

Reply to interactive reviewer comment by reviewer 2 (anonymous)

While we disagree with most of the conceptual concerns raised by reviewer 2, we however appreciate the detailed comments and suggestions on our manuscript. We will reply separately to each of the concerns below.

This study attempts to determine the exposure ages of some large wave-transported boulders at the coast of Rabat, Morocco, using OSL rock surface exposure dating (OSL-RSED). The final exposure ages are however deemed as unreliable (i.e. imprecise and inaccurate) because of large data scatter, resulting in significant fitting uncertainties, and underestimated due to the erosion of boulder surfaces. This is altogether not very surprising, given that neither the selected lithology nor the chosen geomorphic settings are suitable for OSL-RSED technique.

OSL-RSED requires sensitive quartz and feldspar minerals, while the target boulders in this study are calcarenite, a type of limestone that is predominantly composed of carbonate, which does not have the required luminescent properties for OSL dating. OSL-RSED is also based on the sunlight-driven evolution of mm- to cm-scale luminescence-depth profiles beneath rock surfaces, and is thus very susceptible to the effect of erosion, down to sub-mm scales. Such erosion-sensitive profiles cannot be used to derive reliable surface exposure ages from boulders undergoing wave and bio-erosion at rates of $\sim 1 \text{ mm a}^{-1}$, as is the case in this study.

We will address the 5 major points of criticism separately after this general comment, but we feel it is necessary to reply to this specific conceptual comment on site selection already here:

We fully agree that the boulder lithology and the coastal setting used in this study do not provide circumstances that are ideal for OSL-RSED. However, our reasoning for conducting this study was not to apply OSL-RSED to a geomorphological/geological context with ideal preconditions, but to evaluate the potential of the approach for coastal boulder deposits. These deposits indeed potentially represent an important archive for coastal hazard assessment, but they often lack chronological information to be fully exploited. In the absence of alternative dating approaches (which is the case for numerous boulder fields worldwide), any (even relative) chronological information that might be provided by OSL-RSED is useful, because in many locations it is the only chronological information available. In this study we make a first attempt to evaluate the potential of the approach for coastal boulders in general (please note: this is not a dating study), and this includes to accept the challenging conditions and to document how they affect the reliability of the dating approach.

Therefore, we were completely aware of the rather difficult conditions for OSL-RSED of coastal boulders in general when we started the study, and we selected a site that (although not ideal compared to other geomorphological contexts) offered all indispensable prerequisites for the evaluation of OSL-RSED: A lithology containing quartz and feldspar, unambiguous signs of boulder overturning in their taphonomy, and age control at least for some of the boulders. Boulder sites with more appropriate lithologies for OSL-RSED typically lack clear indication of boulder movement and age control, and coastal boulders with better independent chronologies are typically composed of pure limestone that cannot be used for OSL dating. We realize that the reasoning of site selection may not have been explained explicitly enough in the original submission and will add two sentences on this in the introduction of a revised version.

Although not ideal, the properties of these boulders are not as poor as implied by the reviewer comment. Calcarenites are carbonate-dominated and/or carbonate-cemented sandstones (they are predominantly, i.e. $> 50 \%$, composed of carbonate grains). This means that they can contain up to 50 % non-carbonate grains such as quartz and feldspar. At the Rabat coast, the calcarenites generally do contain sensitive quartz and feldspar. This is shown in our study

using pure quartz and feldspar extracts, and it was already documented in other publications prior to this study, e.g. by Barton et al. (2009, Quaternary Science Reviews).

Furthermore, as to the comment on wave- and bio-erosion on boulder surfaces, we have to note that we explicitly did not sample surfaces that were affected by wave- or bio-erosion (except for one case, VAL 1, to investigate the effects of wave- or bio-erosion) under regular/typical non-storm conditions. The samples that are considered for dating are all well above the zone of wave- and bio-erosion. Erosion of their surfaces is driven by atmospheric weathering of the calcarenite, independent of wave- and bio-erosion. Since we selected apparently smooth surfaces with no clear signs of erosion, the quantification of erosion (which in retrospect is larger than expected at least for some of the surfaces) was one aim of this evaluation study.

While I appreciate the amount of effort the authors have put to overcome the challenges arising from this adverse combination of poor luminescence properties and erosion, I am afraid their manuscript, at its present form, is not rigorous enough to be considered for publication in *Esurf*. I could consider this study as a useful methodological contribution to the rapidly growing literature on OSL-RSED if the OSL methods were sound and the data were treated properly. But in my view, this is unfortunately not the case here. In the following, I give an account of both conceptual and methodological issues, which particularly seem problematic to me and try to explain how they could be dealt with differently, where possible. In my opinion, the manuscript may only be considered for publication after addressing these issues properly in a new submission.

Our study is meant as a methodological contribution, not a dating paper. We think we can address all methodological concerns in a revised version, why we think it is suitable for publication. In the following, we will address the five main points, on which the criticism is based on.

Geomorphology and process/hazard information:

The application of OSL-RSED to coastal boulders as is shown in Fig. 1 is oversimplified, as it does not take the effect of reworking into account. If storm surges have enough energy to detach fresh boulders from bedrock, it is very likely that they can rework (slide and overturn) the previously detached boulders sitting loose on the beach as well. It is thus quite conceivable to imagine that some of the surfaces have undergone multiple burial and exposure events, and not only a single continuous exposure event after detachment, as is conceptualised in Fig. 1. In this environment however, the dose rates are low and the burial events are too short (because storm events have high frequency and occur on decadal timescales) to leave a record in the shape of the OSL-depth profiles. Thus, an observed OSL-depth profile measures the cumulative exposure time since the detachment event, and has no record of the subsequent storm events that might have reworked the surface. Consequently, even in the absence of complications due to e.g. erosion and poor luminescence characteristics, such profiles are not particularly useful for deriving process information in similar geomorphic settings. They cannot be used for reconstructing boulder transport histories (as the title suggests), because they do not have a memory of the burial events.

We agree that we can only date the first overturning event of each boulder and not the subsequent movements. So yes, it is right that OSL-RSED of the boulders cannot be used to reconstruct the multiple transportation events that might have moved them to their final position. This is, however, not because of problems to differentiate multiple overturning events. The boulders targeted in this study have most likely been overturned only once. All of the sampled boulders weigh several tons and have a platy shape, corresponding to FI (i.e., flatness index, Nandasena and Tanaka, 2013) values of >1 or mostly even >2 . It is documented in boulder literature that such clasts are usually overturned during storms when detached from the cliff (in this situation storm waves can attack the boulders from below, e.g. Noormets et al.

2004), but that it needs waves with much larger velocities and heights to overturn them once they rest scattered on the supratidal platform (e.g. Nandasena, 2020). The predominant transport mode for a non-cubic subaerial boulder (i.e., such as most boulders in this study, with $FI > 2$) is sliding, not rolling (Imamura et al. 2008; Nandasena and Tanaka, 2013; Liu et al., 2015). While we admit that this could be explained more explicitly in the manuscript (we will do so in a revised version), the current state of the art in boulder transport by storms clearly supports the transport model shown in Figure 1 and contradicts any biasing of our OSL-RSED data by multiple overturning events. Movement of the boulders subsequent to cliff detachment can happen and probably has happened to most of the sampled boulders. But due to the boulder's shape, mass and distance from the cliff, sliding is the most plausible transport mode.

We however admit that the present title indeed may be misleading, and we suggest "Evaluating OSL rock surface exposure dating as a novel approach for reconstructing coastal boulder movement on decadal to centennial timescales" as a new title in a revised version.

They are not good proxy for storm events either, because they only record the single event that detached them from the cliff and not any of the subsequent storm events. One could argue that subsequent events of similar or higher energy are expected to pluck fresh blocks that could also be dated in a similar manner to give a chronology for the storm events. In that scenario, one would expect to see an overall trend of longer exposure events (the so-called "transport ages" here) and thus deeper OSL profiles as one moves farther from the coast, because the storms should gradually push the older boulders inland with time. But this does not seem to be the case; at least not here. For example, according to the age control, sample VAL 6 at a distance of ~80 m from the cliff seems to be younger than sample VAL 4, which is located only ~25 m from the cliff. This presumably implies that boulder detachment is not merely driven by wave power, but is also controlled by other factors such as joint formation and orientation. This inherent geomorphic character can limit the use of OSL-RSED to derive process/hazard information from coastal boulders.

We completely disagree with this opinion, since it contradicts all research on coastal boulder records. The reviewer's argument is clearly opposed by the existing literature on coastal boulders (see e.g. the latest review by Lau and Autret, 2020 and references therein). Coastal boulders have frequently been used as an archive for long-term tsunami and storm hazard assessment (e.g. Terry et al., 2013 and references therein). Regardless of the dating approach used (mainly radiocarbon, U/Th and ESR dating), all of these studies are based on ages for the initial onshore transport of the boulders, i.e. due to detachment from the cliff/reef or due to lifting from subtidal areas to the supratidal platform (e.g. Zhao et al., 2009; Engel and May, 2012; Araoka et al., 2013; Rixhon et al., 2017). While the data presented in these publications do not allow to date each transportation event and consequently not every storm, they show that (i) this limitation is not restricted to OSL-RSED but an inherent problem of all established dating approaches applicable to coastal boulders; (ii) ages of initial onshore transport can give a good impression of the recurrence patterns of storms/tsunami if sufficient boulders are dated, particularly since with increasing age specific events cannot be discriminated chronologically anyway (the fact that the scenario described by the reviewer is not reflected by the small number of ages presented in this study does not mean that the principles behind it do not generally apply); and (iii) boulder movement is often not controlled exclusively by wave power, but it is typically the dominant factor. This means that using coastal boulder records for reconstructing the history of extreme wave events may be limited by some of your concerns, but since they are the best (and often only) archive available for the reconstruction of storm/tsunami impact over geological timescales, these limitations (which apply to all dating approaches, not only OSL-RSED) are widely accepted.

To sum up our reply to the general conceptual issues, it is particularly the potential of OSL-RSED that makes it a promising candidate for providing chronological information on non-limestone, quartz- and/or feldspar-bearing boulder deposits and to make use of the coarse clast record for reconstructing extreme event histories. The exposure dating has also the

potential to provide depositional ages, which is preferred in comparison to dating of marine organisms prone to reworking. We consider this a chance to explore the coastal coarse clast record, and this paper shall present a step forward by evaluating and testing the potential. As we have argued before, the conceptual concerns of reviewer 2 are unsubstantiated.

OSL-RSED data presentation:

I find the presentation of profile data in Figs. 4, A14-16 cluttered and obscure. The mean data points with standard errors include all the information one needs to evaluate the reliability of individual data points and the overall progress of the bleaching front in a given surface. These are also the data points that are fitted to derive either the exposure age or erosion rate. So, in my view, the presentation of individual aliquots and cores in the way it is done in Figs. 4, A14-16 does not provide any useful information and impedes a proper assessment of the quality of the data.

The fits to the profile data that are used to derive the parameter values in Table 2 are not shown. Without the fits, one cannot evaluate their goodness and the reliability of the resulting parameter values.

In order to enable a clear evaluation of the data, my suggestion is to only present the mean data points with standard errors and the fits to the mean data.

Thank you for this comment, we understand the criticism of the way the OSL signal-depth data of the individual samples is presented. Our reasoning for presenting the data the way it is done in the original submission was to show the reader the entire data set he analyses is based on. We, however, realize that this may rather distract from the important information, which are the mean values and the fit of the data. In a revised version we will therefore follow the suggestion of reviewer 2 to adjust Figures A14-16 by (i) presenting only average values for each depth, and (ii) plotting the associated fit of the data to allow evaluation of its reliability.

OSL-RSED calibration:

The data from calibration sample RAB 5-1 CAL in Fig. 5 seem to reach a plateau at ≈ 0.8 and not 1. This makes me wonder i) why this sample was normalised differently and ii) how this apparently different normalisation must have affected the calibration values derived from this sample, and hence the mean calibrated parameter values used to derive the exposure ages/erosion rates. I note that the same (mean) data presented in Fig. A18 seem to have been normalised correctly. This needs to be revised, in case the authors choose to keep this sample in a new analysis of calibration data. Please see my comment below.

Sorry for the confusion. There seems to be a mistake in the axis configuration of this sample in Figure 5, which will be corrected in a revised version. The data set used for the calibration was based on values normalized to 1.0 as it is shown in Figure A18, so the calibration results are not affected by this issue.

The data from calibration samples VAL 4-1 CAL 2 and RAB 5-1 CAL seem to be much more scattered than those from the other samples. Given the goodness (badness?) of the fits to such poor-quality data, I do not think that the parameter values derived from these samples can be deemed as reliable. It is also intriguing that although the data from these samples are much more scattered than those from e.g. sample TEM 3-1 CAL, the relative uncertainties on sample-specific σ_{maphi_0} values derived from these samples are smaller than the uncertainty on the corresponding value obtained for sample TEM 3-1 CAL.

It is absolutely right that these two samples are much more scattered than the others and we agree that individual values fitted using the data are not reliable. We therefore only used them

in combination with the two other samples with flat surfaces to fit mutual σ_{phi_0} values. It is nevertheless a good idea to test a calibration of σ_{phi_0} without these samples (as suggested below).

It is argued that the sample-specific μ values have “huge uncertainties”, and therefore site-specific values of μ have been derived instead as “a reasonable and necessary compromise”. This argument is not supported by the presented data, and is not in accordance with our understanding of μ as a physical parameter.

Firstly, the relative standard deviation (RSD) of sample-specific μ values derived from the calibration samples in Fig. 5 is ~34%, while the RSD of the corresponding σ_{phi_0} values is ~210%. So, if sample-specific μ values can be dismissed because of large uncertainties and overdispersion, how can sample-specific σ_{phi_0} values, which have even greater uncertainties and are more dispersed, be acceptable and taken as a shared parameter between the calibration samples?

Secondly, if μ is dependent on lithology and all samples come from the same calcarenite bedrock, why not sharing μ between all the samples from all the sites? There is no evidence (or at least not presented here) that bedrock lithology varies from one site to another, so I cannot really see the logic behind sharing μ between samples from individual sites, but not between all the samples.

We cannot really follow the argument in this comment. We do not use sample-specific σ_{phi_0} values for calibration. We use a mutual value for all samples (otherwise we would need individual calibration samples for each targeted boulder). Thus we follow the same approach as for μ , i.e. improving the reliability of fitting by sharing the same value for several samples.

While mutual σ_{phi_0} values are, according to current knowledge, a realistic assumption for boulder surfaces from the same area and with the same surface inclination, mutual μ values indeed do not reflect the heterogeneity of rocks even from the same lithological formation (e.g. Gliganic et al., 2019). This is also the case for the study site. Although the lithology is generally similar (all calcarenite) for all boulders targeted in this study, it is not completely uniform along the entire coastline. There are slight differences in granulometry and content of bioclasts. As we explain in the original manuscript version, the best way to account for expected differences in lithology would be to use a sample-specific μ value for each sample. This is, however, impeded by fitting uncertainties, which lead to unreliable sample-specific values. We therefore have to use several samples to derive a mutual μ value. While the lithology is certainly also slightly different between the boulders at each site investigated here, these differences are considered negligible. The more significant differences exist between the different study sites. To account for these rather significant lithological differences between the study sites (for each of them a sufficient number of samples is available), we decided for site-specific μ values. We realized that this is not explicitly mentioned in the original submission and will add some information in a revised version. We, however, also checked the use of a mutual μ value for all samples.

The issues mentioned above make me wonder about the robustness of the calibration approach undertaken here and the reliability of the resulting parameter values. To address these issues, I would reanalyse the calibration data by i) excluding the inferior data of samples VAL 4-1 CAL 2 and RAB 5-1 CAL, and ii) sharing μ between all samples or leaving it as a free sample-specific parameter in fitting.

According to our replies above, we reanalysed the calibration data. (1) We excluded the strongly scattered samples VAL 4-1 CAL 2 and RAB 5-1 CAL when calibrating μ and σ_{phi_0} . (2) In addition to the site-specific μ values used in the original version of the manuscript, we also checked the use of a mutual μ value of $1.39 \pm 0.15 \text{ mm}^{-1}$ for all boulders. While these

modifications change the individual ages of each boulder, the overall chronological pattern of the boulders and, thus, our main conclusions are not affected.

Erosion rate modelling:

The authors have followed a numerical approach (not “analytical” as is mentioned in line 333) to model the OSL erosion rates. But, the OSL erosion rate equation has an exact analytical solution that is already published (see Sohbaty et al., 2018). So, there is no need and no scientific justification for making guesses at the solution numerically as is done here. The parameter values derived from the calibration samples can simply be inserted in the erosion rate equation and fitted to the profiles to give erosion rates.

The approach of Lehmann et al. (2019) that we applied to our samples is indeed a numerical approach. We will correct the wording in the revised version of the manuscript.

We, however, completely disagree that the application of a numerical approach lacks scientific justification while an analytical approach exists. We are aware that an analytical solution for the quantification of erosion from OSL rock surface data was already presented by Sohbaty et al. (2018). The numerical approach of Lehmann et al. (2019) that is used in this study was later published in *Earth Surface Dynamics*, acknowledging the analytical approach but providing an alternative solution for the erosion problem. Both approaches have their advantages and there is no approach that is absolutely superior compared to the other. The analytical solution of Sohbaty et al. (2018) might be more elegant and faster, but the numerical approach chosen in this study (which is not guessing, but inferring results from our data) is able to resolve the problem in time and provides a quantification of misfits and thus uncertainties on the results.

Minor comments:

Line 17: I suggest “wave-driven” instead of “wave-emplaced”. The boulders cannot be “emplaced” by waves and “transported” at the same time.

Thank you for this suggestion. The wording will be changed to “wave-driven” in a revised version.

Lines 48-49: “...these approaches are restricted to certain boulder lithologies and time scales.”. So is OSL RSED; it is largely restricted to lithologies that “contain quartz and/or feldspar” and to timescales of “decades, centuries up to a few millennia” as is mentioned later in lines 61-62.

Thank you for this comment. We realized that we have to be more specific here. Palaeomagnetic dating still suffers from a number of intrinsic methodological limitations, and cosmogenic nuclide dating typically cannot provide sufficient resolution on Late Holocene time scales and is, therefore, of limited benefit for the vast majority of coastal boulders. We will add a sentence explaining these details in a revised version of the manuscript.

Line 63: Does the statement “...to reconstruct...tsunami frequency patterns...” imply that the tsunami events are expected to follow some sort of temporal/spatial patterns?

Yes, tsunamis typically show temporal patterns if they are generated by earthquakes. Since the 1755 Lisbon tsunami was triggered by an offshore earthquake, it is not unlikely that potential predecessors follow a certain temporal pattern that is controlled by the accumulation of seismic strain.

Lines 71-72: Consider to change “...erosion of post-transport exposed boulder surfaces...” to “erosion of boulder surfaces exposed after transportation” or something like that.

Thank you for this suggestion. The wording will be changed accordingly.

Line 79: Add “buried” before “sediment”.

Will be changed as suggested.

Line 94: What Fig. 1 is actually showing is a boulder that is detached from a wave-cut platform and overturned by waves. There is no “transportation” involved in the depicted scenario.

The relocation of the boulder from the cliff edge to the supratidal coastal platform in an overturning movement clearly involves transportation. In Figure 1 the process of overturning during transport is illustrated by showing two successive stages of boulder movement. We nevertheless propose to change the wording in a revised version to better express the fact that we always date the cliff detachment of overturned boulders and not potential transport events following afterwards (which typically take place as a sliding movement for plate-shaped boulders as selected in this study).

Line 147: I cannot see how 2-3 m-high spring tides can reach and exceed the 5-m high first ridge (as is mentioned in line 154) to flood Oulja.

While the first calcarenite ridge shows average heights of about 5 m above sea level, this barrier occasionally shows sections with lower elevations or can even be breached at river mouths. This is where water can enter the depression of the Oulja during high tides.

Lines 189-196: The preheat temperature should also be mentioned somewhere in these lines as Table A2 is in the Appendix.

Thank you for the suggestion. In a revised version, we will mention the preheat conditions (220 °C, 10 s) in this section.

Line 191: The stimulation time in Table A2 is 150 s and not 160 s.

Sorry for this mistake. This should be 160 s as stated in the main text.

Lines 197-208: I suppose the dose recovery and preheat plateau tests described in this paragraph were carried out to guide decision on the most suitable measurement protocol. In that case, this paragraph must precede the previous paragraph in which the actual measurement protocol is explained.

Thank you for the suggestion. We changed the order of arguments to clarify that these experiments were used as a basis for final protocol selection.

Line 207: The “burial ages” suddenly appear here. So far, only OSL RSED is discussed. It is also mentioned (in lines 104-105) that the buried sides of the boulders are inaccessible and “not tried in this study”. So, speaking of burial ages here is confusing to me. In fact, it is first 60 lines further down in the text (line 267) that a careful reader may find out that what here is referred to as burial age, is actually the rock formation age, calculated by dating quartz extracts from deep layers within the boulders that have never seen light after rock formation. These should not be confused by boulder surface burial ages.

Sorry for the confusion. While the ages indeed reflect the timing of sand grain burial during ridge formation, we agree that the term “burial age” may be ambiguous in this study. To differentiate rock surface burial ages (which were not determined in this study) from conventional OSL dating of the sandstone formation (which we refer to here), we will replace “burial ages” by “ages for sandstone formation” in a revised version.

Line 212: I find the use of the term “background level” inappropriate here. Background level in OSL dating is commonly referred to while discussing the stimulation curves. I suggest “plateau” instead.

We absolutely agree that the term “background” may be misleading in this context. The wording will be changed as suggested.

Lines 214-215: This sounds to be a subjective and qualitative approach towards removing the outliers, while there are various quantitative methods to identify them. One common approach that could also be used here is to remove those data points that are different than the mean by three standard deviations.

Thank you for the suggestion. Our approach was indeed somehow subjective. We therefore revised our rejection criteria as follows: (i) Entire cores were excluded, if they did not show any signs of bleaching with depth, while all other cores from the same sample did. (ii) All other data points were classified as outliers according to a deviation from the mean of more than 2 SD. The data did not change significantly. However, we will use this new data set for all analysis in a revised version.

Line 229: Not sure what is meant by “comparable preconditions for sunlight exposure”. If the scenario is as simple as shown in Fig. 1, then all the boulders must have experienced comparable conditions (i.e. detachment and overturn). But if they are likely to have been reworked (i.e. moved and turned over multiple times) then it is very difficult to imagine how they could have had comparable exposure conditions.

We explain the meaning of this term in the second part of this sentence. While all boulders used for this study have been overturned only once (see reply to main comment: the platy boulders used in this study were overturned when detached from the cliff, but moved by sliding only or not at all afterwards), sunlight exposure may also be different due to differential shielding after deposition or due to different exposure angles. We will nevertheless change “preconditions” to “conditions”, since this seems more appropriate.

Line 252: How about “target” instead of “dated”?

We appreciate this suggestion and will change as suggested in a revised version.

Lines 253-256: It is difficult for me to judge this inference by the way the data are presented in Fig. A12. The pure quartz BSL, K-rich feldspar IRSL and polymineral post-IRSL-BSL signals must be normalised and shown on the same graph to enable a direct comparison.

We do not agree with this opinion. What we want to document is: (1) Post-IRSL-BSL signals of polymineralic aliquots are significantly stronger than the IRSL signals measured on the same polymineralic aliquots. This is documented in Fig. A12a, where normalized values of both signals are compared in the same plot (as asked for by the reviewer). (2) The post-IRSL-BSL signals of polymineralic aliquots are dominated by a quartz signal with only minor influence of feldspar signals. This is documented in Fig. A12d, which shows that the IRSL stimulation used in our protocol reduces the potassium feldspar signal to 60% of its initial value (more details are given in the caption of Fig. A12).

Lines 274-275: I assume that calibration was carried out before fitting the actual data? Please present the steps in data analysis in the logical order.

We use this sentence as an introduction to the explanation, why calibration is necessary. To avoid confusion, we will change the wording to “To estimate boulder ages with OSL-RSED, measured post-IRSL-BSL signal-depth data must be fitted with the bleaching model described in Equation (1)”.

Line 279: Sohbaty et al. (2011) is the correct reference.

Thank you for the correction. The reference will be changed accordingly.

Lines 293-295: This is an interesting observation that the calibration sample TEM 3-1 CAL that is collected from an inclined surface yields a σ_{phi_0} value that is ~3 orders of magnitude larger than the corresponding values estimated for the horizontal surfaces. If this conclusion still stands after data reanalysis (see my comments above), it would be useful to report the tilt angle of the surface. At the moment, there is no data on the dependence of σ_{phi_0} on the incident angle of solar radiation in the literature.

After reanalysing the data by excluding the two calibration samples with poorly defined bleaching fronts, there is still a significant difference of one order of magnitude between the horizontal calibration samples and the inclined calibration sample (the angle of the surface is already reported in Table 1 with $\sim 25^\circ$). We agree that such an observation has not been reported and would be worth a more detailed investigation. We are, however, aware that our assumption is only based on a single sample (and a total set of 5 calibration samples even without excluding RAB 5-1 CAL and VAL 4-1 CAL2). It obviously needs a larger dataset and more controlled conditions (e.g. in a bleaching experiment) to evaluate the assumed relationship between inclination of the surface and σ_{phi_0} .

Line 307: What is meant by “inadequate” here?

We will change “inadequate” to “incorrect”.

Line 333: The approach of Lehmann et al. (2019) is numerical not analytical.

This is of course correct. Will be changed to “numerical” in a revised version.

Line 356: “observed” instead of “achieved”?

Thank you for the suggestion. The wording will be changed accordingly.

Line 365: It seems unlikely to me that “mineralogy-induced dose rate differences” can result in the observed scatter in data from such samples. Hot minerals such as zircon and K-rich feldspars are rare, if not non-existent, in calcarenite. Meyer et al. (2018) have attributed similar scatters in their data to the presence of opaque minerals and iron hydroxides, which strongly impede the penetration of light with depth. In the absence of any independent evidence, this seems more reasonable to me as an explanation here.

We agree that this argument will definitely not explain most of the observed scatter. While it may add to the observed scatter of signals (that is why we included the argument originally), it is likely of very minor importance and might involuntarily make the discussion more complicated than necessary. We therefore decided to abstain from using this argument in a revised version of the manuscript.

Lines 366-368: I am not sure I follow. How can the aliquot-to-aliquot variation in feldspar content can give rise to additional scatter in profile data? Does it mean that test dose is not adequately correcting for this possible variation? Why not? What is the evidence?

We do not have direct evidence for this argument. We however know that the IRSL stimulation of our post-IRSL-BSL protocol is removing most of the feldspar signal, but not all of it. Although we assume that the feldspar contribution is insignificant based on our test measurements, it must be expected that the contribution of feldspar signals to the post-IRSL-BSL signal will be slightly different for polymineralic aliquots with different percentages of feldspar. Since this

potential source of scatter will again explain (if at all) only a very minor part of the observed scatter, we again decided to abstain from using this argument in a revised version of the manuscript and focus on the most plausible arguments.

Lines 375-378: While the interpretation that age underestimation could have been caused by unreliable sigma_{phi} values and erosion of the boulder surfaces may be right, it would nevertheless be interesting to see what erosion rates one would get by applying the erosion rate model to samples that do not seem to suffer from age underestimation. The erosion rate of such samples must be negligible compared to the erosion rates of the samples showing age underestimation. This should provide a good basis for your interpretation.

The point developed by reviewer 2 is fair, and we will tackle this question in a revised version of the manuscript.

Line 377: Does “inadequate” mean “unreliable” here?

Yes. We will replace “inadequate” by “unreliable” in the revised version to make this absolutely clear.

Line 388: What is meant by “environmental factors beyond the exposure time”?

The factors referred to here, i.e. post-transport erosion and occasional shielding of the post-transport surface by e.g. water, are explained in the following sections of the manuscript. We will change the wording to “factors different than exposure time” to avoid any confusion.

Lines 418-419: It may be worth mentioning here that, in retrospect, IRSL signals were likely to work better than the post-IRSL-BSL signals for these samples.

We will add a short reference regarding the potential benefits of IRSL signals for some of our samples: “While IRSL signals were not used in this study due to insufficiently bright signals for most samples, in retrospect their use might be advantageous to post-IRSL-BSL signals at least for some of the investigated samples.

Line 422: What is considered as “insufficiently bright signals”? If the post-IRSL-BSL signals shown in Fig. A12 are typical for these samples, they are all well above background by more than 3.

The term “insufficiently bright” refers to the IRSL signals of polymineralic samples. Those shown in Figure A12a are representative for the samples of the different sites and hardly distinguishable from the background (signal < 3 times background for most aliquots).

Line 431-434: 1) The ages obtained from eroding surfaces are “apparent” surface exposure ages. The fact that they underestimate the expected ages, does not mean that they are inaccurate. They may be accurate, but they simply do not reflect the age of the event of interest. 2) There is no scientific basis to support this general statement that the ages from inclined surfaces are inaccurate. Surfaces can be dated regardless of their orientation provided that suitable calibration samples are available.

We agree with and appreciate these arguments. What we want to express is that apparent ages do not agree with age control due to erosion or unreliable calibration samples. We will change the phrasing of this section to better reflect this argumentation in a revised version.