Detection and Explanation of Spatiotemporal Patterns in Late Cenozoic Palaeoclimate Change Relevant to Earth Surface Processes

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Abstract

Detecting and explaining differences between palaeoclimates can provide valuable insights for Earth scientists investigating processes that are affected by climate change over geologic time. In this study, we describe and explain spatiotemporal patterns in palaeoclimate change that are relevant to Earth surface scientists. We apply a combination of multivariate cluster and discriminant analysis techniques to a set of high-resolution palaeoclimate simulations. The simulations were conducted with the ECHAM5 climate model and consistent set-up. A pre-industrial (PI) climate simulation serves as the control experiment, which is compared to a suite of simulations of Late Cenozoic climates, namely a Mid-Holocene (MH, ca. 6.5 ka), Last Glacial Maximum (LGM, ca. 21 ka) and Pliocene (PLIO, ca. 3 Ma) climate. For each of the study regions (Western South America, Europe, Himalaya-Tibet and South Alaska), differences in climate are subjected to geographical clustering to identify dominant modes of climate change and their spatial extent for each time slice comparison (PI–MH, PI–LGM and PI–PLIO). The selection of climate variables for the cluster analysis is made on the basis of their relevance to Earth surface processes and includes 2m air temperature, 2m air temperature amplitude, consecutive freezing days, freeze-thaw days, maximum precipitation, consecutive wet days, consecutive dry days, zonal wind speed and meridional wind speed. We then apply a two-class multivariate discriminant analysis to simulation pairs PI–MH, PI–LGM and PI–PLIO to evaluate and explain the discriminability between climates within each of the anomaly clusters. Changes in ice cover create the most distinct and stable patterns of climate change, and create the best discriminability between climates in western Patagonia. The distinct nature of European palaeoclimates is statistically explained mostly by changes in 2m air temperature (MH, LGM, PLIO), consecutive freezing days (LGM) and consecutive wet days (PLIO). These factors typically contribute 30%-50%, 10%-40% and 10%-30% respectively to climate discriminability. Finally, our results identify regions particularly prone to changes in precipitation-induced erosion and temperature-dependent physical weathering.

Keywords: Cenozoic climate, climate change, ECHAM5, Last Glacial Maximum, Mid-Holocene, Pliocene, discriminant analysis, discriminability, Earth surface processes, erosion

1. Introduction

In the study of Earth surface processes, gaining new quantitative understanding of the atmosphere's interaction with the Earth's surface through erosional processes is limited by the difficulty of establishing reliable palaeoclimatic context for erosion rate histories. Such context is particularly useful when erosion rates are calculated using techniques such as cosmogenic radionuclides and low-temperature thermochronology [e.g. Schaller et al., 2002; Bookhagen et al., 2005; Moon et al., 2011; Insel et al., 2010; Stock et al., 2009], which

integrate over timescales of 10^3-10^{6+} years. Despite recognition of the influence of climate on tectonic processes and landscape evolution through erosion [e.g. Whipple et al., 1999; Montgomery et al., 2001; Willett et al., 2006; Whipple, 2009; Deal et al., 2018], erosion rates calculated from geo- and thermochronological archives are often interpreted under the assumption of modern climate due to insufficient palaeoclimate data [e.g. Starke et al., 2017]. While proxy-based palaeoclimate reconstructions, or reconstructions of climate-controlled variables such as river discharge [e.g. Wickert, 2016], are in some cases able to provide sufficient and plausible context for specific problems, GCMs (general circulation models) offer a complementary and integrative approach to palaeoclimate reconstructions. GCMs complement proxy-based reconstructions in several ways: (1) GCMs have a global coverage [e.g. Salzmann et al., 2011; Haywood et al., 2013; Jeffrey et al., 2013] and therefore provide palaeoclimatological context in regions with sparse proxy records; (2) GCM-based palaeoclimate reconstructions allow the refinement of local proxy-based reconstructions by providing regional means and a broader climatic context; (3) GCMs are able to offer insight into atmospheric drivers of reconstructed local palaeoclimates, because they simulate atmospheric processes based on our physical understanding of the climate system. (4) GCMs allow the conduction of sensitivity experiments to investigate the relationship between climatic drivers and local observations [e.g. Takahashi and Battisti, 2007].

The Paleoclimate Modelling Intercomparison Project (PMIP) coordinates palaeoclimate modelling efforts [Kageyama et al. 2018] and provides experiment designs for the Mid-Holocene and the Last Interglacial [Otto-Bliesner et al., 2017], the last millennium [Jungclaus et al., 2017], the Last Glacial Maximum [Kageyama et al., 2017], and the Pliocene Warm Period [Haywood et al., 2016.], which is part of the Pliocene Model Intercomparison Project (PlioMIP). Despite consistent forcings, experiments carried out with different GCMs and model resolutions yield different results due to GCM specific parameterisation. [Mutz et al., 2018] conducted PMIP-style palaeoclimate experiments with the same GCM (ECHAM5) and resolution, which removes the GCM parameterisation related signal in the differences between simulated palaeoclimates. This experiment framework comprises climate simulations for the pre-industrial (PI, reference year 1850), Mid-Holocene (MH, ~ 6ka), Last Glacial Maximum (LGM, ~ 21ka) and Pliocene (PLIO, ~ 3Ma).

This study takes advantage of the Mutz et al. [2018] simulation framework and quantifies differences between simulated palaeoclimates with regard to variables relevant to Earth surface processes (e.g., rainfall characteristics, temperature derived quantities, wind speed and direction) in order to allow more refined interpretations of potential climatic drivers for changing rates in Earth surface processes. For more effective communication of our methods and results, we separate the PI control simulation from MH, LGM and PLIO in discussion by referring only to the latter three as Late Cenozoic climates. These are time periods over which reconstructed erosion rates typically integrate. Understanding how different these palaeoclimates are from a pre-industrial climate with regard to variables that potentially affect erosion rates is essential in any comprehensive and merited interpretation of erosion rates, and ultimately allows for better assessment of the influence of climatic and tectonic controls on erosion. Three questions are addressed in this study:

1. What are the spatial patterns of climate change in comparisons of pre-industrial with Late Cenozoic time periods?
2. How different were Late Cenozoic palaeoclimates to the pre-industrial climate with regard to variables relevant to Earth surface processes?
3. What constitutes these quantified differences between pre-industrial and Late Cenozoic climates?

We focus on 4 regions that are frequently investigated with regard to erosion histories: South Alaska, Western South America, Himalaya-Tibet and Europe. The first question is addressed by conducting a cluster analyses on the differences between pre-industrial and Late Cenozoic palaeoclimates, which subdivides the 4 study regions into geographical clusters governed by a distinct character of erosion relevant climate change. Whereas Mutz et al. [2018] apply a similar cluster analysis to describe the modes of climate variability in each palaeoclimate simulation, the results of this study consist of maps showing the extent of a particular mode of climate change, and thus provides an overview of climate change over time. The resulting clusters also serve as suitable masks for values used in the discriminant analyses. Questions 2 and 3 are addressed by conducting discriminant analyses on subdivisions of the 4 study areas, which are objectively pre-defined by the aforementioned cluster analyses. The results provide a quantitative assessment and explanation of differences in climate with regard to variables relevant to Earth-surface processes. The third question is answered by conducting a cluster analyses on the differences between pre-industrial and Late Cenozoic palaeoclimates, which yields the extent of regions governed by a distinct character of erosion relevant climate change.

The overarching goal is to provide the Earth surface science community with an overview and quantitative assessment and explanation of how climate changed in the Late Cenozoic with regard to variables relevant to Earth-surface processes. However, the same methods may also be applied to simulations of modern and future climates to detect and explain patterns of climate change that may result in shifts of Earth surface process regimes.

2. Data and Methods

This section describes the data, methods and processing steps. In summary, we apply a cluster analysis that identifies where surface-process-relevant aspects of climate change are likely, and a discriminant analysis that quantifies and explains these changes within the regions identified by clustering. This combination of clustering and discrimination of climate model simulation results yields 5 different sets of results (Fig. 1): Anomaly maps for a set of climate variables (in the supplemental material to this manuscript), multivariate anomaly cluster maps and anomaly cluster characterisations (section 2.2), discrimination scores and a measure of relative contribution by each climate variable to discriminability (section 2.3). While the individual components of the processing chain (Fig. 1) are based on well-established methods, the processing set-up and the particular combination of these methods is tailored to address this study's specific scientific problems.

2.1 ECHAM5 simulations

General Circulation Models (GCMs) simulate global climate based on our physical understanding of atmospheric processes and are primarily used to investigate atmospheric dynamics and contemporary climate change, but have also been applied to improve our understanding of past climates and Earth system dynamics [e.g. Kutzbach et al., 1993; Ehlers and Poulsen, 2009; Maroon et al., 2015, 2016, Mutz et al., 2016, 2018]. GCMs have become well established tools in geoscience, as is reflected by the work of the Palaeoclimate Modelling Intercomparison Project (PMIP) [Kageyama et al. 2018, Bracannot et al., 2012], which adds palaeoclimate related contributions to the Coupled Model Intercomparison Project (CMIP). Palaeoclimate studies address a range of different time scales including the last millenium [e.g. Jungclaus et al., 2010], orbital [e.g. Gong et al., 2013; Lohmann et al., 2013; Pfeiffer and Lohmann, 2016; Wei and Lohmann, 2012; Zhang et

al., 2014] and tectonic time scales [e.g. Knorr et al., 2011; Stepanek and Lohmann, 2012]. Comparisons between palaeoclimate simulations of different time periods are often complicated by use of different models and inconsistencies in model setup. Parameterisations specific to individual GCMs and differences in horizontal and vertical resolution introduce differences between simulations that are not a result of prescribed forcings.

Mutz et al. [2018] conducted a suite of GCM simulations conducted with the same GCM (ECHAM5) and resolution to circumvent the above-mentioned biases. This GCM simulation framework comprises palaeoclimate experiments for pre-industrial times (reference year 1850), the Mid-Holocene (ca. 6.5 ka), the Last Glacial Maximum (ca. 21 ka) and the Pliocene (ca. 3Ma) climates. The experiments were conducted at a spectral resolution of T159 (ca. 80km x 80km), with 31 vertical levels and an output frequency of 1 day. The simulations are based on boundary conditions from coupled AOGCM transient simulations and paleoenvironmental reconstruction initiatives such as PMIPIII [Abe-Ouchi et al 2015], GLAMAP [Sarnthein et al. 2003], CLIMAP [CLIMAP project members, 1981], PRISM [Haywood et al., 2010; Sohl et al., 2009; Dowsett et al., 2010] and BIOME6000 [Prentice et al., 2000; Harrison et al., 2001; Bigelow et al., 2003; Pickett et al., 2004]. For detailed description of model setups, we refer the reader to Mutz et. al [2018] and references therein.

The high resolution and model consistency across all of the time slice experiments in the Mutz et al. [2018] climate simulation set has not been achieved previously in the palaeoclimate modelling community. These GCM experiments therefore represent a unique, state-of-the-art simulation framework suited for investigations of changes in climate controlled processes across the Late Cenozoic. The GCM (ECHAM5) is a well-established model in the climate community, and simulated palaeoclimates are in agreement with other modelling and proxy-based reconstruction efforts. Mutz et al. [2018] use a present-day simulation to establish confidence in the model, and compare palaeoclimate estimates to compilations of proxy-based reconstructions for MH and LGM precipitation over South America and Tibet. The proxy comparisons reveal a satisfactory to good performance of the GCM, thus highlighting the suitability of GCMs in the study of palaeoclimate related problems. The latitudinal gradients and magnitude of difference in temperature and precipitation are in good agreement with results of previous palaeoclimate modelling efforts. We refer the reader to Mutz et al. [2018] for a detailed comparison to other palaeoclimate simulations and proxy-based reconstructions.

2.2 Clustering – Multivariate Anomaly Maps

This study's investigation of differences focusses on 4 regions, which (a) are of interest to the Earth-surface- and palaeo-altimetry communities, and (b) are feasible to work on given the GCM's limitations [cf. Mutz et al., 2018]: South Alaska (52°–68°N, 125°–165°W), Western South America (5°–56°S, 60°–80°W), Europe (26°–65°N, 22°W–66°E) and Hiamaya-Tibet (0°–60°N, 40°–120°E). Climate change within each of the investigated region is spatially heterogenous in both direction and magnitude [Mutz et al., 2018]. In the example of PI–LGM climate change in Europe, southern Norway experiences a strong increase in consecutive freezing days and strong decrease in 2m air temperature, while continental Europe experiences only mild cooling, but strong increases in intra-monthly 2m air temperature variability. We refer to these combinations of changing climate attributes as *modes of (climate) change* in this manuscript. Each of the regions is governed by a number of these distinct modes of climate change. These intra-regional discrepancies merit an informed and objective subdivision of each region into geographical subdomains governed by one of these modes of change prior to the

investigation of differences in climate through time. Geographical clustering allows such subdivisions on the basis of similarities in climate change at different locations within each region. It assesses climate change for each grid box, calculates its similarity to climate change in other grid boxes, and groups them accordingly. More specifically, it allows the grouping of elements (*i*), in this case climate model surface grid boxes, by the co-variability of anomalies of selected climatic element attributes. For each region, the contained elements are subjected to agglomerative hierarchical clustering, followed by k-means clustering corrections [Mutz et al., 2016; Paeth, 2004] to address the inherent shortcomings of a pure hierarchical approach. The Mahalanobis distance [e.g. Wilks, 2011] is used as a measure of similarity (of climate change) between clusters in the entire procedure. Readers are referred to Mutz et al. [2018] for a more detailed description of the procedure and aforementioned shortcomings.

The clustering is conducted on basis of the same climatic attributes as the discrimination between climates through time (section 2.3). In total, *M* (= 9) element attributes, summarised in table 1, are chosen based on (1) their relevance to Earth surface processes, (2) their adequate representation by the GCM, and (3) the feasibility of construction from GCM output: near-surface air temperature (te2m), intra-monthly near-surface air temperature amplitude (t2am), consecutive freezing days (csfd), freeze-thaw days (fthd), maximum precipitation (pmax), consecutive wet days (cswd), consecutive dry days (csdd), zonal near-surface wind speeds (u10) and meridional near-surface wind speeds (v10).

For calculation of consecutive freezing days and freeze-thaw days, a surface temperature of 0°C is taken as a threshold value. The maximum duration of a wet period, i.e. precipitation exceeding 1mm/day [Zolina et al., 2010; Zhang et al., 2011] constitutes the cswd attribute. Inversely, the maximum duration of a dry period, i.e. precipitation failing to exceed the 1mm/day threshold [Zin and Jemain, 2010] constitutes the csdd attribute. These attribute variables are constructed from each palaeoclimate (MH, LGM and PLIO) and reference simulation (PI) output. The climate attribute anomalies, which serve as a basis for the clustering, are then calculated for time slice comparisons PI−MH, PI−LGM and PI−PLIO.

Clustering requires an a priori decision on the number of clusters (*k*) or subdivisions per region. The optimal value of *k* is not known before clustering. Therefore, *k* is varied in our experiments, and the cut off point for the parameter is set once the increase in *k* no longer results in a cluster with distinct climatic character, but instead results in a weakened or strengthened character of an already existing cluster. Since optimal *k* can be expected to roughly scale with region size, the parameter is varied from 3-5 for South Alaska and from 5-8 for the larger regions. The results consist of optimal geographical subdivisions (climate clusters C_1-C_k) with distinct climatic characters, which are described by mean vectors for climate attribute anomalies. Every cluster characterising vector has a length of *M*. For each of the time slice comparisons and clusters, the discriminability and relative contribution to it by each of the *M* attribute variables is quantified in the procedure described in section 2.3.

2.3 Discrimination – Quantifying and Explaining Anomalies

The multivariate linear discriminant analysis (LMD) [e.g. Wilks, 2011] is a statistical tool that allows the investigation and explanation of differences between two or more groups with regard to multiple attribute variables. More specifically, it quantifies the discriminability of the groups and the contribution of each of the attribute variables to this discriminability. The resulting discrimination model can be applied to objectively

categorise an element with unknown group-affiliation. In this study, the time periods (PI, LGM, MH and PLIO) are used as groups, and the aforementioned climate variables relevant to Earth surface processes are chosen as group attributes. Since the focus of this study lies on the assessment of the differences between two specific time periods with regard to multiple climate variables, the problem is treated as a two-group multivariate case.

The centre piece of the analysis lies in finding a discriminant function that best separates the two groups (or paleo-climate time slices). This is carried out for each comparison, i.e. each pair of paleo-climate time slices. This discriminant function can be expressed as a linear combination of the climate attribute variables:

$$Y = \upsilon_0 + \upsilon_1 X_1 + \upsilon_2 X_2 + \ldots + \upsilon_m X_m + \ldots + \upsilon_M X_M \quad (1)$$

where Y is the discriminant function, X_m (m=1 … M) are the climate variables used to assess the differences in paleo-climates, υ_m (m=1 … M) are the discriminant coefficients associated with each variable. υ_0 is a constant (the y-intercept) that is of no relevance to the goodness of separation and will therefore no longer be mentioned. In this case, M=9 (te2m, t2am, csfd, fthd, pmax, cswd, csdd, u10, v10). Each climate variable X_m (cf. table 1) contains elements x_{mn} (n=1 … N), where N is the number of elements in each cluster. Each element (or grid box) is associated with a discriminant value y_n described by:

$$y_n = \upsilon_1 x_{1n} + \upsilon_2 x_{2n} + \ldots + \upsilon_m x_{mn} + \ldots + \upsilon_M x_{MN} \quad (2)$$

In other words, the elements y_n (n=1 … N) are projected onto the discriminant axis Y. The problem of finding a discriminant function that best separates the two groups (or paleo-climates) can therefore also be seen as the process of finding an axis on which the frequency distributions of the projected elements y_n for the two groups show the smallest overlap. Since overlap is a function of distance between groups (**D**) as well as the scatter within them (**S**), the difference between these frequency distributions is described by the distance between the two group centroids (i.e. the group means of the projected elements y_n on the discriminant axis) and the scatter within the group (i.e. sum of squared deviations from the means). This distance, i.e. the discriminant criterion Γ, is maximised in order to find the best discriminant function Y and corresponding discriminant coefficients υ_m (m=1 … M). The problem in the two-group multivariate case of this study can thus be summarised as:

$$\Gamma = \frac{scatter\ between\ groups(D)}{scatter\ within\ groups(S)} = \frac{(\overline{y_{T1}} - \overline{y_{T2}})^2}{\sum_{j=1}^{n_{T1}} (y_{T1j} - \overline{y_{T1}})^2 + \sum_{j=1}^{n_{T2}} (y_{T2j} - \overline{y_{T2}})^2} \rightarrow max \quad (3)$$

Γ is the discriminant criterion, T1 and T2 are the two groups (e.g. PI and LGM in the case of time slice comparison PI–LGM) and n_{T1} and n_{T2} are the number of elements in T1 and T2. We can express the discriminant function as a matrix calculation

$$\mathbf{y} = \boldsymbol{\upsilon_0} + \boldsymbol{\upsilon}\mathbf{X}, \quad (4)$$

where

$$y = \begin{bmatrix} y_1 \\ y_2 \\ \dots \\ y_N \end{bmatrix} \upsilon_0 = \begin{bmatrix} \upsilon_0 \\ \upsilon_0 \\ \dots \\ \upsilon_0 \end{bmatrix} \upsilon = \begin{bmatrix} \upsilon_1 \\ \upsilon_2 \\ \dots \\ \upsilon_M \end{bmatrix} X = \begin{bmatrix} X_{11} X_{12} \cdots X_{1M} \\ X_{21} X_{22} \cdots X_{2M} \\ \vdots \quad \vdots \quad \ddots \quad \vdots \\ X_{N1} X_{N2} \cdots X_{NM} \end{bmatrix}, \tag{5}$$

and solve the above-mentioned optimisation problem via partial derivatives with respect to the discriminant coefficients (**υ**) [e.g. Wilks, 2011]. The discriminant coefficients are then standardised, i.e. put in relation to the standard deviations of respective variables, to yield ω. The relative contribution (ρ) of each of the M attribute variables to discriminability is calculated as:

$$\rho_m = \frac{\omega_m}{M} \tag{6}$$

Finally, the skill of the resulting discrimination model is evaluated. For this, the association of each element to groups T1 (PI) or T2 (MH, LGM or PLIO) is forgotten, and the elements are re-categorised according to the critical discriminant values of the models [e.g. Bahrenberg et al., 1992]. The fraction of correct classifications (the **score**) is calculated and used as a measure of "goodness of separation" given by the models. The described LMD procedure is applied to each time slice comparison T1-T2 (namely PI–MH, PI–LGM and PI–PLIO) and each of the k climate anomaly clusters (C_1, C_2 , ..., C_k) in the 4 study regions. Each calculation yields two variables suitable for addressing the problems treated in this study: 1) a measure for **goodness of discriminability (score)** and 2) a measure for the **relative contribution (ρ)** of each of the M attribute variables to discriminability. A maximum score of 1 indicates perfect separation of all values, whereas a score of 0 indicates that the discrimination model has no explanatory power at all. Attribute variables associated with a ρ value of 1 are solely responsible for the discrimination, whereas those associated with a ρ value of 0 contribute nothing to the discriminability between climates.

2.4 Example Problem

In summary, the clustering of anomalies (section 2.2) reveals geographical clusters (or subdivisions in each of the study regions), in which similar climate change occurs, and describes the mode of climate change in each of these clusters. The LMD (section 2.3) then quantifies the discriminability of climates in these clusters and explains it with the climatic attribute variables. The set of results for each time slice comparisons and region therefore consists of 4 components: 1) *Anomaly (Cluster) Maps* that show the spatial extent of dominant modes of climate change; 2) *Anomaly Cluster Characterisation* that consists of mean vectors of climate change within each cluster and describe the mode of climate change experienced in the grid boxes assigned to the same cluster; 3) *Discrimination Scores* that describe the goodness of discriminability of climates within each of the anomaly clusters; and 4) *Relative Variable Contribution*, which describes the contribution of each of the climatic attribute variables to the discriminability calculated for each of the anomaly clusters. How these sets of results may be used in answering questions pertaining to climate driven Earth surface processes is demonstrated in the simplified example below.

Erosion rates were calculated, for example by means of cosmogenic nuclides, for a region in a specific area of interest circled on the map (Fig. 2). Although they are taken as modern (time T1) erosion rates, the signal

integrates over ten-thousand years and includes erosion rates at time T2. In order to find out if and how significantly erosion rates may actually have been different at time T2, Fig. 2 is consulted.

The anomaly cluster map shows the large-scale spatial patterns of changes in climatic variables relevant to Earth surface processes. Each cluster is associated with a specific mode of climate change, and all locations that fall within it experience this type of climate change. The area of interest lies in cluster C1, so all information not related to C1 and the area of interest are shaded pale grey to remove distractions. The purple-green raster plots reveal the type of climate change associated with C1 and thus with circled area of interest: T2 had a little more rainfall, a lot more (consecutive) wet days, higher temperatures, fewer freezing days and fewer (consecutive) dry days. Does the climate of T2 have a distinct enough signature to tell it apart from T1? The score is reasonably high, which reveals that the climate of T2 does indeed have a distinct signature and consequently, it suggests a good possibility for a different erosional regime or erosive intensity.

In order to assess the consequences of the mode of climate change in C1 for erosion, the circles (Fig. 2) are examined. Those explain which anomaly of which variable is responsible for the distinct signature that has been detected and described above. They indicate that ca. 60% of this "discriminability" can be explained by increases in temperature and ca. 40% by increases in consecutive wet days. Consequently, it is reasonable to assume a priori that erosive processes controlled by these parameters may be affected significantly. The specific sensitivity of conditions and processes at the Earth's surface, e.g. vegetation or critical thresholds in landscape responses, cannot be taken into consideration in these results due to the highly variable nature of it. However, the exact magnitude of those significant (and insignificant) changes are listed in supplemental table T1 to allow the reader to assess the specifics of the impact of climate on processes of the specific area and problem at hand. The above example provides the basic concepts needed to understand the remaining figures in this manuscript.

3. Results

This section contains descriptions of results from the cluster- and discriminant analyses carried out for each of the time slice comparisons and study regions. It is divided into sections for the three major study regions: 3.1 Western South America, 3.2 Europe, and 3.3 Himalaya-Tibet. For brevity, the results for South Alaska are included in the supplemental material (S10). For each subsection, regional results for time slice comparisons PI–MH, PI–LGM and PI–PLIO are presented. The figures accompanying each of those consists of four elements (cf. Section 2.4): 1) The first element to the figures are multivariate anomaly maps that were created by clustering the differences between climatic attribute variables in each of these comparisons (cf. Section 2.2) and show the resulting geographical subdivision into clusters governed by a specific mode of climate change. 2) The second element to the figures is (purple and green) raster plots showing the characterisation of those clusters. They describe the mode of change, i.e. the magnitude and direction of changes observed on average in each of the clusters. For visual clarity, the magnitude of change is scaled by the maximum absolute difference in each region and time slice comparison. Numerical values are listed in tabular format in the supplementary material to this manuscript (T1). 3) The third (grey) element consists of the scores for each cluster. These are based on a discriminant analysis carried out for the cluster (cf. Section 2.3) and describe the goodness of discriminability of the palaeoclimates in comparison to the PI control simulation. 4) The fourth and final element to the figures consists of a measure for the relative contribution (ρ) of each of the 9 climate variables (table 1) to the overall discriminability between two time slices in each geographical cluster. This element is displayed as a layer of

circles on top of element 2 and reveals how much an anomaly of a specific variable (represented by a specific shade of purple or green) contributes to the discriminability between PI and a palaeoclimate in each of the clusters. Larger circle diameters correspond to greater contributions.

For brevity, results for these time slice comparisons PI−MH, PI−LGM and PI−PLIO are simply referred to as results for MH, LGM and PLIO respectively. The notation MH-C_i, LGM-C_i and PLIO-C_i is used to refer to results for the i^{th} geographical subdivision (or cluster) in the respective set of results. Description of changes in climate are implicitly discussed in the context of the control simulation and are therefore descriptions of deviations from the PI climate. Discussions of multivariate anomaly maps and cluster characterisation (elements 1 and 2) focus mostly on the stable and persistent patterns seen in the results, i.e. geographical clusters that are least sensitive to changes in *k* and keep their distinct character. Discussions of scores and relative contribution to discriminability (ρ) focus primarily on clusters with good discriminability. Additional information on climate change on a sub-cluster scale are given in the form of single-variable anomaly maps in the supplementary material to this manuscript (S1-S9).

3.1 Western South America

Large scale patterns and modes of climate change

In western South America, the dominant modes of change for the MH are captured in clusters MH-C_3, MH-C_4, MH-C_5 and MH-C_6 (Fig. 3 a,d). MH-C_3 covers much of central Patagonia and is characterised by decreases in consecutive freeze-thaw days and increases in consecutive freezing days and consecutive wet days. MH-C_4 is the mode of change observed in most of Argentina and the central and southern Andes. It consists of relatively small increases in consecutive wet days and freeze-thaw days. MH-C_5 covers most of the tropics in the region and is characterised by decreases in consecutive wet days and maximum precipitation, and relatively small increases in 2m air temperature and consecutive dry days. Relatively large increases in 2m air temperature and consecutive dry days, and decreases in maximum precipitation and consecutive wet days constitute the mode of change for MH-C_6, which extends over low-altitude subtropics in the region.

LGM-C_1 – LGM-C_3 share a number of characteristics. These modes of changes include decreases and 2m air temperature and increases in 2m air temperature amplitude and consecutive freezing days (Fig. 3 e). Furthermore, the region occupied by LGM-C_1 – LGM-C_3 is covered by ice in the LGM. Differences between these modes of changes include a large increase in maximum precipitation in LGM-C_1 and decreases in consecutive wet days in LGM-C_2. LGM-C_4, covering much of the subtropics in the region, is characterised by relatively little change in all of the investigated variables. LGM-C_5 covers much of eastern Argentina and experiences increases in 2m air temperature amplitude. LGM-C_6 covers much of the tropics of the region and is characterised primarily by decreases in maximum precipitation and consecutive wet days.

The dominant modes of change in the PLIO are described by PLIO-C_1, PLIO-C_2, PLIO-C_3 and PLIO-C_5 (Fig. 3 f). Covering much of eastern Argentina and parts of the central Andes, PLIO-C_1 is characterised primarily by relatively large decreases in consecutive dry days and increases in maximum precipitation and consecutive wet days. The grid boxes assigned to PLIO-C_2, extending over most of the low-altitude tropics and subtropics and the Atacama desert, experience very little change on average. For grid boxes assigned to PLIO-C_3, a decrease in maximum precipitation and consecutive wet days, and increase in consecutive dry days and meridional wind

speeds can be observed. PLIO-C_5 extends over much of the central and southern Andes and is characterised by decreases in freeze-thaw days and increases in 2m air temperature and meridional wind speeds. While PLIO-C_6 experiences some of the largest changes, it only covers the Northern and Southern Patagonian Ice Fields and coincides with the reduction of the ice cover in the PLIO simulation.

Discriminability

The discrimination scores (Fig. 3 d,e,f) are highest for the LGM and lowest for the MH, and changes in temperature, consecutive freezing days, maximum precipitation and consecutive dry days are factors that explain much of the discriminability overall. LGM-C_1 – LGM-C_3 s have the highest scores. In all 3 clusters, decreases in 2m air temperature are one of the primary contributors to the discriminability between LGM and PI climate. It explains 40%-50% of the discriminability in LGM-C_1 and 30%-40% in LGM-C_2 and LGM-C_3. With 20%-30% explained discriminability in LGM-C_1 and 30%-40% explained discriminability in LGM-C_3, increases in consecutive freezing days are a second important factor for discriminability between the climates in western Patagonia. Discrimination with PLIO and PI simulations yields the highest scores for PLIO-C_6, which covers the Patagonian Ice Fields, and PLIO-C_5. Increases in 2m air temperatures and decreases in consecutive wet days and consecutive dry days explain the discriminability in PLIO-C_6 in equal parts (20%-30%). An increase in temperatures and relatively small decrease in consecutive freezing days explain 20%-30% and 40%-50% of the discriminability in PLIO-C_5 respectively.

3.2 Europe

Large scale patterns and modes of climate change

MH-C_1 covers Spain and the region east of the Caspian sea (Fig. 4 a) and is associated with the least change in climate attribute variables. MH-C_2, covering areas in western Europe, western Russia and the Mediterranean coasts (Fig. 4 a), experiences an increase in maximum precipitation (Fig. 4 d). Ukraine, Poland, much of the Baltic Sea coast and southern Scandinavia are assigned to MH-C_3 and experience decreases in freeze-thaw days. MH-C_4 and MH-C_5 primarily cover northern Africa and are characterised by increases in 2m air temperature and maximum precipitation.

LGM-C_1–LGM-C_4 (Fig. 4 b) are all partially characterised by a decrease in 2m air temperature and freeze-thaw days, and increases in consecutive freezing days (Fig. 4 e). It should be noted that grid boxes assigned to these clusters are covered by the Fennoscandian and British Ice Sheets in the LGM simulation. LGM-C_5 extends over much of the Mediterranean region, Spain and European North Russia and characterised by increases in 2m air temperature amplitude, consecutive dry days and relatively small increases in freeze-thaw days and decreases in maximum precipitation and consecutive wet days. Most of central Europe, western Asia and North Africa is assigned to cluster LGM-C_6 and experiences the least change.

The dominant modes of change in the PLIO are captured in PLIO-C_3-PLIO-C_6 (Fig. 4 c). PLIO-C_3 is a mode of change mostly seen in parts of North Africa and characterised by increases in meridional and zonal wind speeds (Fig. 4 f). In the coastal regions north of it, very little change is seen in the PLIO (PLIO-C_4). PLIO-C_5 covers much of central Europe and experiences decreases in freeze-thaw days and 2m air temperature amplitude, and relatively small increases in 2m air temperature. European Russia and parts of Scandinavia are assigned to

PLIO-C_6 and experience increases in freeze-thaw days and 2m air temperature, and decreases in consecutive dry days and 2m air temperature amplitude. PLIO-C7 is mostly distributed along parts of the Mediterranean, Black Sea and Caspian Sea coasts and characterised by increases in consecutive dry days and 2m air temperature, and a decrease in freeze-thaw days. PLIO-C8 covers southeastern Norway and the Alps and is characterised by decreases in consecutive freezing days and 2m air temperature amplitude, and increases in 2m air temperature and freeze-thaw days.

Discriminability

In all time slice comparisons, changes in 2m air temperature explains most of the discriminability in many of the geographical clusters (Fig. 4 d,e,f). Changes in consecutive freezing days and consecutive wet days are also major contributors to discriminability in the LGM and PLIO respectively. LGM-C_1 , LGM-C_2, LGM-C_5 and LGM-C_6 are associated with the highest scores for the LGM. 20%-24% of the discriminability in the clusters can be explained by decreases in temperature, and a similar amount can be explained by increases in consecutive freezing days. Although all PLIO scores are high, the PLIO cluster in central Europe (PLIO-C_5) is associated with the highest value. The discriminability in the cluster can be explained by increases in 2m air temperature (30%-40%), increases in consecutive wet days (20%-30%), decreases in consecutive dry days (10-20%) and decreases in temperature amplitude (10-20%). Discriminability in the high-altitude cluster (PLIO-C_8) can be explained by increases in consecutive dry days (10-20%) and decreases in consecutive wet days (20-30%), maximum precipitation (20-30%) and temperature amplitude (10-20%).

3.3 Himalaya-Tibet

Large scale patterns and modes of climate change

The stable patterns for the MH results include MH-C_1 covering the region south of the Himalayan orogen, MH-C_2 covering central India and Southeast Asia, MH-C_4 in the region around the Caspian sea, and MH-C_5 north of the Caspian and Aral Sea (Fig. 5 a). MH-C_1 is characterised by increases in consecutive wet days and maximum precipitation and decreases in consecutive dry days and 2m air temperature (Fig. 5 d). MH-C_2 mostly experiences changes in meridional and zonal wind speeds. MH-C_4 is characterised by increases in 2m air temperature amplitude and consecutive freezing days and decreases in freeze-thaw days. MH-C_5 grid boxes are associated with relatively large increases in freeze-thaw days and smaller increases in consecutive freezing days and 2m air temperature amplitude.

LGM-C_1 mostly covers the northernmost parts of the region and the Himalayan orogen (Fig. 5 b). The changes associated with it are decreases in 2m air temperature and increases in 2m air temperature amplitude and consecutive dry days (Fig. 5 e). The modes of change described by LGM-C_2 and LGM-C_3 govern large parts of the region, including the Arabian Peninsula, Iran, central Asia, the Tibetan Plateau and Tarim Basin, Mongolia and parts of China. These regions experience relatively small decreases in temperature. Central India and eastern Southeast Asia (LGM-C_4) are associated with decreases in consecutive wet days, maximum precipitation and zonal wind speeds, and increases in consecutive dry days. Parts of Kazakhstan, southern Russia, China, southeast Asia and northern India are assigned to LGM-C_5, which is characterised by increases in consecutive dry days.

The region covered by PLIO-C_1 includes northern India along the Himalayan orogen (Fig. 5 c) and experiences decreases in consecutive dry days and increases in consecutive wet days and maximum precipitation (Fig. 5 f). PLIO-C_2 covers most of the study region and is associated with relatively little change in all climatic attributes except meridional wind speeds. Central Asia is mostly assigned to PLIO-C_3 and experiences an increase in 2m air temperature and decrease in freeze-thaw days. The north of the study region (PLIO-C_4) is characterised by decreases in 2m air temperature amplitude and consecutive freezing days and increases in 2m air temperature and freeze-thaw days. Finally, PLIO-C_5 mostly covers the high altitude locations of the Himalaya-Tibet region that are close to steep topographic gradients, including the Himalayan orogen. This cluster is associated with decreases in wind speeds, consecutive freezing days, consecutive wet days and 2m air temperature, and with increases in consecutive dry days and 2m air temperature.

Discriminability

The significance of the climate attributes in explaining the discriminability the Himalaya-Tibet clusters (Fig. 5 d,e,f) is more variable than in Europe. While changes in 2m air temperature are important in most of the MH and LGM results, there is no clear dominant factor for PLIO clusters. In the LGM, the discriminability in the high-altitude/high latitude cluster (LGM-C_1) is mostly explained by decreases in 2m air temperature (30%-40%), mild increases in consecutive freezing days (20-30%) and mild decreases in consecutive wet days (10%-20%). LGM-C_4 has the second highest discrimination score, and the discrimination in this cluster is explained by decreases in consecutive dry days (10%-20%) and increases in consecutive wet days (10%-20%), maximum precipitation (30%-40%) and 2m air temperature (20%-30%). For PLIO, the type of climate change governing the largest cluster (PLIO-C_2) causes discriminability that is primarily explained by mild decreases in consecutive freezing days (20%-30%) and mild increases in consecutive wet days (20%-30%) and 2m air temperature amplitude (10%-20%). Discriminability in the cluster occupying the region south of the Himalayan orogen (PLIO-C_1) is explained by decreases in consecutive dry days (10%-20%) and 2m air temperature (10%-20%), and increases in consecutive wet days (10%-20%), maximum precipitation (30%-40%) and consecutive freezing days (10%-20%). Cluster PLIO-C_5 is associated with a discriminability best explained by increases in consecutive dry days (30%-40%) and decreases in maximum precipitation (10%-20%) and consecutive freezing days (20%-30%).

4. Discussion

This section describes method-related features and problems, and highlights commonly occurring patterns of change, provides possible explanations for those, and discusses these changes in context of erosional processes.

The role of large scale features

For many of the LGM and some of the PLIO results, changes in 2m air temperature and/or consecutive freezing days significantly contribute to the discriminability in clusters covering mid-latitudes. LGM-C_1 – LGM-C_3 in coastal and high-altitude West Patagonia (South America), LGM-C_1 and PLIO-C_4 in the Himalaya-Tibet region are examples of this. Many of these high-latitude clusters are also characterised by large changes in 2m air temperature and 2m air temperature amplitude in the LGM and PLIO results. The preferential cooling in higher latitudes during the LGM and enhanced meridional temperature gradient [e.g. Otto-Bliesner et al., 2006;

Bracannot et al., 2007; Mutz et al., 2018] can be expected to result in more pronounced seasonality and thus higher variation in near surface temperature amplitude. Inversely, the accentuated warming in higher latitude during the (Late) Pliocene [e.g. Salzman et al., 2011; Ballantyne et al., 2010; Mutz et al., 2018] would result in the opposite. These previously studied large-scale features explain much of the characterisation of high-latitude clusters and the significant contribution of changes in temperature-related variables to regional discriminability. Associated changes in temperature variables can have decisive impacts on physical weathering due to changes in glacial and periglacial processes (see below), as well as to biotic weathering by influencing vegetation cover.

The role of glaciers and periglacial processes

Changes in temperature in high altitude regions can impact physical weathering through glacial erosion [e.g. Egholm et al., 2009, Herman et al. 2013] and periglacial processes [e.g. Hales and Roering, 2007, Andersen et al., 2015, Marshall et al., 2015]. In southernmost South America, northern Europe and South Alaska (cf. supplemental information), the high discriminability and modes of change on the multivariate anomaly maps for the LGM are primarily controlled by the glaciers covering most of the region. Furthermore, many modes of change in the study regions involve consecutive freezing days and freeze-thaw days. Changes from ice-free to ice-covered conditions, such as in LGM-C_1-LGM-C_3 in coastal and high-altitude West Patagonia (South America) and LGM-C_1-LGM-C_4 in different terrains of Europe, are associated with increases in consecutive freezing days and decreases in freeze-thaw days. The opposite is the case for some modes of changes in the PLIO. For example, PLIO-C_6 in high-altitude Patagonia (South America) is associated with changes from ice-covered to ice-free conditions, as well as with an increase in consecutive freezing days. It may therefore shift from glacier to frost-cracking dominated erosional processes. These modes of change in the PLIO mark possible transition from glacier governed processes to periglacial processes and thus increased frost-cracking as the Earth's surface spends more time in the frost-cracking window [e.g. Matsuoka, 2001; Schaller et al., 2002; Andersen et al., 2015, Marshall et al., 2015], whereas many LGM modes of change suggest the opposite. Finally, glacial pre-conditioning of a landscape can modulate the effect of precipitation on landsliding [Moon et al. 2011].

The role of precipitation characteristics

Areas that have been covered by glaciers during the LGM and experienced a post-LGM increase in maximum precipitation or consecutive wet days may be particularly prone to precipitation-triggered landslides. This is the case, for example, in the regions covered by LGM-C_2 and LGM-C_3 clusters in high-altitude Patagonia and LGM-$C1_1$ – LGM-C_3 in northern Europe. More generally, changes in storminess affect erosion through river incision and sediment transport [e.g. Whipple and Tucker, 1999; Hobley et al., 2010]. Maximum precipitation and consecutive wet days are measures of storm intensity and duration respectively, which are primary controls for runoff and relevant for erosion. In most cases, such as the results for the Himalaya-Tibet region in the LGM and PLIO, the co-variability of consecutive wet days, maximum precipitation and consecutive dry days is intuitive: changes in consecutive wet days and maximum precipitation coincide with changes in consecutive dry days in the opposite direction. Even though palaeovegetation is considered in the set-up of the GCM simulations [Mutz et al., 2018], the modulating effect of vegetation on the impact of changes in the precipitation attribute variables on erosion [e.g. Gyssels et al., 2005] cannot be taken into account here, and thus the reader is advised

to do so in their evaluation of the effect of these changes on Earth surface processes. Note also that vegetation modifies river discharge by changing evapotranspiration and infiltration, and modifies hillslope erosion by changing root characteristics [e.g. Schmid et al., 2018]. In absence of significant vegetation changes, areas such as Bhutan, Nepal, Bangladesh and parts of Northern India (MH-C_1 and PLIO-C_1), which experiences strong increases in consecutive wet days and maximum precipitation in the MH and PLIO, are likely to have experienced an increase in such precipitation-induced erosion at these times.

The role of winds

Changes in wind speed components affect aeolian erosion, transport, and deposition, as well as mean raindrop trajectories, which should also be taken into consideration in the assessment of local precipitation-induced erosion [de Lima et al, 1992]. The results of this study reveal that changes in near surface meridional and zonal wind speeds contribute little to the discriminability between climates even in regions that experience wind direction changes due to a different ice cover in Europe [e.g. Siegert and Dowdeswell, 2004], which are reproduced well by the model. Wind speeds only show a significant contribution to discriminability (20%-30%) in the subtropical latitudes of South America due to slower meridional winds in the LGM in the region. The distinctiveness in the character of atmospheric dust transport during the LGM [e.g. Andersen et al., 1998] and thus aeolian erosion may be attributed more to system response and changes in vegetation, which cannot be taken into account in this study, than to a distinctiveness of LGM wind speeds.

Comments on methodical implications

PLIO and LGM clusters are more stable than MH clusters on the multivariate anomaly maps. This stability can be attributed to the relatively large magnitude of climate change in PLIO and LGM time slices. Lower variance of MH anomalies make element attribution to anomaly clusters in the MH more sensitive to randomisation and re-categorisation procedures (section 2.2). Consequently, the nature of MH patterns can be seen as the result of climate change of lower magnitude and less distinctiveness. The most important limitation is the poor representation of precipitation amount in areas of high topography and rainfall [Meehl et al., 2007 and comparisons with ERA-interim and station-based observations not presented here]. However, the threshold for what constitutes a “wet day” or “dry day” is relatively low [Zolina et al., 2010; Zhang et al., 2011; Zin and Jemain, 2010], so that the typical overestimation of total precipitation amount by ECHAM5 in such regions has little or no effect on the attribute variables *consecutive wet days* and *consecutive dry days*, particularly when analysed in comparison to another simulation at a different time (as was done here) which helps reduce any systematic model bias towards high precipitation rates. Comparison of the palaeoclimate simulations with local, terrestrial proxy-based reconstructions in South America and South Asia [Li et al. 2017; Mutz et al. 2018] show an agreement of 54%-59% in the MH, and 65%-100% in the LGM. Erosion-relevant processes that take place on small spatial or short temporal scales, such as intra-storm variations and rainfall characteristics [e.g. Ran et al., 2012], cannot be quantified in this study due to limited model resolution, output frequency and accuracy of such estimates on that scale. The consideration of non-climatic factors, such as local topography, slope and vegetation, is beyond the scope of this study, and the reader is advised to take these into consideration in their assessment of the effect of documented climate change on Earth surface processes. Lastly, since the parameterisation in ECHAM5's land surface scheme creates deficiencies in the representation of the local

hydrology and partitioning of precipitation [e.g. Weiland et al., 2011], variables such as runoff are excluded in this study. Routing GCM output through more sophisticated hydrological models or nesting regional climate models in the GCMs would allow more merited exploration of the regional hydrology.

5. Conclusions

In this study, we quantified the differences between pre-industrial and Late Cenozoic palaeoclimates with regard to variables relevant to Earth surface processes, explained these quantified differences and identified dominant patterns and modes of palaeoclimate change. The key findings of this study are:

- Last Glacial Maximum and Mid-Pliocene climate change is more distinct and more easily quantified than climate change of the Mid-Holocene. This is reflected in the stability of geographical regions (clusters) showing the extent of regions governed by distinct modes of climate change.

- Changes in ice cover result in distinct signatures of climate change. This is reflected by 1) the creation of clusters geographically associated with ice cover changes, 2) the persistence of these clusters when the assigned number of clusters (k) is varied in the procedure, and 3) ice cover changes in South America leading to the best discriminability overall.

- In Europe, changes in 2m air temperature explain most of the discriminability between pre-industrial and all three palaeoclimates (Mid-Holocene, Last Glacial Maximum and Mid-Pliocene). Changes in consecutive freezing days and consecutive wet days are also significant contributors to climate discriminability in Last Glacial Maximum and Mid-Pliocene results respectively. Consequently, these factors lend the Late Cenozoic palaeoclimates their unique signature and should be central in assessments of changes in Earth surface processes.

- Increases in freeze-thaw days and temperature often coincide with decreases in consecutive freezing days, and vice versa. Regions governed by these modes of changes, such as western Patagonia during the Last Glacial Maximum, are prone to changes in erosional process domain from peri-glacial to glacial or vice versa.

- Increases in consecutive wet days and maximum precipitation often coincide with decreases in consecutive dry days. Regions governed by these modes of change, such as locations south of the Himalayan orogen in the Mid-Pliocene, can be expected to be particularly prone to changes in erosion induced by precipitation and storm characteristics.

We note that the methods presented in this study may also be applied to simulations of modern and future climates. Furthermore, the procedure can easily be modified to detect and explain spatiotemporal changes in climate attributes associated with different processes, since the link between climate and the impacted system is established solely on the basis of variable selection. For example, the procedure may be used for the investigation of spatiotemporal changes of climate attributes that are related to specific human adaptation

strategies in the past or future. Such method transfer is merited if (1) the underlying statistical assumptions are still satisfied and (2) the chosen variables adequately represent the relationship between climate and the investigated processes. The procedure may also be applied more broadly to any problem that requires the detection of regions (in either geographical or any n_{th} dimensional variable space) associated with specific modes of change in data, and the subsequent quantification and explanation of changes occurring within identified regions.

Data availability: All data files used in the presentation of results, as well as the model simulation results this study is based on [Mutz et al., 2018], are freely available through the data download section of our work group wiki (esdynamics.geo.uni-tuebingen.de/wiki/) as set of data products tailored to the needs of different geoscientific communities.

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Figures and Captions

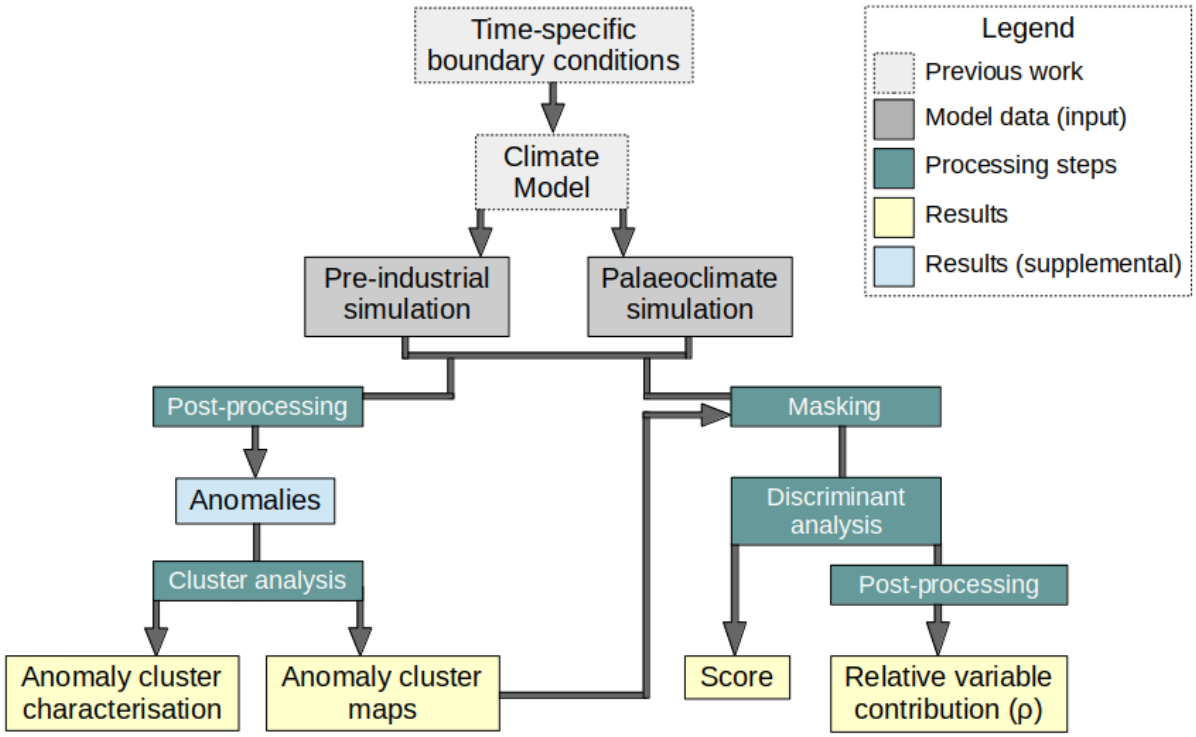

Fig. 1 Anomalies are created from pre-industrial (PI) and Late Cenozoic palaeoclimates (MH, LGM, PLIO). These are subjected to geographical clustering, which results in the identification of distinct modes of climate change (anomaly cluster characterisation) and maps showing the spatial extent of regions governed by these modes (anomaly cluster maps). These are used as geographical masks for palaeoclimate simulations. For each of these anomaly clusters, a discriminant analysis is conducted to quantify the discriminability in each cluster (score) and the relative contribution of each climatic variable to this discriminability.

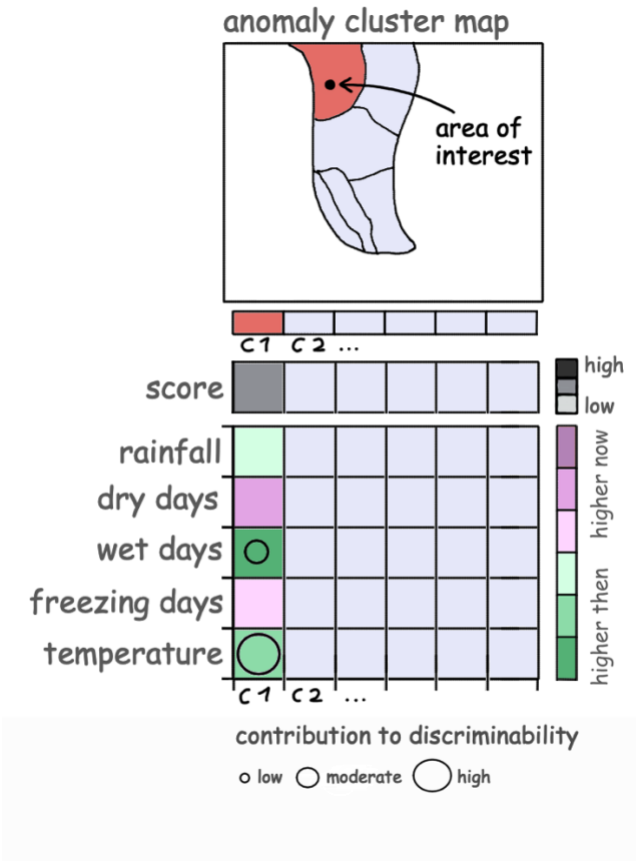

Fig. 2 Example problem: Investigating how climatic boundary conditions for erosion processes have changed in a specific region of interest involves consultation of part of the conceptualised results figure. The location of the region of interest geographically coincides with the region assigned to cluster C1. Consequently, only the results related to C1 are consulted and the rest is greyed out. The clusters were calculated from the differences in the values (delta values) of geomorphically relevant climate variables between two different times (T1 and T2) in the late Cenozoic, and thus represent a specific mode of change. The mode of change associated with the cluster of interest (C1) is revealed by the purple-green column. The relative contributions of the delta values to the overall discriminability between T1 and T2 in cluster C1, indicated by the score, is revealed by the diameters of the circle superimposed onto the delta values in the purple-green column.

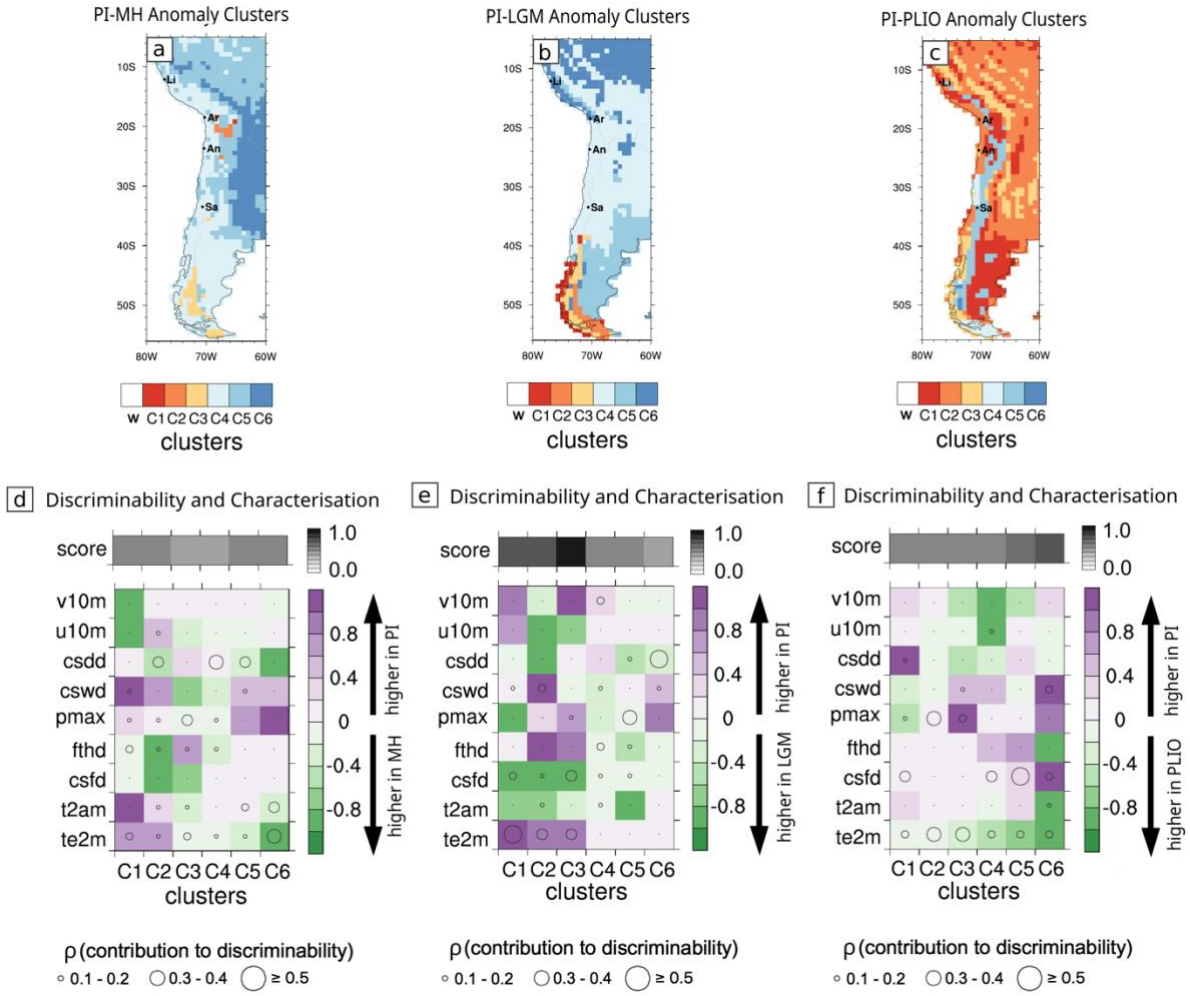

Fig. 3 The multivariate anomaly maps for time slice comparisons PI–MH(a), PI–LGM(b) and PI–PLIO(c) show the geographical coverage of clusters C_1-C_i in Western South America, which describe the spatial extend of regions characterised by similar modes of change. The corresponding modes of change (d,e and f) for each cluster are expressed as relative changes in each of the 9 investigated variables (table 1): 2m air temperature (te2m), 2m air temperature amplitude (t2am), consecutive freezing days (csfd), freeze-thaw days (fthd), maximum precipitation (pmax), consecutive wet days (cswd), consecutive dry days (csdd), zonal near surface wind speeds (u10) and meridional near surface wind speeds (v10). The score (d,e and f) expresses the goodness of discriminability between the palaeoclimate pairs PI–MH(d), PI–LGM(e) and PI–PLIO(f) in each of the anomaly clusters. The size of the circles corresponds to the relative contribution of each of the 9 climatic attribute variables to the measured discriminability in each anomaly cluster for all three time slice comparisons.

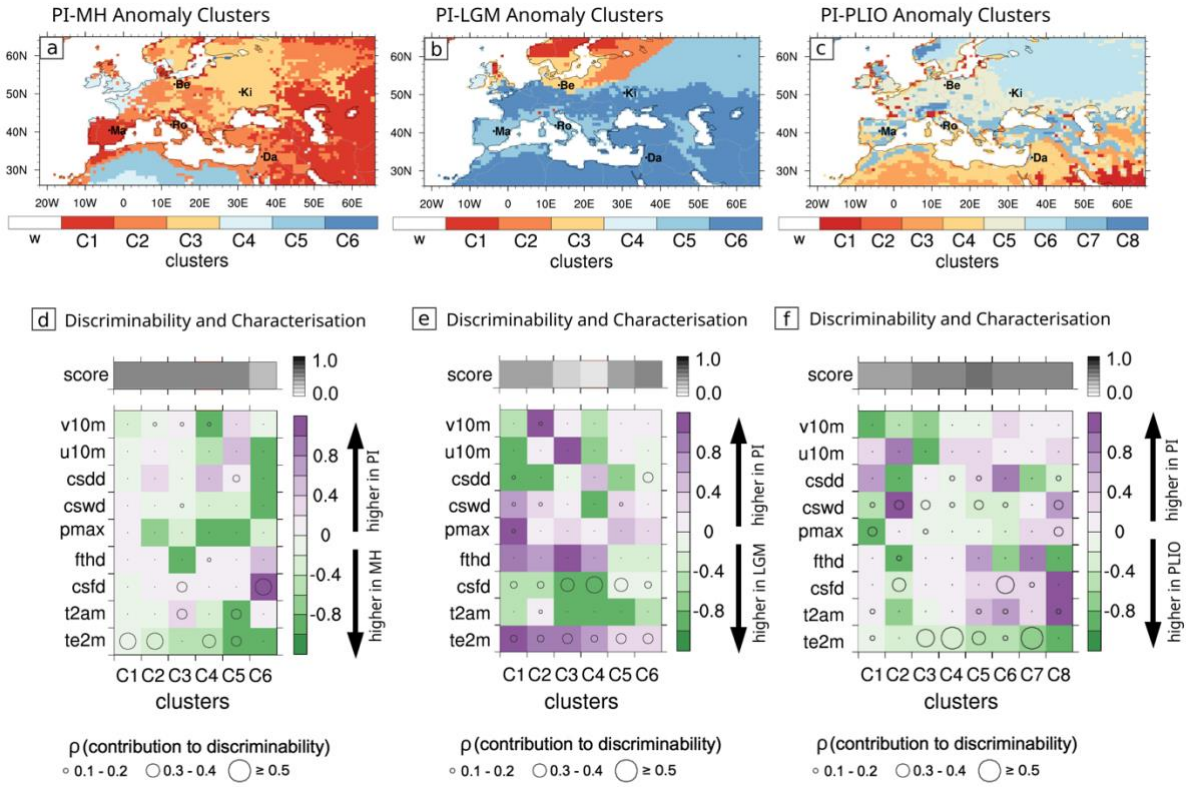

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Fig. 4 The multivariate anomaly maps for time slice comparisons PI–MH(a), PI–LGM(b) and PI–PLIO(c) show the geographical coverage of clusters C_1-C_i in Europe, which describe the spatial extent of regions characterised by similar modes of change. The corresponding modes of change (d,e and f) for each cluster are expressed as relative changes in each of the 9 investigated variables (table 1): 2m air temperature (te2m), 2m air temperature
860 amplitude (t2am), consecutive freezing days (csfd), freeze-thaw days (fthd), maximum precipitation (pmax), consecutive wet days (cswd), consecutive dry days (csdd), zonal near surface wind speeds (u10) and meridional near surface wind speeds (v10). The score (d,e and f) expresses the goodness of discriminability between the palaeoclimate pairs PI–MH(d), PI–LGM(e) and PI–PLIO(f) in each of the anomaly clusters. The size of the circles corresponds to the relative contribution of each of the 9 climatic attribute variables to the measured
865 discriminability in each anomaly cluster for all three time slice comparisons.

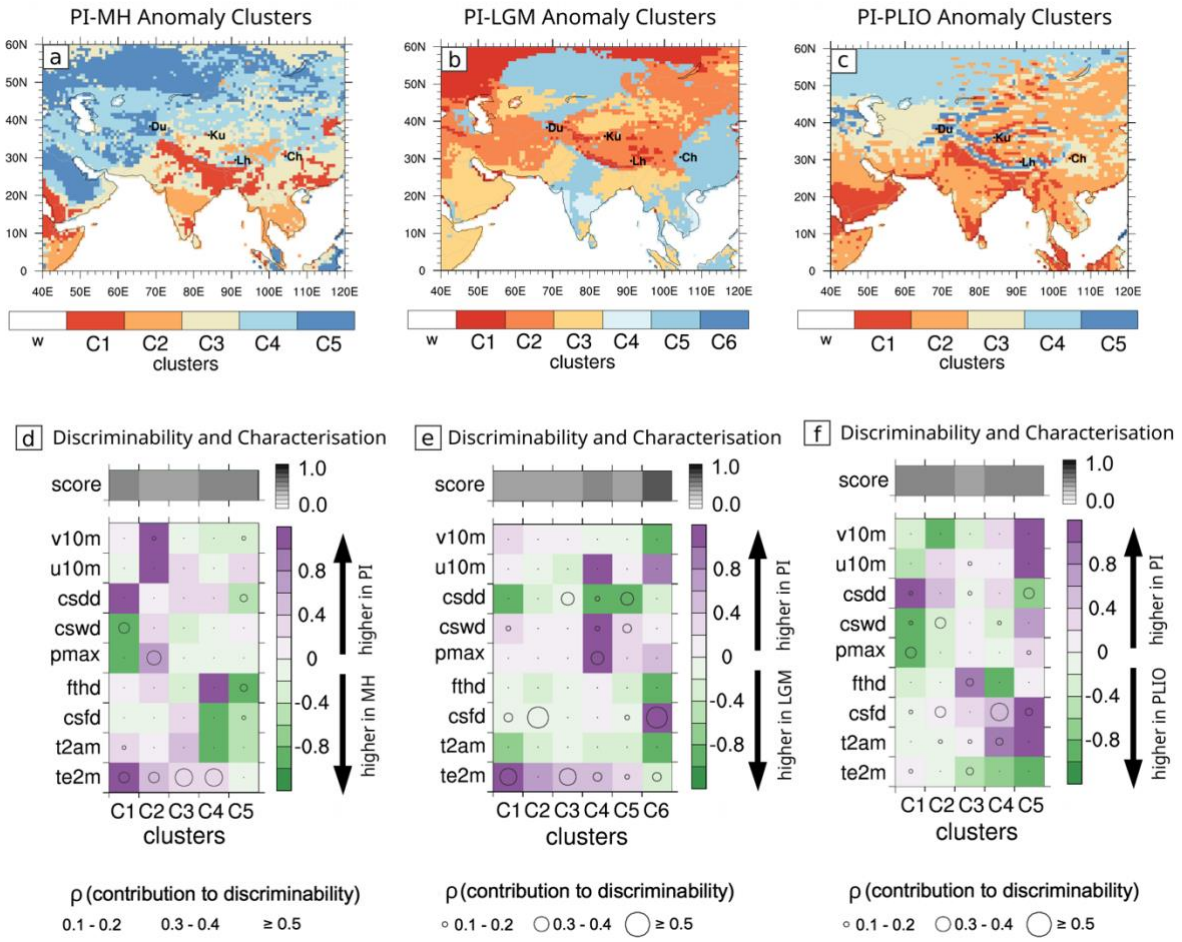

Fig. 5 The multivariate anomaly maps for time slice comparisons PI–MH(a), PI–LGM(b) and PI–PLIO(c) show the geographical coverage of clusters C_1-C_i in Himalaya-Tibet, which describe the spatial extend of regions characterised by similar mod of change. The corresponding modes of change (d,e and f) for each cluster are expressed as relative changes in each of the 9 investigated variables (table 1): 2m air temperature (te2m), 2m air temperature amplitude (t2am), consecutive freezing days (csfd), freeze-thaw days (fthd), maximum precipitation (pmax), consecutive wet days (cswd), consecutive dry days (csdd), zonal near surface wind speeds (u10) and meridional near surface wind speeds (v10). The score (d,e and f) expresses the goodness of discriminability between the palaeoclimate pairs PI–MH(d), PI–LGM(e) and PI–PLIO(f) in each of the anomaly clusters. The size of the circles corresponds to the relative contribution of each of the 9 climatic attribute variables to the measured discriminability in each anomaly cluster for all three time slice comparisons.

		Attribute Variables	
Code	Units	Explanation	Geomorphic relevance
te2m	°C	Mean annual air temperature at 2m height	(peri-)glacial processes, vegetation
t2am	°C	Maximum intra-monthly variation of 2m air temperature	(peri-)glacial processes, vegetation, frost-cracking
csfd	days	Number of consecutive days with surface temperature conditions below 0°C	(peri-)glacial processes, frost cracking
fthd	days	Number of times the 0°C threshold is crossed from day to day	(peri-)glacial processes, frost cracking
pmax	mm/d	Maximum daily precipitation value in a month	landslides, runoff, river incision, vegetation conditions
cswd	days	Number of consecutive days experiencing precipitation exceeding 1 mm/d	landslides, runoff, river incision, vegetation conditions
csdd	days	Number of consecutive days experiencing precipitation below 1 mm/d	landslides, runoff, river incision, vegetation, aeolian erosion
u10	m/s	Zonal (along-latitude) wind speeds at 10m height	aeolian erosion, transport, deposition, raindrop trajectories
v10	m/s	Zonal (along-longitude) wind speeds at 10m height	aeolian erosion, transport, deposition, raindrop trajectories

Table 1: Code, units, explanation and geomorphic relevance of each of the climate attribute variables used in the cluster- and discriminant analysis.

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