Identification of stable areas in unreferenced laser scans for automated geomorphometric monitoring

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Abstract. Current research questions in the field of geomorphology focus on the impact of climate change on several processes causing subsequently natural hazards. Geodetic deformation measurements are a suitable tool to document such geomorphic mechanisms e.g. under the use of terrestrial laser scanners that capture a region of interest in a quasi-laminar fashion which results in a so-called 3D point cloud. In this context the central problem is linked to the transformation of 3D point clouds captured at different points in time (epochs) into a common reference coordinate system. To date, this step has been carried out by usage of artificial targets and/or control points. Several drawbacks are related to this strategy such as the enormous effort to distribute the targets in object space, the required survey by additional geodetic sensors such as total stations or GNSS-receivers as well as the limited extent within the region of interest. In this contribution a surface-based registration methodology, termed the iterative closest proximity algorithm (ICProx), that solely uses points as input, similar to the iterative closest point-algorithm (ICP), and hence does not require any artificial targets or extracted geometric primitives, such as planes. The aim of this study was to automatically determine deformations that occurred at a rock glacier, an ice glacier as well as in a rock fall area. For every example two epochs were processed while the ICProx-algorithm's classification accuracy is 70% on average.

1   Introduction

The task of monitoring surface changes in hazardous areas has been an important task of the geodetic community in the last decades. Due to the predicted increase of occurrences (Anderson and Bausch 2006) this problem domain will hence gain importance in the future. Potential forms of geomorphic natural hazards are rock falls or slides, landslides as well as slow surface movements which may be triggered or enforced by heavy rainfall events. Apart from geomorphic changes the characteristics and geomorphic consequences of glacial retreat are a crucial part of scientific research (Avian et al. 2017). This interest can be justified by e.g. the function of glaciers as water reservoirs (Bradley et al. 2006) that act as water supply for agriculture and communities downstream (Tirado et al. 2010). A valuable tool for the quantitative analysis of the earth's

surface is represented by geomorphometry. The most common forms of geomorphometric monitoring aims to quantify kinematic (Schwalbe et al. 2008), areal (Lambrecht and Kuhn 2007) and/or volumetric changes over time. A particularly appropriate sensor for this task is terrestrial laser scanning (TLS) due to its combination of high precision and information density (Avian et al. 2009). In order to process the acquired TLS data, an appropriate deformation model has to be chosen

5   (Heunecke and Welsch 2000), while generally the so-called congruence model is applied. Deformation measurements require the repeated acquisition of a region of interest (ROI) at different points in time that are referred to as epochs. Based on these multi-temporal measurements geometric changes that occurred in between epochs can be identified. Therefore a stable reference frame is required that is represented by temporally stable points respectively their 3D-coordinates. In order to provide this vital condition, deformed areas within the ROI have to be identified. Stable points or areas within the ROI can

10   then be used to determine transformation parameters. These parameters allow the transformation of a successively captured epoch into the coordinate system of the reference epoch.

For the determination of geomorphometric measures TLS can be seen as an established acquisition technique (Oppikofer et al. 2008), whose measurements yield in so-called point clouds. In order to quantify geometric differences in different epochs

15   it is recommended to carry out all steps of the following processing chain:

- Viewpoint planning (Bechtold and Höfle 2016), (Wujanz et al. 2016a),
- Data acquisition at different points in time (van Veen et al. 2017),
- Transformation of individual epochs into a stable reference frame,
- Quantification of deformation (Jaboyedoff et al. 2012), (Lindenbergh and Pietrzyk 2015).

The most delicate step in this processing chain is the transformation of single epochs into a common reference coordinate system, which is also referred to as registration or matching of point clouds. Erroneous effects that occur in this step have an immediate and systematic impact onto the quantification of deformation. Thus, all conclusions that are drawn based on the generated results are falsified. The detectability of geometric changes itself is influenced by the precision of the applied laser

25   scanner (Böhler et al. 2003), (Wujanz et al. 2017) as well as the chosen algorithm for the actual quantification of deformation, e.g. (Cignoni et al. 1998), (Girardeau-Montaut et al. 2005) or (Lague et al. 2013). While random influences onto the results are inevitable, systematic biases, such as usage of uncalibrated sensors (Neitzel 2006), (Salo et al. 2008), (Al-Manasir and Lichti 2015) or the aforementioned problem when computing transformation parameters, have to be avoided at all costs that otherwise cause bogus deformation.

30   **2   Methodology**

Up to now the transformation into a reference coordinate system has been preferably carried out by usage of artificial targets and/or monumented control points. Several disadvantages can be associated to this strategy such as (i) the effort of

distributing the targets in the object space, (ii) the required survey of the targets in order to determine their 3D-coordinates in a reference system and (iii) the limited extent of the targets within the region of interest. In addition, the size of the targets must increase with growing range between scanner and ROI – for ultra-long-range scanners with ranges of up to some kilometres the targets would have to measure several metres. For built-in systems that monitor a ROI over longer periods of
5 time, rotational biases of the sensor's home position (Lichti and Licht 2006) that may occur in between epochs are usually refined by using artificial targets, see e.g. (Adams et al. 2013). For this, the assumption has to be made that the majority of targets remains geometrically stable over the course of the survey campaign – which might not be the case. A geodetic principle in the context of computing transformation parameters is that corresponding points should surround the ROI. If this is not the case, extrapolative effects occur quite likely which falsify the outcome. The reason why targets are still commonly
10 used in practice despite its numerous drawbacks can be found in the fact that all operational procedures for total stations are directly transferable to TLS. Numerous alternatives for referencing point clouds exist which can be classified as follows:

- Registration based on radiometric features,
- Direct georeferencing of terrestrial laser scanners,
- Use of geometric information.

15 Radiometrically motivated registration approaches have the advantage that they allow the access to well established techniques from the field of image matching. Since the long-term stability of radiometric information captured by TLS has not yet been studied, this strategy is not considered in the following. The successful application over short periods of time has been demonstrated by Böhm and Becker (2007).

20 Methods for direct georeferencing of TLS were initially just of scientific interest (e.g. Paffenholz et al. 2010) while lately several commercial systems emerged (Riegl 2017), (Zoller + Fröhlich 2017). A significant drawback of direct georeferencing is the extension of a scanner's error budget (Soudarissanane 2016), due to the use of additional positioning- and orientation sensors such as GNSS-equipment or electronic compasses. With increasing scanning range the impact onto the relative rotation between two point clouds also increases, which is a result of the limited accuracy of compasses.

25

A vital prerequisite for deformation measurement is the repeated acquisition of a region of interest at different points in time. Consequently the corresponding point clouds feature a large degree of overlap that can be used to transform both epochs into a common reference coordinate system. Apart from the aforementioned method of using radiometric features, geometric information can be used in form of geometric primitives, such as planes (Gielsdorf et al. 2008), (Previtali et al. 2014),
30 (technet 2017), or the overlapping region of two point clouds itself by means of the iterative closest point algorithm (ICP) as proposed by Besl & McKay (1992). A substantial advantage of the last mentioned strategy over the previously mentioned approaches can be associated to the integral use of the information density represented by point clouds. A drawback of the algorithm is its dependence to a sufficient pre-alignment of two datasets.

2.1 Motivation

In order to tackle the problem of providing a satisfactory pre-alignment, several strategies appear to be suitable such as direct georeferencing, manual pre-alignment or the application of pre-alignment algorithms for point clouds. The desired solution should not rely on additional sensors or any form of user-interaction. Hence, a two-tiered strategy was chosen that combines

5 different algorithms for the pre- (Aiger et al. 2008) as well as the fine registration (Besl & McKay 1992) of point clouds. It is well known that ICP-based approaches are sensitive against deformed regions, which yields in erroneous registrations (Wujanz 2012). Thus, it is of vital importance to reveal deformations within the captured datasets and to exclude these areas from the computation of transformation parameters. Kromer et al. (2017) approach this issue by applying Fischler and Bolles' (1981) Random sample consensus (RANSAC) in order to reject correspondences between two epochs which either

10 arise as a consequence of deformation or false matches. Since Kromer et al. (2017) process data which has been captured by a permanent monitoring system at high temporal resolution, comparably low amounts of deformation can be expected between successive epochs. Thus, the risk of potential falsification of the outcome during the registration step by contaminated data should also be rather low, which otherwise may describe an insuperable hurdle for the parameter dependent RANSAC.

The applied solution in this article is termed Iterative Closest-Proximity-algorithm (ICProx) and is based on a segmentation of the object space and identifies deformation respectively stable areas by comparison of derived measures based on locally computed transformation parameters for individual segments (Wujanz et al. 2016b). Congruent stable areas are detected by a combinatorial approach termed the maximum subsample method (Neitzel 2004) which is, in terms of robustness against

20 outliers/deformation, more reliable and hence superior to robust estimators or RANSAC.

2.2 Identification of stable areas in unreferenced laser scans

In order to avoid falsification of the deformation measurement process and consequently the analysis of deformation, it is unavoidable to identify deformed areas within point clouds and to reject them from the computation of transformation parameters. For this task the iterative closest proximity-algorithm (ICProx) was deployed that will be briefly introduced in

25 the following. A detailed description can be found in (Wujanz et al. 2016b). The algorithm consists of three essential phases:

- Phase 1: Local alignment of datasets.
- Phase 2: Identification of both stable and deformation areas
- Phase 3: Computation of transformation parameters based on stable areas.

Phase 1 in turn contains three steps while the first one automatically carries out the pre-alignment of data. Therefore the

30 4PCS-algorithm as proposed by Aiger et al. (2008) is used. In the second step, the coarsely aligned data is segmented in cubes of equal size that are also referred to as octree cells (Samet 2006 p. 211 ff). Subsequently the ICP-algorithm is locally applied within each octree cell. Therefore so-called candidates are sampled in each cell of the reference coordinate system.

For all candidates in the reference point cloud, the ICP will determine corresponding points in the second dataset. Thus every cell receives an individual set of transformation parameters based on the established correspondences. The general idea behind this approach is to locally increase or decrease the portion of deformation in order to identify geometric changes that occurred over the course of time by means of suitable criteria in Phase 2.

Therefore the individual sets of transformation parameters are applied to the corresponding centroid of a cell. Subsequently distances between centroids before and after the application of the computed transformation parameters are compared. This strategy is in general comparable to the congruence analysis of geodetic networks by usage of combinatorial methods (Neitzel 2004, p. 109 ff). Congruent and thus stable regions share the following characteristic which helps to identify them:

10   the distances between centroids before and after application of locally computed transformation parameters remain almost identical. Another unique feature of the algorithm is the automatic computation of uncertainty measures (ISO 2008) that are calculated for the local registration in every octree cell instead of using fixed thresholds. This information is required in order to clarify if an octree cell can be associated to a congruent set of cells or not. The general concept of this feature considers the fact that every scan of a stable object yields in a different point sampling on an object's surface (Wujanz et al.

15   2016b). In Phase 3 the ICP is again applied to both epochs. This time however only stable regions serve as input which are processed as a whole and not in segments. The result of Phase 3 is a set of transformation parameters that forms the basis for the deformation measurement.

Figure 1 depicts a test case on the left where stable areas are coloured in green while deformation is highlighted in yellow.

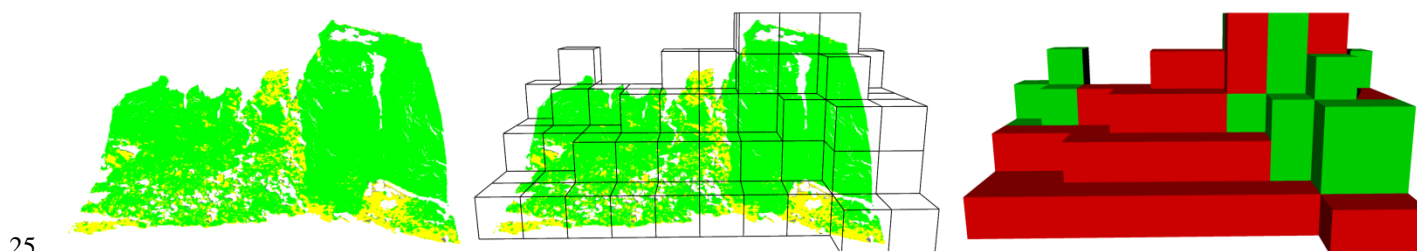
20   The aim of the ICProx-algorithm is identify these yellow areas and to reject them from the final registration. The centre of the figure illustrates the segmentation of the input data. Every cube receives an individual set of transformation parameters by local application of the ICP-algorithm. After congruence analysis all segments are merged to one continuous dataset as depicted by green areas. Red regions have been classified as deformation and are thus rejected from the registration process.

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Figure 1: (Left) Binary inspection map of a test case. (Centre) Octree structure that contains the scene. (Right) Stable (green cubes) and unstable (red cubes) octree cells after congruence analysis.

This contribution demonstrates the automatic registration of multi-temporal point clouds as a vital prerequisite for the computation of geomorphometric measures. Three regions of interest have been processed featuring different landforms such as ice glaciers, rock glaciers and a rock fall area. The study areas are described in detail in section 3 while the computed results are subject of section 4. Section 5 summarises the findings presented in this article and gives an outlook on future

5 endeavours.

3   Study areas

This study comprises three study areas in the Hohe Tauern area, Central Alps, Austria (Figure 2) monitoring three distinct geomorphic process groups:

- Hinteres Langtalkar: Rock glacier movement representing permafrost dynamics.

10 - Pasterze Glacier: Glacier retreat with subsequent proglacial processes.

- Mittlerer Burgstall: Rock fall activity.

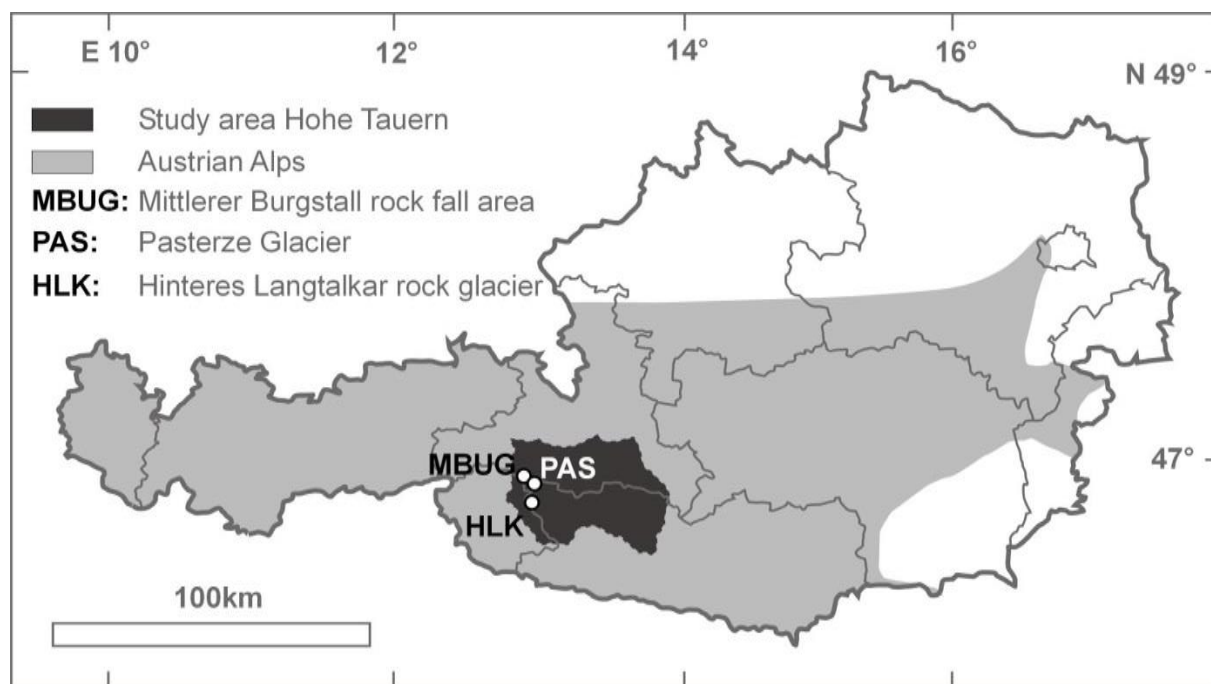



Figure 2: Test sites Pasterze Glacier, Hinteres Langtalkar rock glacier and Mitllerer Burgstall rock fall area within Austria

15   All data sets were acquired using a Riegl LMS Z620 scanner that applies a wavelength of 1550 nm and is capable to measure distances of up to 2.000 m, in applications in the study area 1500 m were achieved. A summary of the characteristics of all epochs is given in Table 1.

Table 1: Summary of all used measurements in every study area comprising date of acquisition, # of points: points per epoch, GSD: ground sampling distance

Year	Date of acquisition	# of points [Mio.]	GSD [cm] and scanning increment [deg]
Hinteres Langtalkar rock glacier			
2011	2011-08-24	1.6	6.6 – 41.8 (0.15 m -100 m, 0.040 deg)
2012	2012-09-05	7.8	4.8 – 33.1 (0.10 m -100 m, 0.029 deg)
Pasterze glacier			
2010	2010-09-13	2.5	20 – 54 cm (0.25 m-600 m, 0,029 deg)
2012	2012-09-10	3.6	16 – 42 cm (0.20 m-600 m, 0,023 deg)
Burgstall rock fall area			
2014	2014-09-08	2.3	20 – 54 cm (0.25 m-600 m,0.029 deg)
2015	2015-09-12	1.6	20 – 54 cm (0.25 m-600 m,0.029 deg)

3.1 The rock glacier Hinteres Langtalkar, Austria

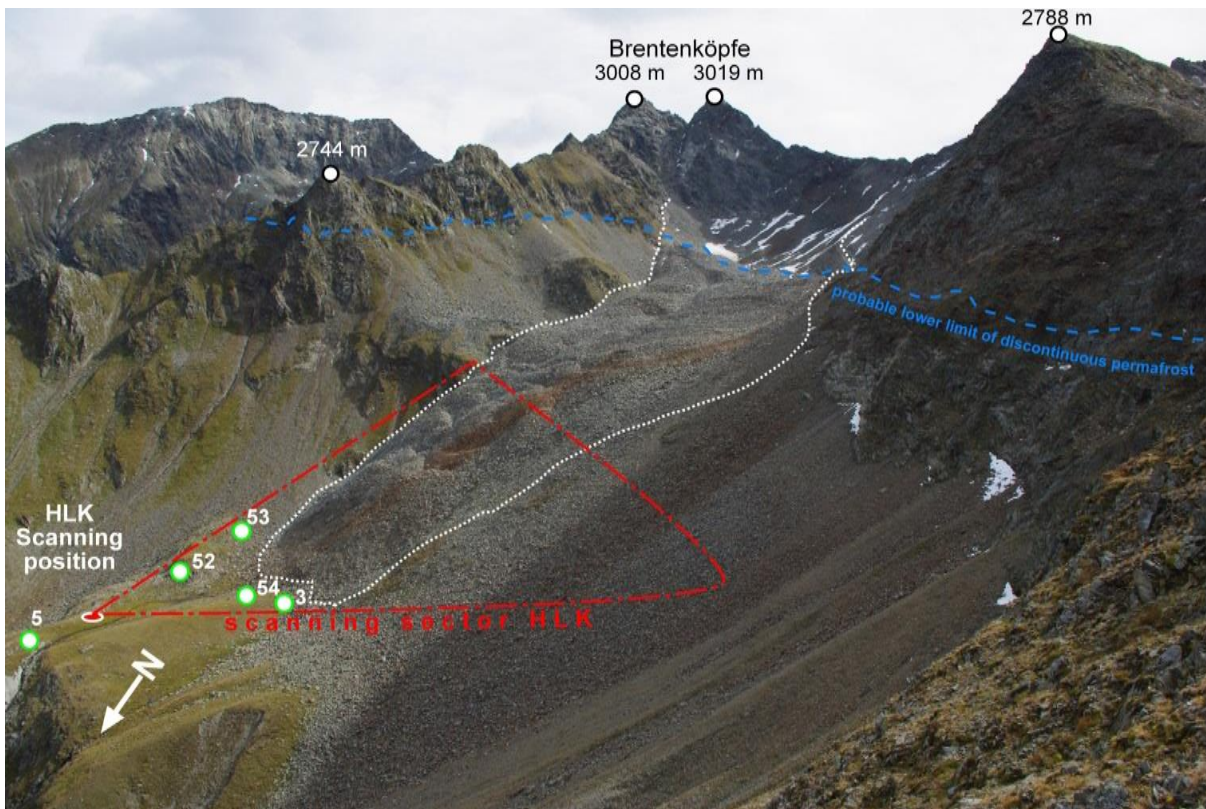
5 The rock glacier Hinteres Langtalkar (N, E) is located in the Schober Mountains in an elevation between 2455 - 2720 m a.s.l. and shows a geometry of c. 900 m in length as well as c. 350 m in width. Using TLS the rock glacier has been monitored since 2000 (Bauer et al., 2003). Extraordinary changes in surface elevation were detected around 2000 (Kaufmann and Ladstädter, 2002, Avian et al. 2005), leading to the assumption of acceleration most likely in 1994 (Avian et al., 2005). In order to geometrically quantify these surface elevation changes a geodetic network was established in 1998 which has been

10 surveyed since annually (Kienast and Kaufmann, 2004). The rock glacier itself shows distinct patterns of surface elevation changes and hence surface displacements and can therefore be categorised into a slower upper part and a rapidly moving lower part with maximum horizontal displacements of up to 2.5 m per year (Avian et al. 2009). The mean horizontal surface displacement rates vary between 0.10 – 0.18 m per year while maximum changes amount up to 3.6 m per year (Kaufmann and Ladstädter, 2009). Delaloye et al. (2008) estimate this rock glacier to be one of the currently fastest moving one in

15 Europe which underlines its importance for the scientific community.

Despite the multiple use of TLS for monitoring geomorphological processes (cp. Oppikofer et al. (2008)), TLS was rarely used for analysis of periglacial effects such as rock glacier movement patterns until the end of the last decade. First annual TLS observations were undertaken by Avian et al (2005, 2009), or Bodin et al., 2008. Data acquisition was carried out from

20 one of two known scanning positions (HLK, 2454,709 m a.s.l. on the left in Figure 3). The lowest and largest ground sampling distance (GSD) for the 2011 campaign summed up to 6.6 cm respectively 41.8 cm while the according measures for the survey of 2013 are 4.8 cm and 33.1 cm (Tab.1). These GSD values were computed based on the closest and furthest distance between scanner and scene. In Figure 3 all applied registration points are highlighted by white circles, the scanner position by a red circle. It can notably be seen that the distribution is rather unfavourable from a geodetic point of view as the

active zone of the glacier, that is highlighted by a white dashed line, is not surrounded by registration points. This may lead to extrapolative respectively leverage effects. The scanning sector is marked by red dashed lines.

5   **Figure 3: Scanner configuration at Hinteres Langtalkar rock glacier. The scanning position HLK is indicated by a red circle, the rock glacier outline by a white dashed line. White circles highlight registration points (image, 28.08.2012).**

3.2   The Pasterze glacier, Austria

The Pasterze glacier (N 47°04', E 12°44') is located in the Hohe Tauern range in the Central part Alps of the Austrian Alps.
10   The geometric behaviour of this glacier is of scientific interest since 1878 (Wakonigg and & Lieb, 1996), which makes Pasterze glacier to one of the longest documented glaciers in the world. The glacier reached its latest maximum around 1852/56 (Little Ice Age, exact date not known, Nicolussi and Patzelt, 2000) with a maximum area of 26.5 km². Since then the glacier constantly lost substance leading to a recent area of appr. 17.3 km² in 2009 (Kaufmann et al., 2015). Accordingly since the end of the Little Ice Age period Pasterze glacier lost approximately 35% of its area and about 60% of its volume.
15   The current retreat of the terminus area of Pasterze glacier using TLS has been carried out since 2001 covering an area of appr. 0.9 km².On average the annual elevation change within the debris covered part sums up to approximately 3.7 m while

6.3 m were detected in the clean ice section between 2011 and 2012. A comprehensive view on the geomorphic consequences of the fast glacier retreat can be found in Avian et al. (2017). The scanning configuration at the Pasterze Glacier terminus area can be seen in Figure 4.

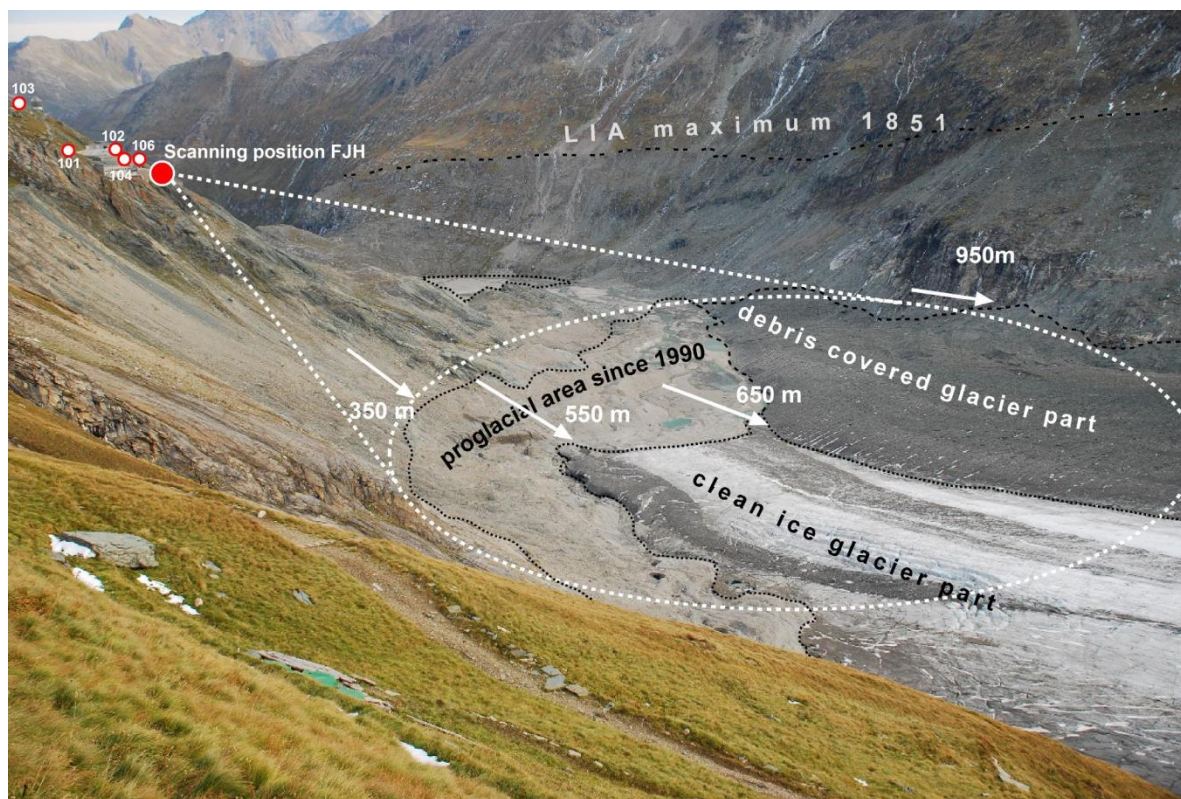



Figure 4: Scanning configuration at Pasterze Glacier terminus area from the scanning position Franz-Josefs-Höhe (FJH). Dots near the scanning position represent the location of the reference points. Numbers near the arrows indicate the distance from the scanning position. (Image 18.09.2008).

The view at scanning position Franz Josefs Höhe (FJH, 2362, 621 m a.s.l.) is excellent for geomorphic applications whereas

10 the availability of stable areas lead to an unfavourable distribution of reference points that do not surround the ROI. As a consequence the computed transformation parameters can only be seen as an approximate solution as they are very likely to be subject of extrapolative effects and hence have to be refined by e.g. surface based registration. The necessity of revealing deformed regions prior to performing co-registration has already been extensively stated so that the ICProx-algorithm is deployed in the following to solve this task for a comparison among the datasets from 2010 and 2012. A summary of quality

15 parameters of TLS measurements at Pasterze Glacier terminus area is given in Avian et al. (2017).

3.3 The Mittlerer Burgstall rock fall area, Austria

The rock fall area Burgstall comprises the two former nunatak mountains of Mittlerer and Hoher Burgstall which were encompassed from Pasterze glacier until the beginning of the 20. century. The Mittlerer Burgstall (N 47° 06´and E 12° 42; 2933 m a.s.l.) shows a distinct NW-SE trending ridge forming two rock faces which were exposed by glacier shrinkage in the last 100 years. In early summer 2007 a large rock fall at the top of ridge released c. 430.000 m³ (V. Kaufmann, Graz University of Technology, personal comment) from a detachment area of c. 3100 m² transporting debris to both sides of the ridge. Results from permafrost distribution modelling and ground temperature monitoring indicate (Kellerer-Pirklbauer et al., 2012) that the detachment area is located within or near the lower limit of permafrost. Consequently in 2010 a geodetic network consisting of one main scanning position (BUG, 2794.958 m a.s.l.) and six reference points was established. From every reference point each point can be viewed that several scanning positions can be chosen to vary scanning geometry. The scanning configuration at scanning position Burgstall (BUG) can be seen in Figure 5.

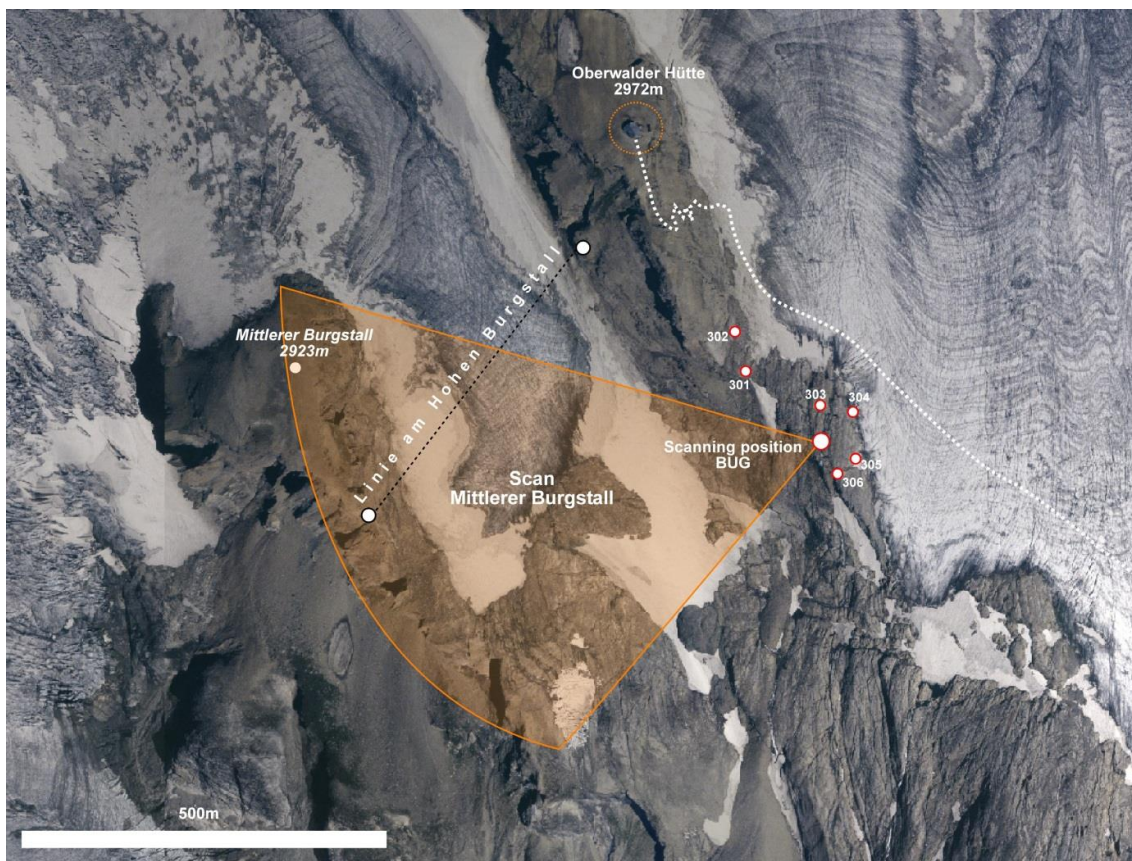



Figure 5: Configuration at scanning position Burgstall ("Standpunkt BUG"). BUG is serving two objects, Mittlerer Burgstall and Hoher Burgstall rock fall area. Codes: 301 – 306: reference points; "Linie am Hohen Burgstall" cross section of traditional glacier surveying. Ortho image Nationalpark Hohe Tauern 1998 (© 2006)

4   Results

Every algorithm is sensitive to chosen settings as well as to the individual characteristics of the given data. In this section the three aforementioned regions of interest have been analysed by the ICProx-algorithm. The input for the ICProx-algorithm consisted of two unreferenced point clouds per ROI given in different local scanner coordinate systems. Pre-alignment of the

5   epochs was carried out by the 4-PCS-algorithm while subsequently the ICProx-algorithm was deployed in order to reveal regions that were subject of deformation. The inherent challenges of the datasets presented in section 3.1 and 3.2 can be associated to their sheer extent. A somewhat different degree of difficulty can be assigned to the ROI that was subject of section 3.3. This scene is demanding due to the scattered distribution of deformation and the comparably small magnitude of geometric changes as a consequence of processes after a huge rock fall.

Since the task of the ICProx-algorithm is essentially a classification problem – identify stable or deformed regions within a point cloud – the generated outcome can be numerically evaluated in form of confusion matrices while a visual inspection is carried out by so-called confusion maps. The reference data for the evaluation process is established by point clouds where deformed areas were manually removed. Based on this data a point cloud was tinted in two colours representing the two

15   classes of interest, namely deformation and stable region. The comparison between reference and the results computed by the ICProx-algorithm finally yields in a confusion matrix respectively confusion maps as depicted for instance on the left of Figure 6. The following colour scheme will be used throughout this contribution and contains four elements, namely:

- Blue: Area is stable, detected as stable (true positive).
- Cyan: Area is stable, detected as unstable (false negative).
20 - Red: Area is unstable, detected as stable (false positive).
- Orange: Area is unstable, detected as unstable (true negative).

Note that a perfect result would only contain the first and last colour in the list. Yet, this aim is not achievable due to the applied octree structure. In the following datasets captured at the sites described in section 3 were processed by using a prototypical implementation of the ICProx-algorithm that was integrated into the software Final Surface (GFaI 2017).

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4.1   Rock glacier Hinteres Langtalkar

After pre-alignment the entire dataset was segmented into octree cells of 30 m so that in total 661 cells emerged. In order to locally determine transformation parameters candidates were sampled with 0.75 m on average. After congruence analysis, the largest subset contains 441 cells that were considered as being stable. Based on these areas an ICP-based registration was

30   carried out leading to average residuals of 15.3 cm between the two point clouds. Afterwards the computed results were compared to the reference yielding in a confusion matrix as well as a colour coded confusion map which is depicted in Figure 6. The left part of Figure 6 shows octree cells where orange represents the rock glacier and red cells are considered as

being stable. This effect occurred predominately in the transition region between glacier tongue and surrounding areas. The reason why these cells have been falsely assigned to the stable class can be found in the very low magnitude of deformation in the transition zone of the rock glacier which has a similar scale as the uncertainty measures (ISO 2008). As a consequence deformations of this size fall below the margin of detectability. The narrow red band on the left side of the scene was falsely

5   assigned as being stable due to its limited spatial expansion.

The remaining parts of Figure 6 show inspection maps based on comparisons between the georeferenced reference dataset and their according versions which were altered in position and orientation after application of the ICProx (centre) and ICP (right). Hence, the motivation for conducting this analysis is to reveal differences between the two surface based approaches

10   against target based geo-referencing. A look at the centre part of Figure 6 shows that the majority of the scene complies within +/- 5 cm while some parts feature deviations between 5 and 14 cm. These differences may be triggered by leverage effects due to an insufficient sampling of artificial targets for geo-referencing. This suspicion substantiates if one has a look at the centre part of Figure 3 where it can be seen that all targets do not surround the glacier tongue and are only located in north-westerly direction (above the glacier tongue in the figure). On average the differences between reference and the result

15   computed based on the ICProx sum up to (i) 5.8 cm for points that are located above the reference surface and (ii) 4.5 cm for points below it while the standard deviation among the two datasets measures 5.7 cm. The right part of the figure again confirms that the ICP is not an appropriate algorithm for registration of point clouds that were subject of deformation. Again the same statistical measures were computed as before where the mean deviation among the datasets sum up to 17.4 cm for points above the reference respectively 23.9 cm for points below the reference surface. The according standard deviation of

20   all residuals is 24.2 cm. It has to be mentioned that the listed statistical measures are somewhat whitewashed due to the fact that only deviations within the specified range between +/- 50 cm were considered but not the grey coloured areas where the discrepancies exceed these limits.
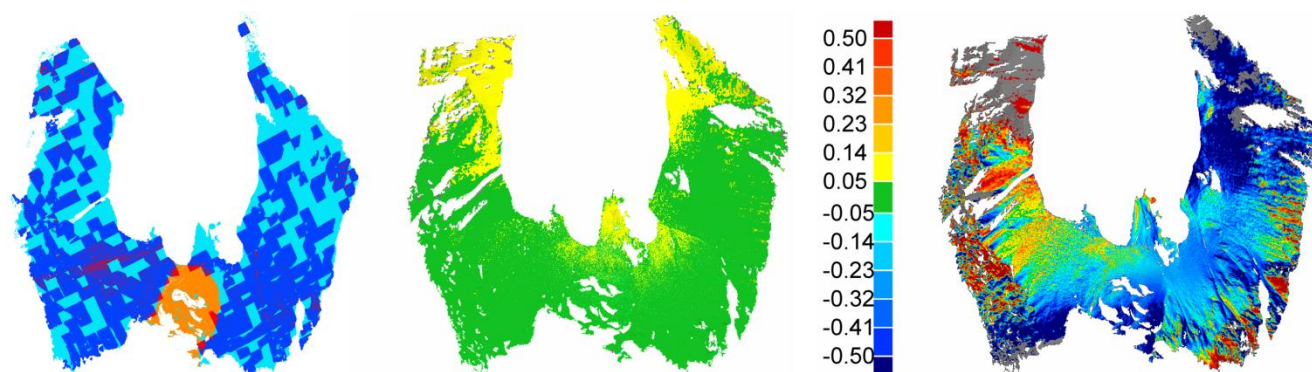
Figure 6: Colour coded confusion map (left), inspection map in metres based on a comparison to the georeferenced outcome by
25   **usage of the ICProx (centre) and the ICP (right) at Hinterer Langtalkar rock glacier.**

The numerical assessment of the generated results is depicted in Table 2 and reveals that a comparably large part of stable areas was not correctly assigned as indicated by the numbers in the cyan coloured cell. The reason for this can be found in

the weak geometric characteristics in the according octree cells. The overall classification accuracy sums up to a moderate 60% = 55% + 5% and can mostly be justified by the aforementioned geometric characteristics of the dataset. For the detection rate of stable areas a value of 60% = 55% / (55% + 37%) was achieved. This result can be rated in two ways. On the one hand it can be seen as a satisfactory result while on the other hand a substantial amount of valid information was not

5   used. The corresponding detection rate of deformation is 62.5% = 5% / (5% + 3%). After analysing the numbers of Table 2 it appears that the allegedly "simple" dataset with only 8% of deformation has turned into a challenging task for the algorithm. However, a second look at Figure 6 reveals that the stable areas surround the major area of deformation. Hence, a stable set of transformation parameters can be expected without extrapolative effects. The influence of 3% of false positive cells can be expected to be little due to the fact that the magnitude of deformation is comparably low in these parts and due to the low

10   ratio between true positive (55%) and false positive (3%) areas.

Table 2: Confusion matrix for the segmentation of the rock glacier

	Area is stable	Deformation
Classification: Area is stable	55%	3%
Classification: Deformation	37%	5%

The inspection map based on the results of the ICProx-algorithm is illustrated in Figure 7 and describes the basis for

15   deformation analysis by Geologists.
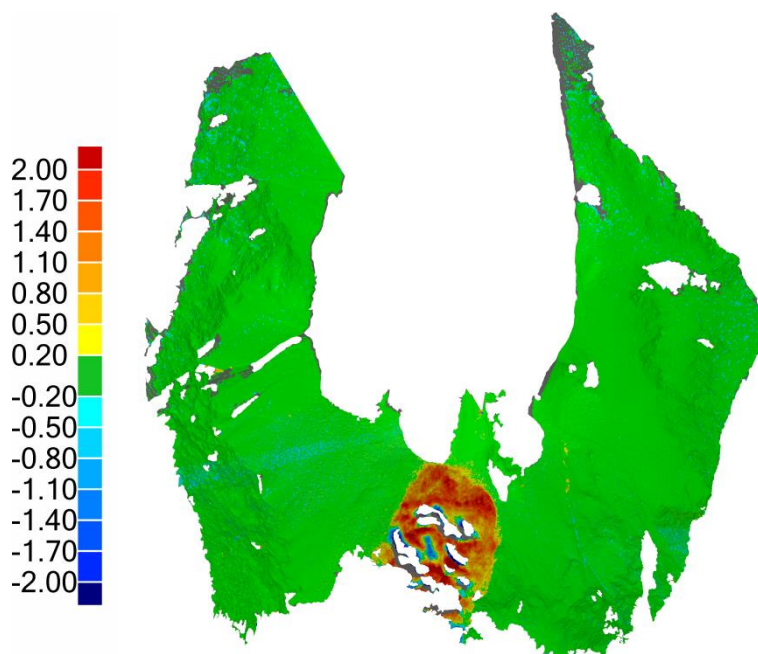
Figure 7: Inspection map in metres of the Hinterer Langtalkar rock glacier based on the ICProx.

4.2 Pasterze glacier terminus area

For this ROI it can be expected that deformation occurred predominantly directly at the glacier's tongue so that comparatively large octree cells were defined. Hence, the size of the cells was empirically set to 52.5 m leading to 923 segments. On average candidates were sampled every 6 m which lead to approximately 40700 candidates. After application

5   of the algorithm a congruent set consisting of 526 octree cells emerged that was finally used to compute the transformation parameters for deformation monitoring. The residuals after registration based on stable regions were 28.8 cm on average. An explanation for this rather large quality measure can be found in the comparably low ground sampling distance in relation to the rugged terrain. Based on expert knowledge a set of reference transformation parameters was computed featuring undoubtedly stable areas. This procedure is very time consuming, yet crucial in order to avoid falsification of the reference

10 dataset by deformation.

After generating the aforementioned reference dataset quality assessment was carried out. Figure 8 depicts the confusion map on the left which was coloured according to the colour scheme proposed at the beginning of section 2.2. In the centre of the left part the glacier tongue is highlighted in orange that caused large geometric changes. On the sides stable areas can be

15 found which have been predominantly detected by the algorithm. Cyan coloured areas can be seen predominantly on the left of the tongue. These regions have not been detected as being stable due to the fact that their local geometric characteristics did not feature notable topographic changes. Consequently the locally applied ICP converged into local minima, as the ICP requires a certain geometric contrast within the datasets. Red areas however have to be rated as being critical for the desired results. These regions can be found primarily in transition areas between glacier terminus and adjacent stable parts of the

20 point cloud. The reason why these areas were falsely classified can be explained by their magnitude of deformation – they fall below the margin of detectability. Single red spots in the otherwise stable flanks left and right of the tongue may were triggered by single sliding or mass wasting processes (Avian et al., 2017). These appearances also fall below the margin of detectability, yet not due to their magnitudes but their limited spatial extent.
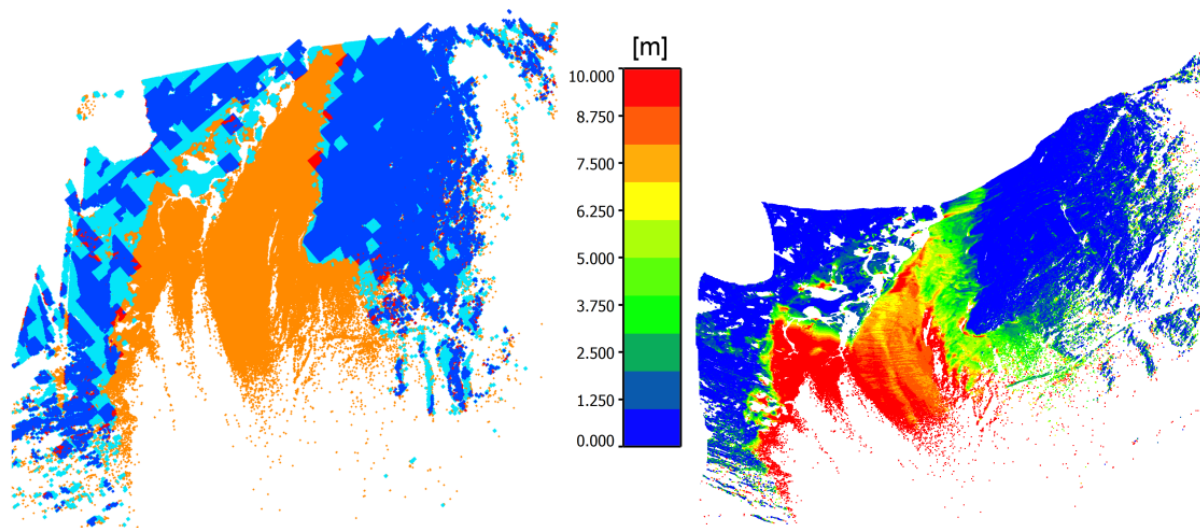
Figure 8: Confusion map based on the segmentation with the ICProx (left) and resulting inspection map (right)

The right part of Figure 8 features a point cloud that is coloured according to geometrical differences that occurred in between epochs. Large differences are notably visible at the glacier terminus area. Two regions need to be highlighted in particular. At first the orange to red coloured part has to be brought to attention vertical surface elevation changes in the magnitude of -6.5 to -10 m/a. In this part of the glacier the terminus area is only sparsely covered by debris. The second notable part is coloured in shades of yellow to green where the glacier tongue is covered by morainic debris. As a consequence the ablation rates are significantly lower compared to the clean ice part with vertical surface elevation changes in the magnitude of – 4 to -6 m/a (Kellerer-Pirklbauer et al. 2008).

The numerical interpretation of the results can be found in

Table 3. In total 19% of the dataset were falsely classified as deformation despite the fact that they actually remained geometrically stable. This effect does not provoke any substantial disadvantages, yet does simply mean that stable areas remain unused for the final registration. The portion of the critically false positive class, which is coloured in red in all tables and figures, sums up to 2%. These regions have undeniably an influence onto the registration yet can be rated as uncritical due to their comparably low magnitudes and low fraction in relation to the stable areas.

Table 3: Confusion matrix for the segmentation of the ice glacier

	Area is stable	Deformation
Classification: Area is stable	43%	2%
Classification: Deformation	19%	36%

All in all the classification has to be rated quite positively with an overall accuracy of 79% = 43% + 36%. The detection rate for stable areas sums up to 69% = 43% / (43% + 19%), the one for deformation is even 95% = 36% / (36% + 2%). Consequently the algorithm can be characterised as being "careful" – in case of doubt areas are rather classified as being deformed.

4.3   Mittlerer Burgstall rock fall area

Just as before the input for the ICProx-algorithm consisted of two unreferenced point clouds, given in different local scanner coordinate systems. In order to decompose the datasets, an octree size of 5 m in all cardinal directions was chosen yielding in 3494 cells. Candidates were sampled with an average spacing of 0.5 m so that about 120000 correspondences had to be

10   iteratively found by the ICP. After application of the ICProx-algorithm a congruent set of 2179 cells arose that was deployed to compute the desired transformation parameters for deformation monitoring. The residuals after registration based on stable regions were 17.2 cm on average. Again the explanation for this rather large quality measure can be found in the comparably low ground sampling distance in relation to the rugged terrain. As before expert knowledge was used to create a set of reference transformation parameters by manually selecting undoubtedly stable areas. Based on this information a

15   binary inspection map with the two classes stable area and deformation was created. This data served as a basis for the confusion map that is depicted in Figure 9. In contrast to the two previously shown examples, several red areas are notable in which deformed regions were falsely classified as being stable. A close look at the red speckles which are located in the major rock fall area highlighted in orange, shows that most of them contain some locally restricted blue regions. As a result, a mixture between stable and predominantly deformed areas occurs within the corresponding octree cells and hence lowers

20   the algorithm's margin of detectability. The reverse case can be seen on the right of the figure that is mostly tinted in blue where small red spots are notable. Another reason for the comparably frequent occurrence of red regions is the low magnitude of deformations.
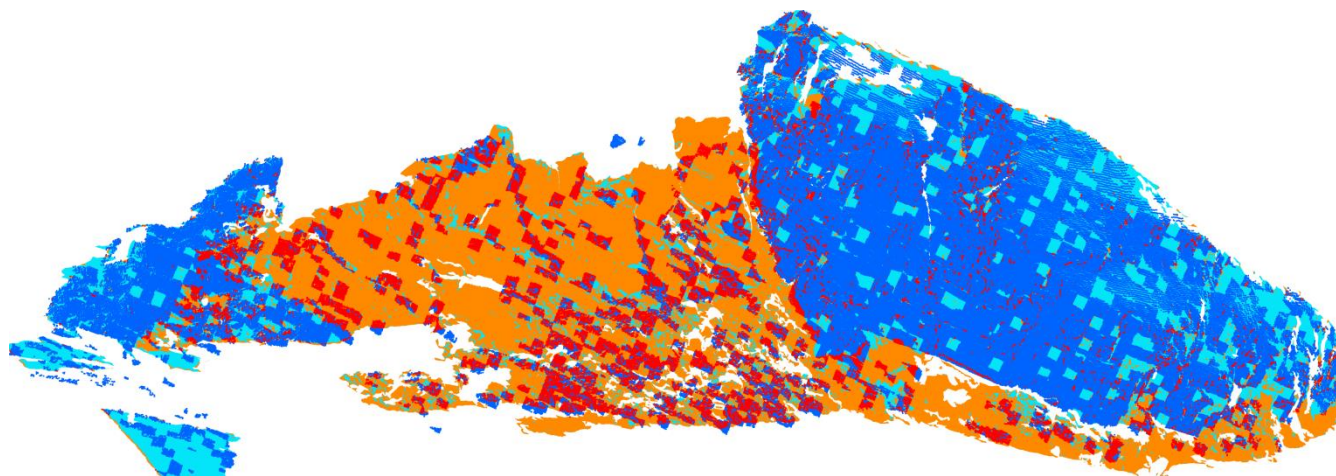
Figure 9: Confusion map of the Burgstall dataset based on the segmentation with the ICProx at Mittlerer Burgstall rock fall area.

The numerical assessment of the generated results is depicted in Table 4 and shows that a comparably large part of deformed

5 areas was not correctly assigned as indicated by the numbers in the red coloured cell. The reason for this is twofold and can be both found in the low magnitude of deformation as well as the heterogeneous occurrence of geometric changes. The overall classification accuracy sums up to 72% = 43% + 29%. For the detection rate of stable regions a value of 74% = 43% / (43% + 15%) was achieved that marks the best value in this contribution. However, the corresponding detection rate of deformation only sums up to 69% = 29% / (29% + 13%).

Table 4: Confusion matrix for the segmentation of the rock fall area

	Area is stable	Deformation
Classification: Area is stable	43%	13%
Classification: Deformation	15%	29%

Based on the reference transformation parameters an inspection map was generated that is depicted in Figure 10. Note that the chosen unit is metres. The majority of geometric changes occurred in the centre of the figure as well as on the bottom of

15 the right part. Apart from larger connected regions, local nests of rock fall can be seen especially on the upper right part of the illustration.
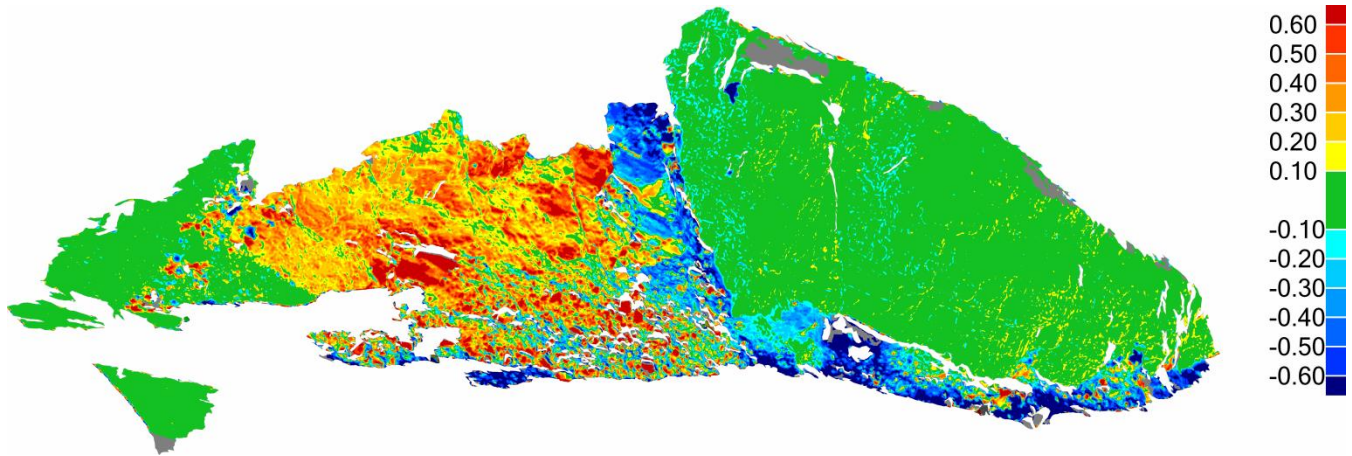
Figure 10: Inspection map in metres based on the detected stable regions – comparison between the datasets from 2014 and 2015

A look at Table 4 shows that 13% of all points were assigned to the false positive class. Hence, its influence onto the outcome is of interest in the following. Therefore the computed set of transformation parameters based on the detected stable regions was applied to the second epoch and compared to the generated reference dataset. On average the absolute deviation between reference and computed result is 7.37 cm. The corresponding inspection map in Figure 11 illustrates the differences in metres. In summary the processed dataset proved to be challenging and revealed boundaries for the ICProx-algorithm. Apart from magnitude and spatial distribution of deformation, the inherent relationship between detectability of geometric changes and local point density had a critical influence onto the outcome, that finally lead to the comparably large amount of falsely classified stable regions. Thus, future surveys in this region should be carried out at higher spatial sampling rates.
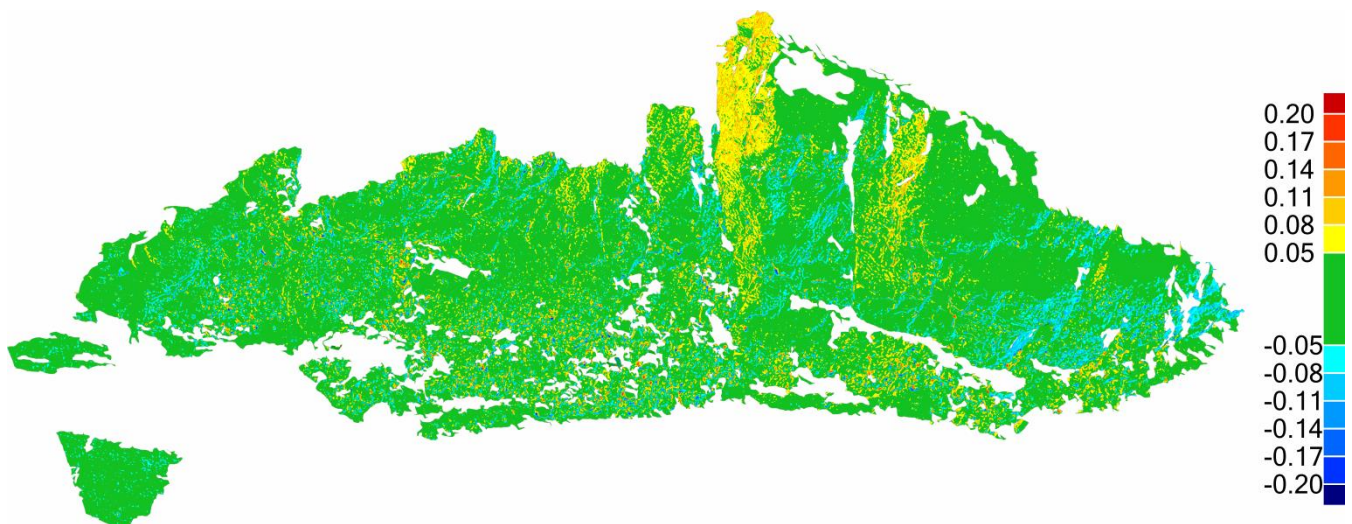
Figure 11: Comparison in metres between reference dataset and results generated based on the ICProx-algorithm

5 Summary and outlook

In this contribution an algorithm, termed the iterative closest proximity-algorithm (ICProx) was used to automatically identify and finally reject deformation from point clouds captured at different points in time. If this diagnostic step is not performed, systematic falsification of the registration is inevitable and consequently wrong conclusions are drawn during

5 analysis of the deformation pattern or computation of geomorphometric measures. Based on the detected stable regions transformation parameters were computed as a vital prerequisite for deformation monitoring. The chosen study areas featured a glacier, a rock glacier, as well as a rockfall area. The datasets itself proved to be demanding due to individual geometric characteristics, extent, relative amount and / or magnitude of deformations. All results were numerically and visually evaluated against reference datasets. The mean classification accuracy for stable regions is 68% with a standard

10 deviation of 7.1%, while the corresponding measure for deformation is 76% with a standard deviation of 17.2%. Thus, the results can be rated as quite promising. However, the importance of carefully acquiring data is particularly apparent in section 4.3 as it demonstrates the inherent link between local sampling density and potential detectability of deformation.

Further research will focus on improving some existing disadvantages of the algorithm as well as the applied assessment

15 procedure. One of them is the dependency of the outcome to the degree of decomposition of the point clouds by usage of an octree structure. Hence, an extension has to be implemented that is capable to determine an optimal octree size under consideration of the local topography. This information could also be used to dismiss certain octree cells due to insufficient geometric properties that would otherwise occur in the algorithm's numerical and visual assessment. In addition, stable cells that border on deformed regions should be split and dismissed from the registration in order to avoid falsification in such

20 transition zones as illustrated in Figure 6 and Figure 8.

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