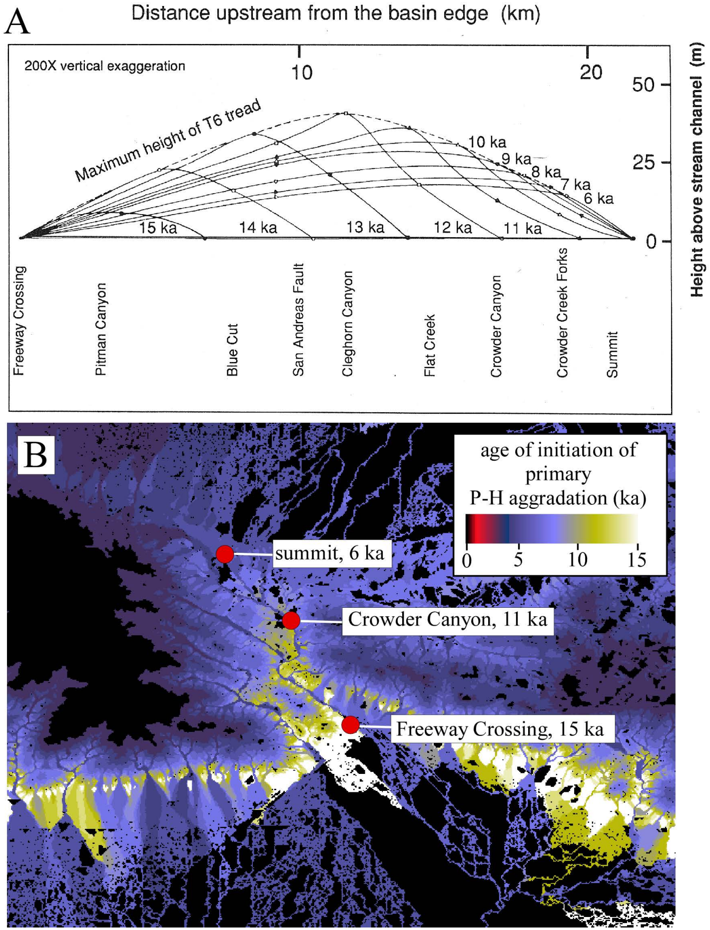

Fig. 10. Comparison of **(A)** measured (from Bull, 1991, after Weldon, 1986) and **(B)** predicted age of initiation of primary aggradation in Cajon Pass, California, according to the PVCH. Location of **(B)** shown in Fig. 4.

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