

Editor

Earth Surface Dynamics

September 20, 2020

Dear Editor,

We would like to thank you for your handling and response to our submission. Please find enclosed the revised manuscript for *Earth Surface Dynamics*, entitled “*Different coastal marsh sites reflect similar topographic conditions for bare patches and vegetation recovery*” [Paper #esurf-2020-56], and detailed list of our responses to the comments of the two reviewers. We highly appreciated the comments made by the reviewers, as they enabled us to greatly improve the manuscript.

Below we give a step-by-step response to the comments. The original comments of the editor and reviewers are copied below and shown in black. Our step-by-step replies are inserted and shown in blue. The line numbers that are mentioned refer to the line numbers in the revised manuscript with tracked changes.

Thank you very much for your continued consideration of this manuscript.

Looking forward to your reply.

Yours sincerely,

On behalf of all co-authors,

Chen Wang

“Different coastal marsh sites reflect similar topographic conditions for bare patches and vegetation recovery” [Paper #esurf-2020-56]

Chen Wang, Lennert Schepers, Matthew L. Kirwan, Enrica Belluco, Andrea D'Alpaos, Qiao Wang, Shoujing Yin, and Stijn Temmerman

List of response to the comments

1. Response to the comments from Reviewer #1

Interactive comment on “Different coastal marsh sites reflect similar topographic conditions for bare patches and vegetation recovery” by Chen Wang et al.

Anonymous Referee #1

This paper is generally well written, studied the topographical conditions determining the presence and revegetation of bare patches in three marsh sites with contrasting tidal range, sediment supply and plant species, respectively distributed in three countries. The introduction is detailed, the methodology is well described, the results are clearly described, and the discussion is well-founded and consistent with current knowledge of the subject. Even so, I suggest some minor changes that I detail below, which I hope would be useful to improve this paper.

Response: We thank the reviewer for this overall positive evaluation.

(1) The information about some materials and methods should be more detailed and clear in the abstract.

Response: This is addressed on lines 19-21 of the revised manuscript (the version with tracked changes), which are modified to emphasize information on the materials and methods:

“Based on GIS analyses of aerial photos and LIDAR imagery of high resolution ($\leq 2 \times 2$ m pixels), we analyzed the topographic conditions under which bare patches occur, including their surface elevation, size, distance to and connectivity to channels.”

(2) The Study area, Materials and data preprocessing, and Data analysis are too long, it's better to make them a bit more concise.

Response: We shortened parts in the Sections of Study area, Materials and data preprocessing, and Data analysis. Especially the Study area was considerable shortened. We refer to the revised manuscript (the version with tracked changes) to see the text revisions.

(3) It's better to detail some of the implications of the results and some useful advice to the policy maker in the discussion.

Response: We added a paragraph in the discussion at the end of section 6.3 (line 523-527):

“Finally, our results may be indicative to decision makers on salt marsh management, as the formation of bare patches may be indicative for marsh degradation towards an unvegetated state that may be difficult to recover. Our study indicates that early signatures for marsh degradation must be particularly monitored in marsh portions, farthest away from main channels and with lowest surface elevations. Monitoring of early signatures is especially advised in systems with very low tidal range and suspended sediment availability.”

(4) I suggest to move some figures to the supporting information, as there are too many figures in the main text now.

Response: We moved figure 7 to the Supplement (Figure S4). Panels (b) of Fig. 1-3 are also moved to the Supplement. Fig. 9 (b) is removed, as suggested by the next reviewer. We also re-arranged Fig. 1-3 to save space.

2. Response to the comments from Reviewer #2

Interactive comment on “Different coastal marsh sites reflect similar topographic conditions for bare patches and vegetation recovery” by Chen Wang et al.

Maarten Kleinhans (Referee)

m.g.kleinhans@uu.nl

1 Main comments

This manuscript presents a data analysis of bare patches in saltmarsh, in particular of the causal variables deemed to govern their formation and possible revegetation. Three different systems are analysed with different tidal ranges and sediment availability. Two main conclusions seem not sufficiently well supported.

The first is that sediment availability and tidal range determine the potential for revegetation, but three study areas are insufficient to isolate one of these two variables, let alone assess their effect in combination. Two of the areas have low sediment availability and the one area

with more sediment also has the highest tidal range.

Response: Indeed we need to emphasize that we only investigated 3 sites, which it is not enough to fully assess the impact of site-differences, such as in tidal range and sediment supply, on occurrence and revegetation of bare patches. We only notice that revegetation only occurred at the site with largest tidal range and sediment supply, while it was not observed at the two other sites with smaller tidal range and sediment supply. But a much higher number of different sites should be investigated, to further assess the effects of tidal range and sediment supply on the occurrence and revegetation of bare patches.

We highlighted this in the discussion, by making the following modifications in the Discussion (section 6.3), lines 477-483:

“Previous modelling has suggested that pond formation increases and pond recovery decreases in marsh sites that are subject to a lower suspended sediment availability, smaller tidal range, and lower rate of relative sea level rise (RSLR) (Mariotti, 2016). First of all, we want to emphasize that we only investigated three sites, which it is not enough to fully assess the impact of site-differences, such as in tidal range, sediment supply, and rate of RSLR, on occurrence and revegetation of bare patches. Yet we notice that revegetation only occurred at the site (Saeftinghe) with largest tidal range and sediment supply, while it was not observed at the two other sites (San Felice and Blackwater) with smaller tidal range and sediment supply.”

It was also highlighted in the conclusion, lines 600-603:

“However, we emphasize that our study only included three sites, and that further research comparing much more sites is needed, to further advance our understanding of why certain marsh sites are more vulnerable to others to formation and persistence of bare patches. Such knowledge will be important to inform decision makers on site-specific priorities for marsh conservation.”

The second conclusion is that the appearance and possible disappearance of unvegetated patches in saltmarsh systems are acting as a bistable state system. While this concept is currently in fashion, the work done here is of interest in its own right and there appears no other support for the idea in this paper than the frequent use of it in other saltmarsh papers.

Response: The reviewer also comes back on this general comment in his more detailed comments below. We agree that our interpretation of results, in context of the alternative stable state theory, was inspired by previous literature on salt marshes, and interpretation of vegetated marshes and unvegetated tidal flats as alternative ecosystem states. We agree that, in this respect, the interpretation of our results is hypothetical. Therefore we followed the advice of the reviewer below, to remove section 6.4 from the discussion, and to remove panel (b) from Figure 9.

Accordingly, we also removed the last two sentences of the original abstract, where we summarized our interpretations in the context of the alternative stable state theory.

Furthermore there are some unanswered questions, such as whether inundation duration would not be a more appropriate biophysical boundary condition than the elevation in the tidal frame. A number of the variables that the study refers to, such as sediment availability, are not measured.

Response: These comments are further detailed below by the reviewer, and we will reply to them below.

Finally work needs to be done on the figures for a clearer presentation of the data and its context. These issues together, further detailed below, suggest that a moderate revision is needed.

Response: These comments are further detailed below by the reviewer, and we will reply to them below.

2 Detailed comments

2.1 Preamble and conclusions

The title does not reflect the contents and is ambivalent (do the similar topographic conditions refer to different coastal marsh sites or to bare patches and vegetation recovery?)

Response: We changed the title to make it more explicit, as follows:

“Different coastal marsh sites reflect similar topographic conditions under which bare patches and vegetation recovery occurs”

The abstract requires some clarification: the sentence "Our results demonstrate that ... distance from the main channels." Do the authors simply mean with 'across' that that all the sites show the same pattern? What kinds of channels are the patches connected to, since these are furthest away from the main channels (whatever they are)?

Response: We changed this sentence to make it clear (line 22-25):

“Our results demonstrate that, for the different marsh sites, bare patches can be connected or unconnected to the channel network, and that the width of the connecting channels increases with the size of the bare patches, in each of the three marsh sites. Further, pixels located in bare patches connected to channels occur most frequently at the lowest elevations and farthest distance from the channels.”

Hence we added here that the analysis is done for individual pixels (which are $\leq 2 \times 2$ m, as explained in the previous sentence in the manuscript), and that the frequency distribution of these pixels, located in bare patches connected to channels, peaks at the lowest elevations and farthest distance from channels. This is because connected bare patches are also the largest patches (this was added in the first sentence above), containing much pixels far away from channels. We don't make a distinction between 'main' channels and 'other' channels, or whatsoever, but we analyzed the frequency distribution of distances to all channels. Hence we removed the word 'main' channels.

The conclusion of the abstract that bare patches may form rapidly and become vegetated rapidly in the unstable zone at intermediate channel distances is based on only one of the sites, which begs the question whether the proposed existence of two stable states can be supported by the data, and how those bare patches at the other sites came about. Were they always unvegetated? Did they die off when the inundation duration increased, as the saltmarsh developed and reduced the outflow at these locations?

Response: Also in response to a previous remark above, we significantly reduced our interpretation of the results in terms of the theory on alternative ecosystem states. Because indeed, in this respect, our interpretation is rather hypothetical. Accordingly, in the abstract, we also removed this reference to alternative stable states.

Line 387 provides an interpretation of why the bare patches sit on higher elevations. This is based on expectations (meaning inferences without evidence), rather than measurement, and not even basic calculations (or readings from the classic wind waveheight plots on the basis windspeed, fetch (here patch size) and depth) are provided. Possibly the ideas here are biased by the reviewed literature as well and other alternative hypothesis could explain the observations. In Brückner et al. (2019, <https://doi.org/10.1029/2019JF005092>, also situated in the Western Scheldt) the modelling shows that expanding saltmarsh may, counterintuitively, lead to increased inundation duration within the marsh, which then leads to die-off. Indeed, the elevation within the tidal frame (as used here) may not be the appropriate measure. I wonder what the inundation duration, or perhaps the hydrodynamic energy, is at the elevations of the connected and the disconnected bare patches, and whether too long inundation has to do with the die-off (assuming these patches were vegetated before), as suggested in Brückner et al. This also fits with the observation that sediment supply is needed to lift up the area and reduce inundation for revegetation.

Response: Indeed we discuss several possible hypotheses quite extensively in section 6.1., by referring to existing insights from the literature. We thank the reviewer for focusing our attention to the paper by Brückner et al., and added it in the hypothetical discussion in line 441-446:

“This micro-topography of levees close to channels and depressions further away from channels, is often associated with an increasing inundation duration after high tides, and decreasing soil drainage/aeration during low tides, with increasing distance from channels (e.g., Ursino et al., 2004). Also, a modelling study suggested that marsh vegetation expansion can lead to increased inundation time, and as such can feedback on increased stress and

chance for vegetation die-back (Brückner et al. 2019). This may be all mechanisms that may contribute to increased chance for occurrence of bare patches within marshes at farther distances from channels.”

2.2 Comments on results

Figures 1 to 3 show insufficient context. One key variable for the authors is connectivity of the patches in terms of distance to channels, so the bigger context of the study areas must be shown to see the bigger and smaller channels. This would be more interesting information than the photographs in the panels (which have different meanings for colour anyway). An image or lidar map showing the surrounding landscape including the channels would be more useful here, and the original images can go to the online supplement. For Saeftinghe I checked and the study location in Figure 7 is quite close to the embanked boundary of the system (so the white band on the bottom left is in fact an embankment). In fact the right zone is quite close to an old embankment within the area and one wonders whether that leads to enhanced ponding and a modified channel pattern like one can see further east along the dike.

Response: We moved the aerial images (panels (b) in Figs. 1 to 3) to the online supplementary materials. Also, we added images of the larger surrounding landscape for the three study sites. We added these images also in the online supplementary materials, instead of in the main manuscript, as it contains already quite a lot of figures, and the other reviewer suggested reducing the number of figures in the main manuscript.

Why are there bare patches not considered in Fig. 3?

Response: This is explained in the methods section 3.4, line 237-241:

“In the Blackwater study site, we selected a study area away from the influence of roads and uplands (Fig. 3). The small study area (marked with shading in Fig. 3) was chosen for the field survey. A larger study area (the entire colored region in Fig. 3) was later considered in order to increase the number of bare patches connected to channels wider than 1 m. Bare patches that are connected with narrow channels (< 1 m) and that are located outside of the small study area (blue polygons in Fig. 3) were not considered in the analysis.”

Figure 4 has a lot of redundant header and axis text information and the real information is hidden on a few square centimeters. Likewise for Figure 5, where removal of the horizontal axis texts for panels a and b makes it possible to have higher plots on the same space, so that the data are more clearly shown and comparable. This is necessary, because what happens in the tails of these skewed distributions is interesting: the connected bare patches plot above the other distributions.

Response: In Figure 4, we removed the label “relative elevation” at the upper X-axis for panel (b) and (c). Further, we removed the label “Elevation (m MSL)” at the bottom X-axis for panels (a) and (b). As a result we could make the figures a bit larger in the vertical direction.

In Figure 5, we followed the same advice, by removing the label “distance to closest channel (m)” from the horizontal axis, and making the figures a bit larger.

Figure 6 contains novel information and shows interesting trends. However, the relative vertical axis per channel width class leads to a biasing emphasis on a very small number of cases for the largest channel widths. Perhaps another presentation would solve that problem: a matrix (pcolor in matlab) with log(patch size) on the horizontal axis, log(channel width) on the vertical, and log(number or fraction of total) as the colour scale. The channel width classes are not consistent with the possibly logarithmic distribution of the number of patches against channel width and I suggest to simply use classes of a 2-base log or something here, which would also improve the horizontal axis in Figure 8c from non-equidistant class to a true width scale.

Response: We tried alternative representations of the results, in line with the suggestions made here, but it did not improve the representation. Actually, the pattern of increasing connecting channel width with increasing patch size, becomes less clear. Therefore, we chose to keep the same figure format.

Figure 8 needs to mention in the caption that this concerns the Saeftinghe site only.

Response: Indeed this was done.

Is distance to the closest channel calculated from a map of channels or from the DEM?

Response: Channels are mapped based on aerial images, which is explained in section 3.1

How is the information in panel c obtained; is that the same as in Figure 6a but then split up for the permanent and temporary bare patches?

Response: Indeed.

Why is there no data for the other areas?

Response: For the considered time periods (see sections 3.2 to 3.3) we did not observe revegetation of bare patches in the other two areas (San Felice and Blackwater).

I suppose there are older images so this is open for analysis. As it stands now, there is very little data and support for the conclusions about stability and revegetation, especially since this plot is only for the system where the authors claim that revegetation is most likely. How do they know?

Response: Indeed our interpretations in the discussion, in the context of alternative stable state theory, were strongly hypothetical. Also in response to a previous similar remark by the reviewer, we left out this part of the discussion.

Figure 9a has four variables mentioned on the top arrow to the right, but is width of the connecting channel really increasing to the right, away from the widest main channel and into the bifurcating network? That is only possible if the reduction in depth goes much more rapid.

Response: Indeed, this may have been confusing. Therefore, the formulation was changed as

“connectivity of bare patches to channels”. This is indeed small close to the main channel in the drawing, and increases with increasing distance from the main channel.

Is erosion the right term here?

Response: we rephrased it as “resuspension”, as contrasting with the “sedimentation”.

How is it possible that sediment disappears in such a strongly converging flow (meaning very low velocity in the patches landward of the first bit of well-defined channel)? Are waves important here, as high up on the marsh in a very shallow, vegetated and micro-fetch area? Waves are known to be important in this sort of system, but that is on saltmarsh edges where there is fetch and depth to generate waves. It is not simply saltmarsh collapse and disappearance of organic material that causes the bare patches?

Response: previous studies have proposed that waves play a significant role in resuspension (or erosion) of pond bottom material in marsh ponds, and in lateral expansion of ponds by wave-induced erosion of the pond-marsh boundaries (Mariotti, 2016; Ortiz et al. 2017). But these studies identified a critical interior marsh pond size of 200–1000 m for wave-induced erosion of vegetated pond edges (i.e. effect of fetch length). However, many of the bare patches considered in our study, are smaller, and hence wave-induced erosion is expected to be very small. Tidal currents are then the expected dominant control over hydrodynamics, eventually responsible for resuspension and export of sediments from bare patches.

In Figure 9b, the horizontal axis provides two complex variables: sediment supply and soil drainage, but how do you know that it concerns these two and not the many others mentioned in lines 77-80? These are entirely inferred here but not measured. Any concentration from literature such as in line 125 is meaningless because of the very large spatial variation and the sediment settling in the marsh so far from the channels. So the position of the blue and green curves in the graphic is really unknown and we cannot know whether there are really two disconnected lines or simply a single continuum. And that means that the connection to the bistable state diagram is entirely speculative. I know it is attractive to try and see the landscape through the filter of the concept from complexity theory (citations here go back to Scheffer but the idea is already reviewed in Thorn and Welford 1994 <https://www.jstor.org/stable/2564149>), but this connection needs to be supported by the data. At present, it is not, and removal of this panel and section 6.4 of the discussion would in my opinion increase the quality of the paper.

Response: We followed this advice, and removed panel (b) of Figure 9, and removed section 6.4 from the paper.

The lidar images in the supplement are barely useful as presented here. The gray scale and small image size, and the lack of colour scale bar makes it very hard to see anything at all here.

Response: We replace them by aerial images of the larger surrounding landscape.

2.3 Suggestions for the text

The present objective (line 108) is now to determine the topographic conditions determining the presence of bare patches, but the idea also seems to determine whether they can revegetate, so I suggest 'presence and dynamics'.

Response: we followed this suggestion.

The authors define two kinds of bare patches, but surely this is a continuum and there is a certain image resolution. They need to indicate what size of connecting channel is the cutoff for an isolated or connected patch earlier than in line 397 in the discussion.

Response: Indeed, we also mentioned this threshold channel width in the methods section 3.1 in line 186-189.

The size of bare patches is important for the discussion (line 427) but size is not plotted in Fig. 6, only number of pixels and that could also indicate many small patches. A plot of patch size, and possibly analyses with patch size as a variable, are needed to make this argument.

Response: Patch size is indeed plotted on the X-axis of Figure 6.

Different coastal marsh sites reflect similar topographic conditions under which bare patches and vegetation recovery occurs~~for bare patches and vegetation recovery~~

Chen Wang^{1,2}, Lennert Schepers², Matthew L. Kirwan³, Enrica Belluco⁴, Andrea D'Alpaos⁵, Qiao Wang^{6,1}, Shoujing Yin¹, and Stijn Temmerman²

¹Satellite Application Center for Ecology and Environment, Ministry of Ecology and Environment / State Environmental Protection Key Laboratory of Satellite Remote Sensing, Fengde East Road 4, Beijing 100094, China

²Ecosystem Management Research Group, University of Antwerp, Universiteitsplein 1, 2610 Wilrijk, Belgium

³Virginia Institute of Marine Science, PO Box 1346, 1375 Greate Road, Gloucester Point, Virginia 23062, USA

⁴Department of Civil, Environmental, and Architectural Engineering, University of Padova, via Loredan 20, 35131 Padua, Italy

⁵Department of Geosciences, University of Padua, Via Gradenigo 6, 35131 Padua, Italy

⁶Faculty of Geographical Science, Beijing Normal University, Xijiekouwai Street 19, Beijing 100875, P.R.China

Correspondence to: Qiao Wang (wangqiao@mee.gov.cn) and Shoujing Yin (shoujingy@163.com)

Abstract. The presence of bare patches within otherwise vegetated coastal marshes is sometimes considered to be a symptom of marsh die-back and the subsequent loss of important ecosystem services. Here we studied the topographical conditions determining the presence and revegetation of bare patches in three marsh sites with contrasting tidal range, sediment supply and plant species: the Scheldt Estuary (the Netherlands), Venice Lagoon (Italy), and Blackwater Marshes (Maryland, USA). Based on GIS analyses of aerial photos and LIDAR imagery of high resolution (< 2 x 2 m pixels), we analyzed the topographic conditions under which bare patches occur, including their surface elevation, size, distance to and connectivity to channels.~~We analyzed topographic properties of bare patches, including elevation, size, distance and connectivity to channels based on GIS analyses of aerial and LIDAR imagery.~~ Our results demonstrate that, for the different marsh sites, bare patches can be connected or unconnected to the channel network, and that the width of the connecting channels increases with the size of the bare patches, in each of the three marsh sites. Further, pixels located in bare patches connected to channels occur most frequently at the lowest elevations and farthest distance from the channels.~~Our results demonstrate that, for across the different marsh sites, the size distribution of bare patches shows a pattern of larger bare patches, which are connected to channels, than bare patches that are not connected to the channel network. Further, pixels located in bare patches connected to channels occur most frequently at the lowest elevations and farthest distance from the main channels.~~ Pixels in bBare patches disconnected from channels occur most frequently at intermediate elevations and distances from channels, and vegetated marshes dominate at highest elevations and shortest distances from channels. Revegetation in bare patches is observed in only one site with the highest tidal range and highest sediment availability, and preferentially occurs from the edges of small unconnected bare patches at intermediate elevations and intermediate distances from channels. ~~Our results are discussed within the alternative stable state theory. We suggest the existence of two stable~~

35 ~~states, a high elevated vegetated state close to channels that tends to remain high and vegetated, and a low elevated state of bare connected patches far from channels that tends to remain bare, with an unstable state at intermediate channel distances where bare patches may form and rapidly become revegetated. Although our study is only for three different marsh sites with large variations in local conditions, such as tidal range, sediment availability and plant species, it suggests that similar topographic conditions determine the occurrence of bare patches. Such insights may inform decision makers on coastal marsh management on where to focus monitoring of early signatures of marsh degradation.~~

40 **1 Introduction**

Tidal marshes are coastal ecosystems that provide many valuable ecosystem services such as fishery production (Barbier et al., 2011), sequestration of CO₂ (McLeod et al., 2011), protection against shoreline erosion and mitigation of flood risks during storm surges (Barbier et al., 2008; Wamsley et al., 2010; Gedan et al., 2011; Temmerman et al., 2013; Temmerman and Kirwan, 2015). However, tidal marshes and their valuable ecosystem services can be lost when marshes die-off, for
45 instance, as a consequence of sea level rise. Large-scale tidal marsh loss by conversion of marshes into bare tidal flats, open water or bare patches within marshes has been reported from different locations around the world (Baumann et al., 1984; Day et al., 2000; Kearney et al., 2002; Carniello et al., 2009; Kirwan and Megonigal, 2013). Bare patches within marshes, which are often covered by standing water and then referred to as pools, ponds (Stevenson et al., 1985) or marsh basins (Mariotti and Fagherazzi, 2013), are a common feature in salt marshes around the world. In some regions, bare patches are
50 dynamic features that develop but also recover and revegetate (e.g., New England; Wilson et al., 2009, 2010, 2014). In other areas, however, bare patches do not revegetate and are causing permanent marsh loss on a large scale (e.g. Mississippi Delta Penland et al., 2000; Morton et al., 2003).

Marsh loss and recovery is of particular concern because there is growing evidence that vegetated marshes and bare flats behave as alternative stable ecosystem states (Fagherazzi et al., 2006; Kirwan and Murray, 2007; Marani et al., 2007; van
55 Wesenbeeck et al., 2008; Marani et al., 2010; D'Alpaos, 2011; McGlathery et al., 2013; Wang and Temmerman, 2013; Moffett et al., 2015; D'Alpaos and Marani, 2016; van Belzen et al., 2017), which implies that recovery after marsh loss would be very difficult (Hu et al., 2015a; van Belzen et al., 2017). Observations have shown that vegetated marshes and bare flats occupy different elevation ranges (Marani et al., 2007, 2010; Carniello et al., 2009; Wang and Temmerman, 2013) and that shifts from the low-lying bare state to the high-elevated vegetated state occur rapidly once a threshold elevation has been
60 exceeded (Wang and Temmerman, 2013). Moreover, models indicate that the system would shift abruptly between the high elevation vegetated state and low-lying bare state when a threshold value is reached in elevation, sediment input, or rate of sea level rise (Fagherazzi et al., 2007; Kirwan and Guntenspergen, 2010; Marani et al., 2010; D'Alpaos et al., 2011; D'Alpaos and Marani, 2016). Previous studies further suggest that the state shift can be irreversible because of a hysteresis effect (Kirwan and Murray, 2007; Marani et al., 2010; Kirwan et al., 2011), where the threshold conditions to revert the
65 ecosystem back to the original state are far more difficult to reach than the threshold conditions that caused the shift

(Scheffer et al., 2001; Scheffer and Carpenter, 2003). Field experiments have also demonstrated that vegetation recovery after disturbance is slower under increased tidal inundation, which further suggests the applicability of alternative stable state theory to vegetated and bare areas in intertidal zones (van Belzen et al., 2017).

70 The two stable states of marshes and tidal flats can be explained by positive feedback mechanisms which are strongly mediated by the presence or absence of marsh vegetation. As long as vegetation is present on the marsh, waves and tidal currents are effectively attenuated by vegetation-induced friction over several meters of continuously vegetated marsh surfaces (Neumeier and Amos, 2006; Mudd et al., 2010; Vandenbruwaene et al., 2011; Yang et al., 2012; Hu et al., 2014). As a consequence, suspended sediment is deposited on the marsh surface and marshes can maintain a high position in the tidal frame, even with sea level rise (Kirwan and Guntenspergen, 2010; D'Alpaos et al., 2011; Fagherazzi et al., 2012).
75 Above- and belowground plant material further contributes to marsh accretion (Nyman et al., 2006; Kirwan and Guntenspergen, 2012). When vegetation is absent, however, organic matter accumulation is strongly reduced, and increased tidal currents and waves may prevent sedimentation or even trigger erosion (Fagherazzi et al., 2006; Kirwan and Murray, 2007; Marani et al., 2007; Mariotti and Fagherazzi, 2010; Temmerman et al., 2012). In large lagoons or extensive tidal basins, the low elevation of the tidal flats is mainly maintained by wave erosion (Fagherazzi et al., 2006; Mariotti and
80 Fagherazzi, 2010; Hu et al., 2015b). The existence of these two alternative stable states has been empirically observed on the large scale of whole tidal basins where large areas ($\sim\text{km}^2$) of marshes and tidal flats may coexist next to each other (Marani et al., 2007; Carniello et al., 2009; Wang and Temmerman, 2013). However, the existence of alternative stable states has not yet been empirically explored to explain marsh loss and recovery by formation and revegetation of bare patches ($\sim 10\text{--}100$ m^2), which is addressed in this paper.

85 Bare patches are defined here as non-vegetated areas in the interior of otherwise vegetated marshes. Here we consider two types of bare patches: (i) connected bare patches that have a connection to the tidal channel network and (ii) isolated bare patches that are separated from the channels by surrounding marsh vegetation. Literature suggests that unconnected bare patches start as areas with vegetation die-off, by increased flooding stress and inadequate drainage, high salinity stress (DeLaune et al., 1994; Wilson et al., 2009, 2014), coverage by drifted plant material (Harshberger, 1916; Miller and Egler,
90 1950; Redfield, 1972), physical disturbance by ice, or herbivory by crabs, nutria, muskrats, geese or snails (Harshberger, 1916; Stevenson et al., 1985; DeLaune et al., 1994; Silliman, 2005; Argow and FitzGerald, 2006). Subsequent elevation loss due to the collapse of the root structure or decomposition and disintegration of soil organic matter can deepen the bare patches (DeLaune et al., 1994; Wilson et al., 2014). Connected bare patches form by creek-bank erosion at the creek heads (Kearney et al., 1988) and subsequent connection of channel heads to bare patches (Redfield, 1972) or by expansion of
95 unconnected bare patches that ultimately reach a channel and become hydraulically connected to the channel network (Wilson et al., 2014; Mariotti, 2016).

Nevertheless, it is not fully understood under which topographic conditions connected and unconnected bare patches occur, and especially under which conditions they recover through re-establishment of vegetation. For example, we may hypothesize that unconnected bare patches are buffered from tidal currents and waves by the surrounding marsh vegetation,

100 and therefore are less prone to erosion and more suitable for vegetation recovery. On the other hand, they might also receive less sediment input since sediment is efficiently trapped by the surrounding vegetation buffer (Mudd et al., 2010; Moskalski and Sommerfield, 2012). The opposite applies for connected bare patches: they receive direct sediment input through the channels, but experience higher flow velocities that may cause sediment resuspension and erosion. Some studies show that marsh plants might recolonize bare patches when they become connected, drain and if vertical accretion elevates the bare patches sufficiently for plant establishment (Redfield, 1972; Wilson et al., 2009, 2014). However, higher flow velocities and therefore a decrease in accretion by reduced mineral sediment deposition or erosion may inhibit the recovery of vegetation in connected bare patches (DeLaune et al., 1994; Mariotti, 2016).

Hence, despite the fact that bare patches are often recognized as symptoms of marsh loss (Kearney et al., 1988; DeLaune et al., 1994; Fagherazzi et al., 2013; Mariotti and Fagherazzi, 2013; Wilson et al., 2014; Mariotti, 2016), there are relatively few studies on the dynamics of bare patches. For example, the modelling study by Mariotti (2016) simulates that pond expansion is favored under conditions with low tidal range, low sediment supply and high relative sea level rise. Apart from this study, there is poor empirical evidence on the conditions that determine the presence and/or recovery of bare patches, especially across marsh sites that differ in characteristics such as tidal range, sediment supply and plant species. In this paper, we first study the topographic conditions determining the presence and dynamics of bare patches. Next, we study the topographic conditions determining the marsh vegetation recovery (i.e., the re-establishment of vegetation) in bare patches. To identify the topographic conditions determining the presence of bare patches, we compared the surface elevation, bare patch size and distance to channels for connected and unconnected bare patches in three different sites, located in the Scheldt Estuary (a river-dominated estuary in the Netherlands, 4.8 m tidal range), Venice lagoon (a back-barrier lagoon in Italy, 1.0 m tidal range) and Blackwater Marshes (a submerging tidal marsh, in Maryland, USA, < 0.5 m tidal range). To identify the conditions determining the revegetation of bare patches, we carried out a time series analysis in the Scheldt Estuary, the only site where revegetation was observed and searched for relations between the rate of revegetation of bare patches and topographic conditions including surface elevation, distance to channels and the width of connecting channels. Our hypotheses are that (1) bare patches across all three study sites are found at similar elevation relative to the tidal frame, distance to tidal channels and degree of connectivity to tidal channels; (2) low elevation relative to the tidal frame and wide channel connection lead to larger bare patches that are more difficult to revegetate.

2 Study area

~~We studied three marsh sites that have different characteristics, including different tidal range, sediment supply and plant species: (i) Saeftinghe (the Netherlands), (ii) San Felice (Italy) and (iii) the Blackwater Marshes (USA). Within each of the three marsh sites, a specific study area was selected based on data availability and the presence of bare patches on aerial images. In the next paragraph we give more background information on the three marsh sites. Detailed information about the aerial images is provided in Sect. 3.~~

2.1 Saeftinghe marsh, Scheldt estuary, the Netherlands

The Scheldt estuary is a river-dominated estuary located in the southwest of the Netherlands and the northwest of Belgium (Fig. 1). The Saeftinghe marsh (51.33° N, 4.17° E) is a 3000 ha brackish tidal marsh ~~within the brackish zone of the estuary~~. It is subject to a semi-diurnal tidal regime with a local mean tidal range of 4.88 m, a salinity of 5–18 PSU, and a mean suspended sediment concentration (SSC) of 30–60 mg L⁻¹ (Temmerman et al., 2003a; van Damme et al., 2005). In the last 80 years, a long-term rise of mean high water level (MHWL) was observed in the Saeftinghe marsh at a rate of 5.7 mm/yr, while the vegetated marsh regions expanded in area and increased in elevation steadily and continuously (Wang and Temmerman, 2013). ~~The lower areas are colonized by the pioneer-Dominant plant species include *Spartina anglica*, and *Salicornia europaea*, *Scirpus maritimus*, is found in depressions of higher marshes. *Elymus athericus*, and *Phragmites australis*, etc is present on natural levees along creek edges. The highest parts are dominated by *Phragmites australis*.~~ Marsh vegetation is observed between -2 m and +1 m relative to MHWL, ~~with the highest frequency of vegetation presence centered around MHWL~~ (Wang and Temmerman, 2013). Parts of the Saeftinghe marsh have been converted to bare patches. This is partly attributed to geese grubbing for below-ground tubers (Elschot et al., 2017). In addition, bare patches are formed at places with poor drainage ~~and temporary ponding of water~~ after high overmarsh tides, ~~This is especially the case at marsh platforms~~ near the head of the smallest tidal channels, ~~i.e. where drainage towards channels is least developed~~. This is the case in the selected study site, ~~situated in the south of the Saeftinghe marsh~~, covering an area of 35 ha (Fig. 1).

2.2 San Felice marsh, Venice lagoon, Italy

The Venice lagoon is a back-barrier tidal lagoon situated in the northeast of Italy, ~~and is~~ characterized by a micro-tidal semi-diurnal regime with a mean tidal range of about 1.0 m (Day et al., 1999) and a maximum tidal range of 1.5 m (Rinaldo et al., 1999a, 1999b; Marani et al., 2007). The long-term rate of relative sea level rise varies around 3–4 mm/yr (Carbognin et al., 2004). The marsh systems area in the Venice lagoon ~~are degrading with~~ has decreased by about 75 % ~~reduction in marsh area~~ since 1901 ~~(from a 64 km² to 43 km²)~~, caused by both drowning and lateral erosion of marshes (Tommasini et al., 2019). The San Felice salt marsh (45.48°N, 12.46°E) ~~is located in the northern part of the Venice Lagoon, close to the Lido inlet (see Marani et al., 2003 for further details on the study site) and is considered to be~~ one of the best preserved marshes in the Venice Lagoon, being capable of keeping pace with current relative sea level rise (e.g., Roner et al., 2016; Marani et al., 2003). The average salinity varies between 24 and 33 PSU (Gieskes et al., 2013; Zirino et al., 2014), and the average SSC is between 10 and 20 mg/l (Zaggia and Ferla, 2005; Defendi et al., 2010; Venier et al., 2014). The salt marsh is occupied by halophytic species, ~~such as *Salicornia veneta*, *Spartina maritima*, *Limonium narbonense*, *Sarcocornia fruticosa*, *Puccinellia palustris*, *Inula crithmoides*, *Juncus maritimus*~~ (Silvestri et al., 2005; Marani et al., 2006). ~~The pioneer species present on the lowest elevations are mainly *Salicornia veneta* and *Spartina maritima*. *Limonium narbonense* covers slightly higher elevations. *Sarcocornia fruticosa* dominates the highest elevations, such as natural levees, together with *Puccinellia palustris* and *Inula crithmoides*. *Juncus maritimus* is observed within a broad range of elevations (Silvestri et al., 2005).~~ The elevation

of the salt marsh ranges from 0 m to 0.7 m relative to mean sea level (MSL). Our ~~specific~~ study site ~~in the interior of the San~~
165 ~~Felice marsh~~ has an area of 72.3 ha (Fig. 2).

2.3 Blackwater marshes, Chesapeake Bay, USA

The Blackwater marshes (38.40° N, 76.08° W) ~~are located along the ,part of the largest marshland in the~~ Chesapeake Bay
~~(Maryland, USA), are situated at the confluence of the Blackwater and Little Blackwater Rivers.~~ They cover an area of about
6000 ha with an average SSC of about 50 mg/L and an average salinity of 10 PSU (Stevenson et al., 1985; Ganju et al., 2013;
170 Kirwan and Guntenspergen, 2015). Long-term local sea level rise is currently 3.7 mm/yr (NOAA station 8571892,
<http://tide-sandcurrents.noaa.gov/sltrends>, 12/19/2016). ~~Extensive marsh loss was reported in the Blackwater system, where~~
~~a~~ About half of the interior marshes have disappeared since 1938, mainly by the development and enlargement of bare
patches, which are occurring as interior marsh pools (Stevenson et al., 1985; Kearney et al., 1988; Kirwan and
Guntenspergen, 2012; Schepers et al., 2017). The ~~marsh loss and~~ pool expansion has been attributed to submergence by sea
175 level rise, ~~and~~ vegetation disturbance by invasive herbivores and subsequent open-water expansion (Stevenson et al., 1985;
Kendrot, 2011). Changes in water level are mainly driven by ~~wind setup and barometric pressure fluctuations during~~
meteorological events (~~wind and air pressure~~), while the astronomical tidal range is about 0.25 m at our study site. Brackish
vegetation dominates ~~the Blackwater marshes~~, with species such as *Scirpus americanus*, ~~and~~ *Spartina alterniflora*
~~occupying low elevations, and~~ *Spartina patens*, *Distichlis spicata*, *Spartina cynosuroides*, and *Phragmites australis*
180 ~~occupying higher elevations~~ (Pendleton and Stevenson, 1983; Kirwan and Guntenspergen, 2012). Our ~~specific~~ study area ~~in~~
~~the Blackwater Marshes~~ covers an area of about 699.8 ha (Fig. 3).

3 Materials and data preprocessing

3.1 General procedure

For all three study sites, aerial photographs were digitized, georeferenced and manually classified into vegetated marshes,
185 unconnected bare patches, connected bare patches and tidal channels (Figs. 1-3). Bare patches that were smaller than 1 m²
were not considered in this study. Given the resolution of the images (see below), bare patches were classified as to be
connected to the channel network when the connecting channel was at least 0.5 m wide. Hence our classification of
unconnected bare patches ~~includes truly unconnected patches, but~~ may also include patches with a small connecting channel
(less than 0.5 m wide) ~~that was impossible to detect on the aerial images~~. LIDAR data (~~Figs. S1-S3~~) was used to analyze the
190 elevation differences between vegetated marshes, unconnected bare patches, connected bare patches and tidal channels.
When bare patches were inundated during the LIDAR survey, the soil surface elevation within the bare patches was
measured with field surveys (methods are explained below for the different study sites). Generally, LIDAR data have larger
and more homogeneous spatial coverage and higher spatial resolution. Field surveys only include selected locations, but with

greater vertical accuracy, especially for vegetated areas where LIDAR partially reflects on the vegetation canopy, and open
195 water where LIDAR reflects on the water surface. All the spatial analyses were done using ArcGIS.

3.2 Saeftinghe

For the Saeftinghe study site, a time series of false-color aerial images was used, from 1990, 1998, 2004 and 2008. The four
images were selected considering the data availability and to detect dynamic changes from vegetated marsh portions into
bare patches and vice versa. All the photos were processed in a similar way, by scanning, georeferencing and mosaicking
200 them into digital pictures with a minimum resolution of 0.5 m. All the aerial images were provided by Rijkswaterstaat (the
Dutch governmental institute for water management) (Huijs, 1995; van der Pluijm and de Jong, 1998; Reitsma, 2006; Bakker
and Bijkerk, 2009). From all the available aerial photographs, we extracted two sample areas (Fig. 1) free from drifted plant
debris, which were analyzed together. The digitized aerial images in the sample areas were classified into vegetation, water
and bare soil based on supervised maximum likelihood classification, and then further classified visually into vegetated
205 marshes, channels, connected bare patches and unconnected bare patches. For elevation data in Saeftinghe, we used a Digital
Terrain Model (DTM) with a resolution of 2×2 m (Fig. S1), which was obtained from a LIDAR survey performed in 2004
during low tide with a maximum vertical error of 0.2 m (Alkemade, 2004). The measurement point density of the LIDAR
survey varied from 1 point per 16 m^2 to several points per m^2 . The DTM data were also provided by Rijkswaterstaat. We
used only one LIDAR dataset to derive the elevations of bare patches and marshes over the period 1990-2008 because
210 previous research in the area showed that during that period, elevation changes were limited with maximum rates of 1 cm/yr
(Wang and Temmerman, 2013). This implies that over the considered time scale (1990-2008), maximum elevation changes
(~18 cm in 18 years) are of the same order of magnitude as the vertical error of the LIDAR data (~20 cm). Therefore, we
decided to use one LIDAR-based DTM for 2004, which is considered to be representative to characterize the approximate
time-averaged elevation of marshes and bare patches over the period 1990-2008. No field survey data were used for
215 Saeftinghe since all bare patches drain completely during neap tides so that soil surface elevations were recorded by LIDAR.

~~We note that even so-called unconnected bare patches may have drainage via connecting channels < 0.5 m wide (see above)
or truly unconnected bare patches may drain at neap tides by subsurface drainage towards nearby channels, that are typically
1-3 m deep and also dry during low tides in this macro-tidal setting (mean tidal range of 4.9 m).~~

3.3 San Felice

220 For the San Felice study site, our analysis was based on a vegetation map classified from a hyperspectral image with a
resolution of 1.3 m, which was acquired in 2002 by the airborne CASI sensor (15 bands in the visible and near infrared
portion of the spectrum) (Belluco et al., 2006). ~~The vegetation map distinguished water, bare soil and four vegetation classes.~~
It was visually reclassified into channels, connected bare patches, unconnected bare patches and vegetated marshes (Fig. 2).
For the latter, we consulted a black-white aerial photograph acquired in 2000 with a resolution of 16 cm, and a 1-meter
225 resolution pan-sharpened multispectral IKONOS image acquired in 2006. For elevation data in San Felice, we used both a

DTM obtained from a LIDAR survey (Fig. S2) and field measurements. The LIDAR survey was performed during low tide in 2002 with a mean measurement point density of about 48 points/m² and a vertical accuracy better than 0.15 m (Wang et al., 2009). From these data we constructed a gridded DTM with a spatial resolution of 1×1 m. Field elevation measurements from the Venice Water Authority in 2000 were also used, because some bare patches were inundated during the LIDAR survey. Data were collected with stereo aerial photography for marshes, stadia rods with GPS for areas close to marshes and mudflats, and single-beam echo-sounder for shallow waters (Sarretta et al., 2010). In total, 340 elevation measurements were located in vegetated marshes, and 95 measurements in bare patches. The boundary of the study area was delineated by channels and creeks as shown in Fig. 2, considering the availability of data. Since almost no vegetation recovery in bare patches was observable on aerial images from the San Felice marsh, we did not do a time series analysis on vegetation recovery.

3.4 Blackwater

In the Blackwater study site, we selected a study area away from the influence of roads and uplands (Fig. 3). The small study area (marked with shading in Fig. 3) was chosen for the field survey. A larger study area (the entire colored region in Fig. 3) was later considered in order to increase the number of bare patches connected to channels wider than 1 m. Bare patches that are connected with narrow channels (< 1 m) and that are located outside of the small study area (blue polygons in Fig. 3) were not considered in the analysis. We used false color aerial photographs with a spatial resolution of 0.3 m obtained in 2010 and provided as digitized and georeferenced mosaic by the United States Department of Agriculture (USDA). ~~Similar as in Saeftinghe, we~~ classified the photos into ~~water, vegetation and bare soil using a supervised maximum likelihood classification procedure, and then we visually classified them into~~ vegetated marshes, connected bare patches, unconnected bare patches and channels, ~~using the same method for Saeftinghe~~. We also used data acquired from a LIDAR survey (Fig. S3) and a field elevation survey. The LIDAR data were obtained in 2003 with an average area sampling density of about 0.8 points per m² and a mean vertical accuracy of 0.14 m. The DTM was provided with a resolution of 2×2 m by the U.S. Geological Survey and Maryland Department of Natural Resources. As most bare patches were covered by water during the LIDAR survey, a field survey was carried out in 2012 in the small study area using RTK-GPS with ±1.5 cm accuracy. In total, 36 elevation measurements were collected in 5 unconnected bare patches, 31 measurements in 5 connected bare patches and 93 measurements in the vegetated marshes. An overview of the number of data points (LIDAR and GPS measurements) that fall within marshes and bare patches are given for the different study sites in Table 1. ~~We did not do a time series analysis on vegetation recovery because o~~Other studies ~~in the Blackwater Marshes~~ have demonstrated that recovery ~~of marsh vegetation within bare patches~~ is absent (Scheepers et al., 2017), ~~and therefore we did not analyze a time series of aerial images~~.

4 Data analysis

4.1 Topographic conditions determining the presence of bare patches

In order to identify the topographic conditions determining the presence of bare patches or marsh vegetation, we analyzed the frequency distributions of surface elevation and distance to channels for connected and unconnected bare patches, and compared them with the vegetated marsh portions, for the three study sites. The surface elevation was analyzed using LIDAR data and field data. The distance to channels was calculated as the Euclidean Distance from the edge of channel polygons. Bare patches smaller than 1 m² were excluded from the analysis. The nonparametric Mann-Whitney U test was conducted to test whether or not there were significant differences between elevations of vegetated marsh portions, connected and unconnected bare patches.

Elevation classes of 10 cm were used, since smaller elevation classes were not deemed to be reasonable considering the vertical accuracy of LIDAR data. Surface elevation relative to the local mean low and high water levels (i.e., the tidal frame), is an important factor for vegetation because it determines the frequency, depth and duration of tidal flooding and is widely considered as a crucial ecological condition for marsh plant growth. Therefore, in order to allow comparisons between the three marsh sites with largely different tidal ranges, we rescaled the surface elevation relative to the tidal frame using the following relationship:

$$RE = \frac{E - MLWL}{MHWL - MLWL} \quad (1)$$

where RE is the relative elevation (a dimensionless proportion of the local tidal frame), E is the actual elevation (in m relative to a fixed datum), $MLWL$ and $MHWL$ are the mean low water level and mean high water level, respectively (in m relative to the same datum). Hence RE is 0 for elevations equal to $MLWL$ and is 1 for elevations equal to $MHWL$.

In addition, the frequency distribution of bare patch sizes was calculated and related to the widths of channels that were connected to bare patches. The channel width was measured on the aerial photographs at the connection with the bare patch for each single patch and classified into categories with 5 m spacing. Unconnected bare patches (channel width < 0.5 m) and bare patches connected with small channels (channel width between 0.5 m and 1 m) were classified as two separate categories because of their large number. We combined all bare patches with a connection > 80m in the highest class, since there were only 0, 1 and 2 patches for this category in the Saeftinghe, San Felice, and Blackwater marsh sites, respectively.

4.2 Topographic conditions determining the revegetation of bare patches in Saeftinghe

We studied revegetation of bare patches in the Saeftinghe marsh during the last two decades. We did not include the San Felice and Blackwater marshes in this analysis, because there was almost no revegetation recognizable on the aerial photographs during this period of the last two decades. ~~This does not necessarily mean that revegetation of bare patches is not taking place in the latter two study areas, but at least suggests that it is not occurring on the time scale of the last two decades.~~ Between each aerial photograph in 1990, 1998, 2004 and 2008, we identified areas that changed from vegetated to

bare surfaces, areas that revegetated from bare to vegetation, and areas that remained bare or vegetated. From this data, we determined the rate of revegetation of bare areas. We made a distinction between the following classes:

- (1) permanent bare patches that never revegetated within the considered time period from 1990 to 2008;
- 290 (2) rapidly revegetated bare patches, identified as bare in only one image, either 1998 or 2004, and observed as vegetation in the other three images;
- (3) permanent marsh areas, classified as vegetation throughout the time series.

In order to identify the topographic conditions for rapid or no revegetation of bare areas, the frequency distribution of elevation was calculated for these three classes (permanent bare patches, rapidly revegetated bare patches, and permanent
295 marsh areas), as well as the frequency distributions of the distance to the closest channel. In addition, we also determined the width of the channels connecting to the bare patches. For permanent bare patches, the channel width is calculated as the mean value for 1990, 1998, 2004 and 2008. For rapidly revegetated bare patches, the channel width is the value when the bare patches occurred, either in 1998 or in 2004. In order to identify the relationship between the rate of revegetation and the width of connecting channels, the frequency distribution of channel widths was compared between permanent bare patches
300 and rapidly revegetated bare patches.

5 Results

5.1 Topographic conditions determining the presence of bare patches

In order to identify the topographic conditions determining the presence of bare patches, we tested relationships between their presence and three topographic variables, which are (1) elevation of the bare soil surface, (2) distance of the bare
305 patches from channels and (3) channel width for bare patches connected to channels (channel width <0.5 m for unconnected bare patches). We first tested whether these three topographic variables are independent from each other. The correlations were low (Pearson's $r < 0.5$) and not significant ($p > 0.05$) between all variables and for all field sites. Only for the Blackwater marsh, the correlation between the elevation and the channel width was high (Pearson's $r = -0.9$), but this correlation is based on a very low number of connected ($n=5$) and unconnected bare patches ($n=5$).

310 5.1.1 Elevation

In Saeftinghe, the connected bare patches, unconnected bare patches and vegetated marshes fall within the elevation ranges of 2.3–3.5 m above MSL, which is close to the local MHWL (relative elevation $RE = 0.91–1.16$) (Fig. 4a). The differences in elevation between the vegetated marshes, connected and unconnected bare patches were statistically significant between each two of the three features ($p < 0.001$ based on the Mann-Whitney test). The peaks of the elevation distribution for the
315 vegetated marshes and unconnected bare patches are 0.1 m higher than for the connected bare patches (or the difference between relative RE , $\Delta RE = 0.02$). The mean elevation of the vegetated marshes is highest (2.97 m above MSL, $RE = 1.05$),

whereas this is 0.14 m lower for the unconnected bare patches ($\Delta RE = 0.03$) and 0.23 m lower for the connected bare patches ($\Delta RE = 0.05$).

320 In San Felice, the connected bare patches, unconnected bare patches and vegetated marshes are situated in different ranges of elevations between -0.5 and +0.7 m relative to MSL ($RE = 0-1.2$, Fig. 4b). The differences in elevation distributions of these three categories are also statistically significant ($p < 0.001$ based on the Mann-Whitney test). The elevation measured in the field is lower than that from the LIDAR survey for both the connected bare patches and vegetated marshes. The peaks of the elevation distribution of the vegetated marshes and unconnected bare patches are about 0.15 m lower than MHWL ($RE = 0.85$) based on LIDAR data, and about 0.3 m or 0.5 m higher than connected bare patches ($\Delta RE = 0.3$ or 0.5) based on
325 LIDAR or field data, respectively. The mean LIDAR elevation of the vegetated marshes is 0.35 m relative to MSL ($RE = 0.85$), which is 0.04 m higher than unconnected bare patches ($\Delta RE = 0.04$) and 0.28 m higher than connected bare patches ($\Delta RE = 0.28$).

In the Blackwater Marshes, connected bare patches, unconnected bare patches and vegetated marshes occupy significantly different ranges of elevations ($p < 0.001$ based on Mann-Whitney test) between -0.7 m and +0.5 m relative to MSL ($RE = -$
330 $0.9-1.5$, Fig. 4c). The peaks of the elevation distribution of the vegetated marshes are 0.1 m lower than MHWL ($RE = 0.8$), 0.3 m higher than unconnected bare patches ($\Delta RE = 0.6$) and 0.6 m higher than connected bare patches ($\Delta RE = 1.2$). The mean elevation is the highest for the vegetated marshes (0.13 m relative to MSL, $RE = 0.76$), 0.23 m lower for the unconnected bare patches ($\Delta RE = 0.46$) and 0.6 m lower for connected bare patches ($\Delta RE = 1.2$).

Together these results indicate that connected bare patches, unconnected bare patches, and vegetated marshes tend to occupy
335 different elevation ranges at each site ($p < 0.001$ by Mann-Whitney), with the largest absolute elevation differences in Blackwater, the smallest in Saeftinghe, and intermediate values for San Felice. Connected bare patches always lie within the lowest elevation range, whereas vegetated marshes always dominate the highest elevation range around MHWL. Unconnected bare patches are always found in the intermediate elevation range, which is about 0.1–0.5 m higher than the connected bare patches. The difference in RE (relative to the tidal frame) between the connected and unconnected bare
340 patches is about 0.02 in Saeftinghe, 0.2–0.5 in San Felice, and 0.6–0.8 in Blackwater.

5.1.2 Distance to channels

The frequency distribution of the distance between a bare patch and closest channel shows similar results for the three marsh sites (Fig. 5). Vegetated marshes rather than bare patches occur near channels. With increasing distance from channels, marsh vegetation becomes less frequent and unconnected bare patches become more frequent. Connected bare patches occur
345 most frequently at large distances from the channels. The peak of the distribution is situated at 1.0 m for vegetated marshes in all three sites; at 8 m for unconnected bare patches and over 10 m for connected bare patches in both Saeftinghe and San Felice; and at 82 m and 89 m for unconnected and connected bare patches in Blackwater, respectively.

5.1.3 Bare patch size in relation to connectivity to channels

Bare patch size generally increases with increasing width of connecting channels, whereas the number of bare patches
350 decreases with increasing channel widths (Fig. 6). The unconnected bare patches in Saeftinghe, San Felice and the small
study area of Blackwater, occupy 63 %, 36 % and 67 % of the total number of bare patches, respectively, but only 2 %, 1 %
and 3 % of the total area of bare patches, respectively. Hence, unconnected bare patches are numerous but small. The
number of connected bare patches, in contrast, is in most cases smaller and they become less abundant with increasing width
of the connecting channels.

355 5.2 Topographic conditions determining the revegetation of bare patches in Saeftinghe

The multi-temporal analysis for Saeftinghe shows that bare patches have been dynamically expanding or shrinking between
the four images of 1990-1998-2004-2008 (Fig. 7 Fig. S4). We focused on bare areas with two extreme rates of revegetation,
which are permanent bare areas (which never revegetated throughout the time series) and rapidly revegetated bare areas
(only present in 1998 or 2004 and revegetated by the next time step). The spatial distribution of these bare categories (Fig.
360 7 Fig. S4) suggests that the inner portion of big connected bare patches tends to be stable and never revegetated within the
studied period, while rapidly recovering bare areas are mainly present at the edge of small bare patches.

5.2.1 Elevation

The elevation distribution showed that permanently bare areas (i.e. remaining bare over the studied 18-year period) occupy
the lowest range of elevations, whereas permanent marsh areas have the highest range of elevations (Fig. 8 Fig. 7a). At
365 intermediate elevations, bare patches become rapidly revegetated (i.e. within 4 to 6 years after their first appearance (Fig.
8 Fig. 7a).

5.2.2 Distance to channels

The frequency distribution of the different bare categories with distance to the channels (Fig. 8 Fig. 7b) shows that stable
marshes are closest to channels with a peak around 1-2 m from channels. Bare areas that revegetated quickly have an
370 intermediate distance around 8 m from channels, whereas permanent bare areas are located farthest from the channels with a
peak at 21 m.

5.2.3 Connectivity to channels

Permanent bare areas are always connected to channels, and tend to be associated with wide channels (Fig. 8 Fig. 7c). The
percentage of bare areas that become revegetated increases with decreasing channel width (Fig. 8 Fig. 7c).

Bare patches within otherwise vegetated coastal marshes are often recognized as symptoms of marsh loss in many places around the world (Kearney et al., 1988; Fagherazzi, 2013; Mariotti and Fagherazzi, 2013; Ortiz et al., 2017; Schepers et al., 2017), but comparative studies among different marsh systems to better understand the conditions that determine their presence and potential vegetation recovery are relatively scarce (e.g., Mariotti, 2016). ~~Fig. 9~~[Fig. 8](#) provides a schematic summary of our results, and of our interpretations that are discussed here. For three marsh sites with different tidal ranges, sediment input and plant species, we showed that: (1) bare patches connected to channels occur most frequently at the lowest surface elevations and farthest distances from creeks; unconnected bare patches most frequently occupy intermediate elevations and distances from creeks, and are smaller in size and larger in number; and vegetated marshes dominate at the highest surface elevations and closest to creeks. (2) The elevations of connected and unconnected bare patches tend to be lower relative to the tidal frame in sites with a smaller tidal range, although our analysis only included three sites. (3) Recovery of vegetation in bare patches at the time scale of the last two decades was only observed in the site with high tidal range and high sediment input. Here vegetation recovery is hampered by low surface elevations relative to the local tidal frame, by farther distance from channels and by a high connectivity to the channel network. Below we will further substantiate these findings, and discuss interpretations and potential hypotheses that may explain mechanisms of formation and recovery of bare patches.

6.1 Topographic conditions determining the presence of bare patches

Our results suggest that bare patches exist under qualitatively similar topographic conditions across three different marsh systems. We found that marshes have a higher elevation than bare patches, in accordance with previous studies (DeLaune et al., 1994; Erwin et al., 2006; Wilson et al., 2014). In addition to existing insights, we found that unconnected bare patches are most frequently found at higher elevations and shorter distances from channels as compared to connected bare patches (Figs. 4 and 5). Additionally, we found a positive relationship between patch size and the width of the connecting channel (Fig. 6). These different observations may be interpreted as follows. First, the positive relationship between bare patch size and connecting channel width may be due to the difference in tidal prism (i.e. the total water volume that floods into, and drains out of, the bare patches during a tidal cycle). A larger bare patch implies a larger tidal prism, which means that higher volumes of water are transported into and out of the bare patches. Assuming that most of the water is transported through the connecting channel, a larger tidal prism would be associated with larger channel-forming discharges and therefore wider channels (e.g., Rinaldo et al., 1999b; Kirwan et al., 2008; D'Alpaos et al., 2010; Vandenbruwaene et al., 2013). Secondly, our finding that unconnected bare patches occur most frequently on higher elevations than connected bare patches, may be interpreted as follows. We expect that connected bare patches experience higher incoming and outgoing flood and ebb flow velocities as they are directly connected to the channels. Furthermore, we found that connected bare patches are larger (Fig. 6), and hence we may expect more potential for erosion of surface sediments induced by waves (because of

larger wind fetch length). Wave erosion in interior marsh ponds has been found to be related to the size and wind fetch length of marsh ponds (Mariotti and Fagherazzi, 2013; Mariotti, 2016; Ortiz et al., 2017). Hence larger bare patches are likely to experience more wave-induced erosion and are found in this study to be connected through wider connecting channels, which may facilitate the tidal export of the eroded sediments from connected bare patches, and therefore may explain the lower surface elevation of connected bare patches. In contrast, we hypothesize that unconnected bare patches, which are typically smaller (Fig. 6), may be expected to experience less wave erosion (smaller fetch length) and much weaker flow velocities (as flow is obstructed by surrounding vegetation). With respect to the latter effect, we notice that our classification of unconnected patches may also include patches with connecting channels smaller than 0.5 m but impossible to detect on the aerial images. Nevertheless, also in the case of such small connecting channels < 0.5 m wide, one can expect that drainage of the bare patches after overmarsh tides is much slower, with lower ebb flow velocities, as compared to bare patches with wide connecting channels (up to several tens of meters wide, see Fig. 6), facilitating faster drainage, higher ebb flow velocities and potentially leading to larger tidal export of eroded sediments. Unconnected bare patches were also found to occur most frequently at shorter distances from channels as compared to connected bare patches, and this may facilitate higher sediment supply to unconnected bare patches closer to channels, as suspended sediment concentrations typically decrease with increasing distance from channels (Leonard, 1997; Christiansen et al., 2000; Temmerman et al., 2003b). Therefore, higher sediment supply and lower magnitude of waves and tidal currents in smaller, unconnected bare patches at shorter distance from channels, may facilitate the settlement of suspended sediments and reduce erosion, and as such may explain our finding of higher surface elevations of unconnected bare patches as compared to connected bare patches (Fig. 9Fig. 8). This finding is also in accordance with the model of Mariotti (2016) proposing that, in what is called the “pond collapse regime”, the depth of connected marsh ponds would be larger than the depth of unconnected ponds.

Thirdly, our results indicate that connected bare patches are predominantly located farther away from channels than unconnected bare patches (Fig. 5). One potential explanation is that connected bare patches are generally larger than unconnected bare patches (Fig. 6), so that a larger fraction of the connected bare patches is located at a farther distance from channels. Presence of bare patches in relation to distance from channels has been previously studied on large regional scales (10^2 – 10^4 m) considering only large estuarine channels (Turner and Rao, 1990; Kearney and Rogers, 2010). On a smaller scale (10 – 10^2 m), Redfield (1972) qualitatively reported that big bare patches are located relatively far from channels. Adamowicz and Roman (2005) observed that bare patches were located at around 11 m from the nearest channel in both ditched and unditched marshes in New England. Such a value is similar to that found for Saeftinghe and San Felice, but smaller than the value obtained for Blackwater. The elevation difference between connected and unconnected bare patches probably relates to their difference in distance to channels. Marshes typically have a micro-topography of higher levees along channels and lower depressions farther away from channels as a consequence of progressive suspended sediment deposition during tidal flooding of marshes from channels (e.g., Reed, 1988; Covi and Kneib, 1995; Leonard, 1997; Esselink et al., 1998; Reed et al., 1999; Allen, 2000; Temmerman et al., 2004; D’Alpaos et al., 2007; Bartholdy, 2012). In accordance with this micro-topography, the lower-elevation connected bare patches are located farther away from channels than the

higher-elevation unconnected bare patches (Fig. 9 Fig. 8). This micro-topography of levees close to channels and depressions further away from channels, is often associated with an increasing inundation duration after high tides, and decreasing soil drainage/aeration during low tides, with increasing distance from channels (e.g., Ursino et al., 2004). Also, a modelling study suggested that marsh vegetation expansion can lead to increased inundation time, and as such can feedback on increased stress and chance for vegetation die-back (Brückner et al. 2019). This may be all mechanisms that may contribute to increased chance for occurrence of bare patches within marshes at farther distances from channels.

In addition, the frequency distribution of distance to the closest channels is observed to be exponential for the vegetated marsh surfaces in all three marsh sites, which is analogous to the results by Marani et al. (2003) and holds only for the vegetated marsh surfaces.

Finally, our results demonstrated that the size of bare patches is negatively related to the number of bare patches (Fig. 6).

Such finding has also been observed in other marsh systems (Turner and Rao, 1990; Schepers et al., 2017). This may be indicative for initial formation of many small bare patches that grow and merge together through time, hence leading to a decreasing number of larger patches. This process of merging of initially small bare patches into larger patches has been documented for the Blackwater study site from an analysis of time series of aerial pictures over the period 1938-2010 (Schepers et al., 2017).

In conclusion, we observed qualitatively similar topographic conditions for the presence of bare patches across the three study sites, albeit that the elevations of connected and unconnected bare patches tend to be lower relative to the tidal frame in sites with a smaller tidal range. The latter agrees with earlier findings that micro-tidal marshes have in general a lower surface elevation than macro-tidal marshes (Kirwan et al., 2010; D'Alpaos et al., 2011). Our finding suggests that feedback mechanisms between vegetation and topography are important in regulating the position of the bare patches, and perhaps generalizable across systems. However, we emphasize that our analysis is based on only three study sites, and more research is needed to assess the degree to which this finding is universal.

6.2 Topographic conditions determining the revegetation of bare patches

The comparison between bare patches with two extreme revegetation rates (i.e., permanent bare patches over the studied 18-year period and rapidly revegetated bare patches within 4-6 years) for Saeftinghe suggests that fast revegetation preferentially occurs by expansion of the vegetated edge into small, higher elevation, unconnected bare patches, whereas the central areas of big, lower elevation, connected bare patches tend to remain unvegetated over the considered time period of 18 years. These results are consistent with previous studies. For example, only small bare patches were invaded by vegetation in ditched marshes in Louisiana, although large bare patches were permanent over a study period of 22 years (Turner and Rao, 1990). In several New England marshes, re-establishment of vegetation started within 1-2 years after unconnected bare patches merged with the channel network and became drained (Wilson et al., 2009, 2014). Additionally, some studies find that unconnected bare patches expand and merge quickly, while connected bare patches are relatively stable (Kearney et al., 1988). These disparate observations in different marsh sites may be due to different environmental conditions, such as differences in relative sea level rise, tidal range, and sediment availability (Mariotti, 2016). In a

modelling study, Mariotti (2016) demonstrated that vegetation recovery in marsh ponds is favoured under conditions of slow
475 relative sea level rise, large tidal range, and large inorganic sediment supply.

6.3 Vulnerability for bare patch formation and resilience for bare patch recovery

480 ~~Previous modelling has suggested that pond formation increases and pond recovery decreases in marsh sites that are subject to a lower suspended sediment availability, smaller tidal range, and lower rate of relative sea level rise (RSLR) (Mariotti, 2016). First of all, we want to emphasize that we only investigated three sites, which it is not enough to fully assess the impact of site-differences, such as in tidal range, sediment supply, and rate of RSLR, on RSLR, on occurrence and revegetation of bare patches. Yet we notice that revegetation only occurred at the site (Saeftinghe) with largest tidal range and sediment supply, while it was not observed at the two other sites (San Felice and Blackwater) with smaller tidal range and sediment supply. Although our study includes only three marsh sites, differences in tidal range, sediment availability, and rate of relative sea level rise (RSLR) between the sites allow us to explore their impact on marsh~~
485 ~~vulnerability and resilience in terms of bare patch formation and recovery.~~ As explained in the description of the three study sites (see Sect. 2), the average tidal range and suspended sediment concentrations vary from highest in the Saeftinghe marsh (4.9 m and 30-60 mg/l respectively), intermediate in San Felice marsh (1 m and 10-20 mg/l), to lowest in the Blackwater marshes (0.5 m and 50 mg/l). Long-term RSLR rates in the San Felice and Blackwater marshes are within the same range of
490 3–4 mm/yr, while mean high water level rise in Saeftinghe is 5.7 mm/yr. In Saeftinghe, marsh elevations are mostly above MHWL, while they are mostly below MHWL in San Felice and Blackwater (Fig. 4). Probably this may explain why the proportion of bare surface area is larger in San Felice and Blackwater (34.33 % and 42.58 %, resp.) than in Saeftinghe (15.72 %). ~~This may also suggest that marshes are more prone to presence of bare patches in sites with lower tidal range and sediment availability (Figs. 1-3), and is furthermore in accordance with the modelling study of Mariotti (2016), demonstrating that pond formation increases and pond recovery decreases with decreasing sediment availability, decreasing tidal range, and increasing rate of RSLR. Based on the comparison of our three study sites, we could not draw clear conclusions on the role of RSLR rate.~~

Marsh resilience inferred by revegetation of bare patches was only observed in Saeftinghe where the mean tidal range is 4.9
m, and bare patches have high elevations relative to the tidal frame (average RE = 1.002 for connected bare patches and
500 average RE = 1.02 for unconnected bare patches; Fig. 4). Revegetation of bare patches has been observed in other systems with high tidal ranges (Millette et al., 2010; Wilson et al., 2014), which facilitates well-drained conditions during low tide and enables vegetation regrowth. In contrast, in Blackwater where the mean tidal range is about 0.5 m, bare patches have a much lower elevation relative to the tidal frame, even below the MLWL (average RE = -0.462 for connected bare patches and average RE = 0.284 for unconnected bare patches; Fig. 4), which means that there is no drainage at low tide so that
marsh vegetation cannot recover. Bare patches also tend to be permanent in other systems under low tidal ranges, such as in
505 Louisiana and mid-Atlantic US salt marshes (Wilson et al., 2014; Ortiz et al., 2017). Clearly, the same elevation loss in a marsh with small tidal range will result in a higher increase in tidal inundation frequency and duration, and consequently in

510 more stress on vegetation growth, as compared to a marsh with a large tidal range. Hence, if marsh vegetation and elevation loss occur, it would be easier to recover for marsh vegetation in a higher tidal range environment, such as that of Saeftinghe, as compared to situations with a lower tidal range, such as the Blackwater and San Felice marshes. This interpretation is in agreement with previous studies. Microtidal marshes were reported to be particularly vulnerable to bare patch formation and expansion (Kearney et al., 1988; Mariotti and Fagherazzi, 2013). Marshes with larger tidal ranges also have bare patches but they are generally more dynamically forming and recovering, while the whole marsh system is relatively stable (Redfield, 1972; Wilson et al., 2009). The model of Kirwan and Guntenspergen (2010) suggested that extensive bare patches occur, expand quickly and become permanent under small tidal ranges but not under large tidal ranges, because the elevation range suitable for vegetation growth is smaller in low tidal range environments. In general, marsh stability is positively related to tidal range (Kirwan et al., 2010; D'Alpaos, 2011), and numerical modelling indicates that high sediment concentrations are necessary for recovery of bare patches (Mariotti, 2016). However, Mariotti (2016) only considers recovery after connection to the tidal channel network, not the recovery of isolated bare patches. In our study, we observed that bare patches unconnected to the tidal channel network all recovered at the Saeftinghe site. Complete drainage of the Saeftinghe bare patches during ebb tides might explain this apparent discrepancy. We suggest that the close distance to channels (see Figs. 5 and 8b) (e.g., Ursino et al., 2004) and coarser sediment associated with channel levees (Allen, 2000) enable the unconnected bare patches to drain completely in Saeftinghe through subsurface drainage, and allow vegetation recovery.

525 Finally, our results may be indicative to decision makers on salt marsh management, as the formation of bare patches may be indicative for marsh degradation towards an unvegetated state, that state that may be difficult to recover. Our study indicates that early signatures for marsh degradation must be particularly monitored in marsh portions, farthest away from main channels and with lowest surface elevations. Monitoring of early signatures is especially advised in systems with very low tidal range and suspended sediment availability.

6.4 Marshes and bare patches as two stable states in intertidal areas

530 ~~High elevation, vegetated marsh areas and low elevation, unvegetated tidal flat areas have been previously identified as alternative stable states, primarily at the large scale of whole tidal basins (km²) (e.g., Fagherazzi et al., 2006; Marani et al., 2007, 2010; McGlathery et al., 2013; Wang and Temmerman, 2013; Moffett et al., 2015). Field evidence for alternative stable state behavior is however scarce for the smaller scale of vegetated and bare patches within marshes (10–100 m²).~~

535 ~~Here we evaluate and discuss our results within the framework of alternative stable state theory, based on the conceptual model presented in Fig. 9. The state variables that are considered are (1) the vegetation biomass (high for the vegetated state versus zero for the bare state) and (2) the surface elevation. According to field studies and models of marsh evolution, both state variables interact through feedback loops: presence of vegetation will promote accretion of mineral and organic sediments, therefore resulting in higher surface elevation (Neumeier and Amos, 2006; Mudd et al., 2010; Vandenbruwaene et al., 2011; Yang et al., 2012; Hu et al., 2014), which further stimulates biomass production, up to a point where a high equilibrium elevation is reached where biomass productivity reaches an optimum (Morris et al., 2002; D'Alpaos et al., 2007;~~

540 Kirwan and Murray, 2007; Marani et al., 2007, 2010); while vice versa, absence of vegetation will facilitate erosion induced
by tidal currents and waves, hence lowering surface elevation, which further prohibits vegetation growth, until a low bare
equilibrium elevation is reached where erosion and deposition by waves and tidal currents are in dynamic equilibrium
(Fagherazzi et al., 2006; Kirwan and Murray, 2007; Marani et al., 2007, 2010). Such internal feedback loops between
545 vegetation and elevation, i.e. the state variables, are indicated along the Y axis in Fig. 9b, and can lead to two stable states, a
high elevation vegetated state (the green curve in Fig. 9b) or low elevation bare state (the grey curve in Fig. 9b). The
conditions under which these two states can develop, are indicated along the X axis of Fig. 9b, such as the sediment supply
and level of soil drainage/aeration. Previous modelling studies have shown that high sediment supply and low sea level rise
rates (i.e. allowing long periods of intertidal soil drainage and aeration) favors the evolution towards a high vegetated state,
while low sediment supply and high sea level rise rates (i.e. leading to shorter periods of intertidal soil drainage and aeration)
550 promote formation of the low bare state (Marani et al., 2007, 2010; Mariotti, 2016). Within the context of our analysis, such
conditions are spatially varying within marshes as a function of distance from tidal channels, as both sediment supply (e.g.,
Christiansen et al., 2000; Temmerman et al., 2003a) and soil drainage/aeration (e.g., Ursino et al., 2004) typically decrease
with increasing distance from channels. Additionally, connectivity of bare patches to the channel network may further
influence sediment import into or export from bare patches, as discussed before.

555 Hence, we propose that at short distance from channels, conditions are favorable for a high, vegetated stable state, while at
large distance from channels, conditions are favorable for a low, bare stable state; and that at intermediate distances from
channels, both vegetated and bare states may co-exist next to each other as alternative stable states (Fig. 9a and b). Our
results provide suggestive support for this hypothesis and conceptual framework, although we emphasize that our discussion
here is very indicative and far from conclusive. Our results indicate that the occurrence of vegetated marshes peaks at the
560 highest elevations (Fig. 4) and shortest distances from channels (Fig. 5), which may be indicative for the high elevation
vegetated stable state (left part of Fig. 9b). Connected bare patches peak at the lowest elevation (Fig. 4) and farthest
distances from channels (Fig. 5), and may be indicative for the low elevation bare stable state (right part of Fig. 9b). At
intermediate distances from channels, vegetated and bare marsh portions can exist next to each other (Fig. 5; and see e.g. Fig.
3). In the one field site with observed vegetation recovery (Saeftinghe), the bare areas that were never revegetated in the
565 studied 18 year period predominantly occurred in large connected bare patches at the lowest elevations and farthest distances
from channels (Fig. 8). This may be indicative for the existence of bare stable states that are most frequently found at far
distances from channels (right part of graph in Fig. 9b). In that same field site, fast revegetation (within 4-6 years) was
predominantly observed in small unconnected bare patches at intermediate elevations and intermediate distances from
channels (Figs. 7 and 8). This may be indicative for disturbances of vegetation cover that quickly recover to the vegetated
570 state, at intermediate distances (left part of graph in Fig. 9b). Comparison between the three study sites, suggests that
conditions for the bare stable state are more likely in sites with lower tidal range (i.e. less soil drainage/aeration) and low
sediment supply, as revegetation was not observed in the two sites with lowest tidal range and sediment supply (Blackwater
and San Felice), and only occurred in the site with highest tidal range and sediment supply (Saeftinghe) (Fig. 9b).

575 ~~Finally, we note that the bio-geomorphic feedback mechanisms (i.e. vegetation-elevation feedbacks) leading to the elevated marsh state are similar in large-scale or small-scale studies, but there might be differences leading to the low, unvegetated state. In large lagoons or extensive tidal basins, the low elevation of the tidal flats is mainly maintained by feedbacks between wave erosion and water depth (Fagherazzi et al., 2006; Mariotti and Fagherazzi, 2010). Wave-induced erosion is, however, expected to be a negligible process in small bare patches such as those of Saeftinghe (max. 60 m wind fetch length) that completely drain at ebb tide. Hence wave erosion is not likely to be a relevant mechanism controlling the depth of small~~
580 ~~bare patches or the evolution of small bare patches in Saeftinghe. Our results indicate that a connection with the tidal channel system is associated with a lower elevation of the bare patches, probably because of higher flow velocities that decrease sedimentation rates or cause erosion. Other work suggests that the substrate of bare ponds within marshes might be eroded and exported through the connecting channels (e.g., Day et al., 2011; Schepers et al., 2017). Nevertheless, further research should test these hypotheses.~~

585 7 Conclusions

In this paper, we studied the topographical conditions for presence and revegetation of bare patches within three coastal marsh sites that are largely different in tidal range, sediment supply and plant species. The analyses of aerial photographs, LIDAR data and field topographic measurements showed that the topographic conditions (i.e. elevations, distances from channels and connectivity to channels) for presence of bare patches were qualitatively consistent among the three marsh sites.

590 We found that bare patches connected to channels occur most frequently at the lowest surface elevations and farthest away from creeks; unconnected bare patches most frequently occupy intermediate elevations and distances from creeks, and are smaller in size and larger in number; and vegetated marshes dominate at the highest surface elevations and closest to creeks. Further, we showed that the elevations of connected and unconnected bare patches tend to be lower relative to the tidal frame with increasing tidal range, although our analysis only included three sites. Revegetation of bare patches was only observed
595 in one site, which was the site with the highest tidal range and the largest sediment supply. For that site, we found that the chance of bare patch revegetation decreases with increasing width of channels that connect bare patches to the tidal channel network. The latter is associated with lower bare patch elevation, farther distance to channels and bigger bare patch size. Finally, in the context of sea level rise, our results suggest that the marsh site with the highest tidal range and highest sediment input is less vulnerable to bare patch formation and more resilient in terms of revegetation of bare patches than the
600 two other marsh sites with lower tidal range and lower sediment supply. However, we emphasize that our study only included three sites, and that further research comparing much more sites is needed, to further advance our understanding of why certain marsh sites are more vulnerable to others to formation and persistence of bare patches. Such knowledge will be important to inform decision makers on site-specific priorities for marsh conservation.

Data availability

605 The aerial images and DTM data for Saeftinghe can be downloaded from Rijkswaterstaat (<https://geoservices.rijkswaterstaat.nl>). The IKONOS data for San Felice can be accessed at Planetek Italia s.r.l. (<https://www.planetek.it>). The field elevation data for San Felice can be accessed at Venice Water Authority in Italy (<http://provveditoratovenezia.mit.gov.it>). The aerial images for Blackwater can be downloaded from earthexplorer.usgs.gov. The LIDAR data for Blackwater can be downloaded from <https://inport.nmfs.noaa.gov/inport/item/49781>.

610 Supplement link

The supplement related to this article is available online at:

Author contribution

ST and CW designed the study. CW prepared the manuscript with contributions from all co-authors.

Competing interests

615 The authors declare that they have no conflict of interest.

Acknowledgements

We dedicate this paper to the late Jose Busnelli who contributed to parts of the analyses presented in this paper. We thank Rijkswaterstaat (Dick De jong) in the Netherlands, Venice Water Authority in Italy, and United States Department of Agriculture and Maryland Department of Natural Resources in the USA for providing vegetation maps, aerial photographs,
620 as well as elevation data of LIDAR surveys and field surveys. We also would like to thank Glenn Guntenspergen, Melissa Duvall, Patrick Brennand and Kyle Derby for the field measurements in the Blackwater Marshes.

Financial support

This work has been supported by the Project 41501116, 51761135022, and 41401413 supported by National Natural Science Foundation of China, the European Union Programme Erasmus Mundus External Cooperation Window (EMECW)–Lot 14–
625 China and by the FWO research community “Functioning of River Ecosystems by Plant-Flow-Soil interactions”, Research Foundation Flanders (FWO PhD grant L.S., 11S9614N), the Technology Foundation for Selected Overseas Chinese Scholar from Ministry of Personnel of China (2015), and the U.S. National Science Foundation (1237733, 1426981, 1654374, 1832221 and 1529245).

References

- 630 Adamowicz, S. C. and Roman, C. T.: New England salt marsh pools: a quantitative analysis of geomorphic and geographic features, *Wetlands*, 25, 279–288, <https://doi.org/10.1672/4, 2005>.
- Alkemade, I. S. W.: *Kwaliteitsdocument Laseraltimetrie, Projectgebied Westerschelde, Ministerie van Verkeer en Waterstaat, Rijkswaterstaat, Delft, NL, 2004.*
- Allen, J.: Morphodynamics of Holocene salt marshes: a review sketch from the Atlantic and Southern North Sea coasts of Europe, *Quat. Sci. Rev.*, 19, 1155–1231, [https://doi.org/10.1016/s0277-3791\(99\)00034-7, 2000](https://doi.org/10.1016/s0277-3791(99)00034-7, 2000).
- 635 Argow, B. A. and FitzGerald, D. M.: Winter processes on northern salt marshes: evaluating the impact of in-situ peat compaction due to ice loading, Wells, ME, *Estuar. Coast. Shelf Sci.*, 69, 360–369, <https://doi.org/10.1016/j.ecss.2006.05.006, 2006>.
- Bakker, R. B. and Bijkerk, W.: *Toelichting Bij de Geomorfologische Kartering Westerschelde 2008 op Basis van False Colour-Luchtfoto's 1:10.000. Report, Ministerie van Verkeer en Waterstaat, Rijksinstituut, Adviesdienst Geo-Informatie & ICT, Den Haag, Delft, NL, 2009.*
- 640 Barbier, E. B., Hacker, S. D., Kennedy, C., Koch, E. W., Stier, A. C., and Silliman, B. R.: The value of estuarine and coastal ecosystem services, *Ecol. Monogr.*, 81, 169–193, <https://doi.org/10.1890/10-1510.1, 2011>.
- Barbier, E. B., Koch, E. W., Silliman, B. R., Hacker, S. D., Wolanski, E., Primavera, J., Granek, E. F., Polasky, S., Aswani, S., Cramer, L. A., Stoms, D. M., Kennedy, C. J., Bael, D., Kappel, C. V., Perillo, G. M. E., and Reed, D. J.: Coastal ecosystem-based management with nonlinear ecological functions and values, *Science*, 319, 321–323, <https://doi.org/10.1126/science.1150349, 2008>.
- 645 Bartholdy, J.: Salt marsh sedimentation, in: *Principles of Tidal Sedimentology*, edited by: Davis, R. A. and Dalrymple, R. W., Springer, Dordrecht, NL, 151–185, 2012.
- 650 Baumann, R. H., Day, J. W., and Miller, C. A.: Mississippi deltaic wetland survival: sedimentation versus coastal submergence, *Science*, 224, 1093–1095, <https://doi.org/10.1126/science.224.4653.1093, 1984>.
- Belluco, E., Camuffo, M., Ferrari, S., Modenese, L., Silvestri, S., Marani, A., and Marani, M.: Mapping salt-marsh vegetation by multispectral and hyperspectral remote sensing, *Remote Sens. Environ.*, 105, 54–67, <https://doi.org/10.1016/j.rse.2006.06.006, 2006>.
- 655 [Brückner, M. Z. M., Schwarz, C., van Dijk, W. M., van Oorschot, M., Douma, H., and Kleinhans, M. G.: Salt marsh establishment and eco-engineering effects in dynamic estuaries determined by species growth and mortality, *J. Geophys. Res.: Earth Surf.*, 124, 2962–2986, <https://doi.org/10.1029/2019JF005092, 2019>.](https://doi.org/10.1029/2019JF005092, 2019)
- Carbognin, L., Teatini, P., and Tosi, L.: Eustacy and land subsidence in the Venice Lagoon at the beginning of the new millennium, *J. Mar. Syst.*, 51, 345–353, <https://doi.org/10.1016/j.jmarsys.2004.05.021, 2004>.
- 660 Carniello, L., Defina, A., and D'Alpaos, L.: Morphological evolution of the Venice Lagoon: evidence from the past and trend for the future, *J. Geophys. Res.*, 114, F04002, <https://doi.org/10.1029/2008jf001157, 2009>.

- Christiansen, T., Wiberg, P. L., and Milligan, T. G.: Flow and sediment transport on a tidal salt marsh surface, *Estuar. Coast. Shelf Sci.*, 50, 315–331, <https://doi.org/10.1006/ecss.2000.0548>, 2000.
- 665 Covi, M. P. and Kneib, R. T.: Intertidal distribution, population dynamics and production of the amphipod *Uthlorchestia spartinophila* in a Georgia, USA, salt marsh, *Mar. Biol.*, 121, 447–455, <https://doi.org/10.1007/bf00349453>, 1995.
- D'Alpaos, A.: The mutual influence of biotic and abiotic components on the long-term ecomorphodynamic evolution of salt-marsh ecosystems, *Geomorphology*, 126, 269–278, <https://doi.org/10.1016/j.geomorph.2010.04.027>, 2011.
- 670 D'Alpaos, A., Lanzoni, S., Marani, M., and Rinaldo, A.: Landscape evolution in tidal embayments: modeling the interplay of erosion, sedimentation, and vegetation dynamics, *J. Geophys. Res.*, 112, F01008, <https://doi.org/10.1029/2006jf000537>, 2007.
- D'Alpaos, A., Lanzoni, S., Marani, M., and Rinaldo, A.: On the tidal prism–channel area relations, *J. Geophys. Res.*, 115, F01003, <https://doi.org/10.1029/2008jf001243>, 2010.
- 675 D'Alpaos, A., Mudd, S. M., and Carniello, L.: Dynamic response of marshes to perturbations in suspended sediment concentrations and rates of relative sea level rise, *J. Geophys. Res.*, 116, F04020, <https://doi.org/10.1029/2011jf002093>, 2011.
- D'Alpaos, A. and Marani, M.: Reading the signatures of biologic–geomorphic feedbacks in salt-marsh landscapes, *Adv. Water Resour.*, 93, 265–275, <https://doi.org/10.1016/j.advwatres.2015.09.004>, 2016.
- 680 Day, J. W., Britsch, L. D., Hawes, S. R., Shaffer, G. P., Reed, D. J., and Cahoon, D.: Pattern and process of land loss in the Mississippi delta: a spatial and temporal analysis of wetland habitat change, *Estuaries*, 23, 425–438, <https://doi.org/10.2307/1353136>, 2000.
- Day, J. W., Kemp, G. P., Reed, D. J., Cahoon, D. R., Boumans, R. M., Suhayda, J. M., and Gambrell, R.: Vegetation death and rapid loss of surface elevation in two contrasting Mississippi delta salt marshes: the role of sedimentation, autocompaction and sea-level rise, *Ecol. Eng.*, 37, 229–240, <https://doi.org/10.1016/j.ecoleng.2010.11.021>, 2011.
- 685 Day, J. W., Rybczyk, J., Scarton, F., Rismondo, A., Are, D., and Cecconi, G.: Soil accretionary dynamics, sea-level rise and the survival of wetlands in Venice Lagoon: a field and modelling approach, *Estuar. Coast. Shelf Sci.*, 49, 607–628, <https://doi.org/10.1006/ecss.1999.0522>, 1999.
- Defendi, V., Kovačević, V., Arena, F., and Zaggia, L.: Estimating sediment transport from acoustic measurements in the Venice Lagoon inlets, *Cont. Shelf Res.*, 30, 883–893, <https://doi.org/10.1016/j.csr.2009.12.004>, 2010.
- 690 DeLaune, R. D., Nyman, J. A., and Patrick, J. W. H.: Peat collapse, ponding and wetland loss in a rapidly submerging coastal marsh, *J. Coast. Res.*, 10, 1021–1030, 1994.
- Elschot, K., Vermeulen, A., Vandenbruwaene, W., Bakker, J. P., Bouma, T. J., Stahl, J., Castelijn, H., and Temmerman, S.: Top-down vs. bottom-up control on vegetation composition in a tidal marsh depends on scale, *PLoS One*, 12, e0169960, <https://doi.org/10.1371/journal.pone.0169960>, 2017.

- Erwin, R. M., Cahoon, D. R., Prosser, D. J., Sanders, G. M., and Hensel, P.: Surface elevation dynamics in vegetated *Spartina* marshes versus unvegetated tidal ponds along the Mid-Atlantic Coast, USA, with implications to waterbirds, *Estuaries Coasts*, 29, 96–106, <https://doi.org/10.1007/bf02784702>, 2006.
- Esselink, P., Dijkema, K. S., Sabine, R., and Geert, H.: Vertical accretion and profile changes in abandoned man-made tidal marshes in the Dollard Estuary, the Netherlands, *J. Coast. Res.*, 14, 570–582, 1998.
- Fagherazzi, S.: The ephemeral life of a salt marsh, *Geology*, 41, 943–944, <https://doi.org/10.1130/focus082013.1>, 2013.
- 700 Fagherazzi, S., Carniello, L., D'Alpaos, L., and Defina, A.: Critical bifurcation of shallow microtidal landforms in tidal flats and salt marshes, *Proc. Natl. Acad. Sci. U. S. A.*, 103, 8337–8341, <https://doi.org/10.1073/pnas.0508379103>, 2006.
- Fagherazzi, S., Kirwan, M. L., Mudd, S. M., Guntenspergen, G. R., Temmerman, S., D'Alpaos, A., van de Koppel, J., Rybczyk, J. M., Reyes, E., Craft, C., and Clough, J.: Numerical models of salt marsh evolution: ecological, geomorphic, and climatic factors, *Rev. Geophys.*, 50, 294–295, <https://doi.org/10.1029/2011rg000359>, 2012.
- 705 Fagherazzi, S., Mariotti, G., Wiberg, P., and McGlathery, K.: Marsh collapse does not require sea level rise, *Oceanography*, 26, 70–77, <https://doi.org/10.5670/oceanog.2013.47>, 2013.
- Fagherazzi, S., Palermo, C., Rulli, M. C., Carniello, L., and Defina, A.: Wind waves in shallow microtidal basins and the dynamic equilibrium of tidal flats, *J. Geophys. Res.*, 112, F02024, <https://doi.org/10.1029/2006jf000572>, 2007.
- Ganju, N. K., Nidzieko, N. J., and Kirwan, M. L.: Inferring tidal wetland stability from channel sediment fluxes: observations and a conceptual model, *J. Geophys. Res.: Earth Surf.*, 118, 2045–2058, <https://doi.org/10.1002/jgrf.20143>, 2013.
- Gedan, K. B., Kirwan, M. L., Wolanski, E., Barbier, E. B., and Silliman, B. R.: The present and future role of coastal wetland vegetation in protecting shorelines: answering recent challenges to the paradigm, *Clim. Change*, 106, 7–29, <https://doi.org/10.1007/s10584-010-0003-7>, 2011.
- 715 Gieskes, J. M., Elwany, H., Rasmussen, L., Han, S., Rathburn, A., and Deheyn, D. D.: Salinity variations in the Venice Lagoon, Italy: results from the SIOSED project, May 2005–February 2007, *Mar. Chem.*, 154, 77–86, <https://doi.org/10.1016/j.marchem.2013.05.011>, 2013.
- Harshberger, J. W.: The origin and vegetation of salt marsh pools, *Proc. Am. Philos. Soc.*, 55, 481–484, 1916.
- Hu, Z., Suzuki, T., Zitman, T., Uittewaal, W., and Stive, M.: Laboratory study on wave dissipation by vegetation in combined current–wave flow, *Coast. Eng.*, 88, 131–142, <https://doi.org/10.1016/j.coastaleng.2014.02.009>, 2014.
- 720 Hu, Z., van Belzen, J., van der Wal, D., Balke, T., Wang, Z. B., Stive, M., and Bouma, T. J.: Windows of opportunity for salt marsh vegetation establishment on bare tidal flats: the importance of temporal and spatial variability in hydrodynamic forcing, *J. Geophys. Res.: Biogeosciences*, 120, 1450–1469, <https://doi.org/10.1002/2014jg002870>, 2015a.
- 725 Hu, Z., Wang, Z. B., Zitman, T. J., Stive, M. J. F., and Bouma, T. J.: Predicting long-term and short-term tidal flat morphodynamics using a dynamic equilibrium theory, *J. Geophys. Res.: Earth Surf.*, 120, 1803–1823, <https://doi.org/10.1002/2015jf003486>, 2015b.

- Huijs, S. W. E.: Geomorfologische Ontwikkeling van Het Intergetijdgebied in de Westerschelde, 1935-1989. Rapport R 95-3, Universiteit Utrecht, Utrecht, NL, 1995.
- 730 Kearney, M. S., Grace, R. E., and Stevenson, J. C.: Marsh loss in nanticoke estuary, Chesapeake Bay, *Geogr. Rev.*, 78, 205–220, <https://doi.org/10.2307/214178>, 1988.
- Kearney, M. S. and Rogers, A. S.: Forecasting sites of future coastal marsh loss using topographical relationships and logistic regression, *Wetl. Ecol. Manag.*, 18, 449–461, <https://doi.org/10.1007/s11273-010-9178-y>, 2010.
- 735 Kearney, M. S., Rogers, A. S., Townshend, J. R. G., Rizzo, E., Stutzer, D., Stevenson, J. C., and Sundborg, K.: Landsat imagery shows decline of coastal marshes in Chesapeake and Delaware Bays, *Eos Trans. Am. Geophys. Union*, 83, 173–178, <https://doi.org/10.1029/2002eo000112>, 2002.
- Kendrot, S. R.: Restoration through eradication: protecting Chesapeake Bay marshlands from invasive nutria (*Myocastor coypus*), in: *Island Invasives: Eradication and Management. Proceedings of the International Conference on Island Invasives (Occasional Papers of the IUCN Species Survival Commission Occasional Papers of the IUCN Species Survival Commission)*, edited by: Veitch, C. R., Clout, M. N., and Towns, D. R., IUCN, Gland, Switzerland, 313–319, 2011.
- 740 Kirwan, M. L. and Guntenspergen, G. R.: Influence of tidal range on the stability of coastal marshland, *J. Geophys. Res.: Earth Surf.*, 115, F02009, <https://doi.org/10.1029/2009jf001400>, 2010.
- Kirwan, M. L. and Guntenspergen, G. R.: Feedbacks between inundation, root production, and shoot growth in a rapidly submerging brackish marsh, *J. Ecol.*, 100, 764–770, <https://doi.org/10.1111/j.1365-2745.2012.01957.x>, 2012.
- 745 Kirwan, M. L. and Guntenspergen, G. R.: Response of plant productivity to experimental flooding in a stable and a submerging marsh, *Ecosystems*, 18, 903–913, <https://doi.org/10.1007/s10021-015-9870-0>, 2015.
- Kirwan, M. L., Guntenspergen, G. R., D'Alpaos, A., Morris, J. T., Mudd, S. M., and Temmerman, S.: Limits on the adaptability of coastal marshes to rising sea level, *Geophys. Res. Lett.*, 37, 58–94, <https://doi.org/10.1029/2010gl045489>, 2010.
- 750 Kirwan, M. L. and Megonigal, J. P.: Tidal wetland stability in the face of human impacts and sea-level rise, *Nature*, 504, 53–60, <https://doi.org/10.1038/nature12856>, 2013.
- Kirwan, M. L. and Murray, A. B.: A coupled geomorphic and ecological model of tidal marsh evolution, *Proc. Natl. Acad. Sci. U. S. A.*, 104, 6118–6122, <https://doi.org/10.1073/pnas.0700958104>, 2007.
- 755 Kirwan, M. L., Murray, A. B., and Boyd, W. S.: Temporary vegetation disturbance as an explanation for permanent loss of tidal wetlands, *Geophys. Res. Lett.*, 35, L05403, <https://doi.org/10.1029/2007gl032681>, 2008.
- Kirwan, M. L., Murray, A. B., Donnelly, J. P., and Corbett, D. R.: Rapid wetland expansion during European settlement and its implication for marsh survival under modern sediment delivery rates, *Geology*, 39, 507–510, <https://doi.org/10.1130/g31789.1>, 2011.
- 760 Leonard, L. A.: Controls of sediment transport and deposition in an incised mainland marsh basin, Southeastern North Carolina, *Wetlands*, 17, 263–274, <https://doi.org/10.1007/bf03161414>, 1997.

- Marani, M., Belluco, E., D'Alpaos, A., Defina, A., Lanzoni, S., and Rinaldo, A.: On the drainage density of tidal networks, *Water Resour. Res.*, 39, 5029–5035, <https://doi.org/10.1029/2001wr001051>, 2003.
- 765 Marani, M., D'Alpaos, A., Lanzoni, S., Carniello, L., and Rinaldo, A.: Biologically-controlled multiple equilibria of tidal landforms and the fate of the Venice Lagoon, *Geophys. Res. Lett.*, 34, 224–238, <https://doi.org/10.1029/2007gl030178>, 2007.
- Marani, M., D'Alpaos, A., Lanzoni, S., Carniello, L., and Rinaldo, A.: The importance of being coupled: stable states and catastrophic shifts in tidal biomorphodynamics, *J. Geophys. Res.*, 115, F04004, <https://doi.org/10.1029/2009jf001600>, 2010.
- 770 Marani, M., Silvestri, S., Belluco, E., Ursino, N., Comerlati, A., Tosatto, O., and Putti, M.: Spatial organization and ecohydrological interactions in oxygen-limited vegetation ecosystems, *Water Resour. Res.*, 42, 387–403, <https://doi.org/10.1029/2005wr004582>, 2006.
- Mariotti, G.: Revisiting salt marsh resilience to sea level rise: are ponds responsible for permanent land loss? *J. Geophys. Res.: Earth Surf.*, 121, 1391–1407, <https://doi.org/10.1002/2016jf003900>, 2016.
- 775 Mariotti, G. and Fagherazzi, S.: A numerical model for the coupled long-term evolution of salt marshes and tidal flats, *J. Geophys. Res.*, 115, F01004, <https://doi.org/10.1029/2009jf001326>, 2010.
- Mariotti, G. and Fagherazzi, S.: Critical width of tidal flats triggers marsh collapse in the absence of sea-level rise, *Proc. Natl. Acad. Sci. U. S. A.*, 110, 5353–5356, <https://doi.org/10.1073/pnas.1219600110>, 2013.
- 780 McGlathery, K., Reidenbach, M., D'Odorico, P., Fagherazzi, S., Pace, M., and Porter, J.: Nonlinear dynamics and alternative stable states in shallow coastal systems, *Oceanography*, 26, 220–231, <https://doi.org/10.5670/oceanog.2013.66>, 2013.
- McLeod, E., Chmura, G. L., Bouillon, S., Salm, R., Björk, M., Duarte, C. M., Lovelock, C. E., Schlesinger, W. H., and Silliman, B. R.: A blueprint for blue carbon: toward an improved understanding of the role of vegetated coastal habitats in sequestering CO₂, *Front. Ecol. Environ.*, 9, 552–560, <https://doi.org/10.1890/110004>, 2011.
- 785 Miller, W. R. and Egler, F. E.: Vegetation of the wequetequock-pawcatuck tidal-marshes, Connecticut, *Ecol. Monogr.*, 20, 143–172, <https://doi.org/10.2307/1943548>, 1950.
- Millette, T. L., Argow, B. A., Marcano, E., Hayward, C., Hopkinson, C. S., and Valentine, V.: Salt marsh geomorphological analyses via integration of multitemporal multispectral remote sensing with LIDAR and GIS, *J. Coast. Res.*, 265, 809–816, <https://doi.org/10.2112/jcoastres-d-09-00101.1>, 2010.
- 790 Moffett, K., Nardin, W., Silvestri, S., Wang, C., and Temmerman, S.: Multiple stable states and catastrophic shifts in coastal wetlands: progress, challenges, and opportunities in validating theory using remote sensing and other methods, *Remote Sens.*, 7, 10184–10226, <https://doi.org/10.3390/rs70810184>, 2015.
- Morris, J. T., Sundareshwar, P. V., Nietch, C. T., Kjerfve, B., and Cahoon, D. R.: Responses of coastal wetlands to rising sea level, *Ecology*, 83, 2869–2877, [https://doi.org/10.1890/0012-9658\(2002\)083\[2869:rocwtr\]2.0.co;2](https://doi.org/10.1890/0012-9658(2002)083[2869:rocwtr]2.0.co;2), 2002.

- 795 Morton, R. A., Tiling, G., and Ferina, N. F.: Causes of hot-spot wetland loss in the Mississippi delta plain, *Environ. Geosci.*, 10, 71–80, <https://doi.org/10.1306/eg100202007>, 2003.
- Muskalski, S. M. and Sommerfield, C. K.: Suspended sediment deposition and trapping efficiency in a Delaware salt marsh, *Geomorphology*, 139–140, 195–204, <https://doi.org/10.1016/j.geomorph.2011.10.018>, 2012.
- 800 Mudd, S. M., D'Alpaos, A., and Morris, J. T.: How does vegetation affect sedimentation on tidal marshes? Investigating particle capture and hydrodynamic controls on biologically mediated sedimentation, *J. Geophys. Res.*, 115, F03029, <https://doi.org/10.1029/2009jf001566>, 2010.
- Neumeier, U. R. S. and Amos, C. L.: The influence of vegetation on turbulence and flow velocities in European salt-marshes, *Sedimentology*, 53, 259–277, <https://doi.org/10.1111/j.1365-3091.2006.00772.x>, 2006.
- 805 Nyman, J. A., Walters, R. J., Delaune, R. D., and Patrick, W. H.: Marsh vertical accretion via vegetative growth, *Estuar. Coast. Shelf Sci.*, 69, 370–380, <https://doi.org/10.1016/j.ecss.2006.05.041>, 2006.
- Ortiz, A. C., Roy, S., and Edmonds, D. A.: Land loss by pond expansion on the Mississippi river delta plain, *Geophys. Res. Lett.*, 44, 3635–3642, <https://doi.org/10.1002/2017gl073079>, 2017.
- Pendleton, E. and Stevenson, J.: Investigations of Marsh Losses at Blackwater Refuge: Final Report, University of Maryland Center for Environmental and Estuarine Studies, Cambridge, MA, 1983.
- 810 Penland, S., Wayne, L., Britsch, L. D., Williams, S. J., Beall, A. D., and Butterworth, V. C.: Geomorphic Classification of Coastal Land Loss between 1932 and 1990 in the Mississippi River Delta Plain, Southeastern Louisiana. USGS Open File Report 00-417, U.S. Geological Survey, Coastal Marine Geology Program, Woods Hole, MA, 2000.
- Redfield, A. C.: Development of a New England salt marsh, *Ecol. Monogr.*, 42, 201–237, <https://doi.org/10.2307/1942263>, 1972.
- 815 Reed, D. J.: Sediment dynamics and deposition in a retreating coastal salt marsh, *Estuar. Coast. Shelf Sci.*, 26, 67–79, [https://doi.org/10.1016/0272-7714\(88\)90012-1](https://doi.org/10.1016/0272-7714(88)90012-1), 1988.
- Reed, D. J., Spencer, T., Murray, A. L., French, J. R., and Leonard, L.: Marsh surface sediment deposition and the role of tidal creeks: implications for created and managed coastal marshes, *J. Coast. Conserv.*, 5, 81–90, <https://doi.org/10.1007/bf02802742>, 1999.
- 820 Reitsma, J. M.: Toelichting Bij de Vegetatiekartering Westerschelde 2004 op Basis Vanfalse Colour-Luchtfoto's 1:5000/1:10000, Report, Ministerie van Verkeer en Waterstaat, Rijksinstituut, Adviesdienst Geo-Informatie & ICT, Den Haag, Delft, NL, 2006.
- Rinaldo, A., Fagherazzi, S., Lanzoni, S., Marani, M., and Dietrich, W. E.: Tidal networks: 2. Watershed delineation and comparative network morphology, *Water Resour. Res.*, 35, 3905–3917, <https://doi.org/10.1029/1999wr900237>, 1999a.
- 825 Rinaldo, A., Fagherazzi, S., Lanzoni, S., Marani, M., and Dietrich, W. E.: Tidal networks: 3. Landscape-forming discharges and studies in empirical geomorphic relationships, *Water Resour. Res.*, 35, 3919–3929, <https://doi.org/10.1029/1999wr900238>, 1999b.

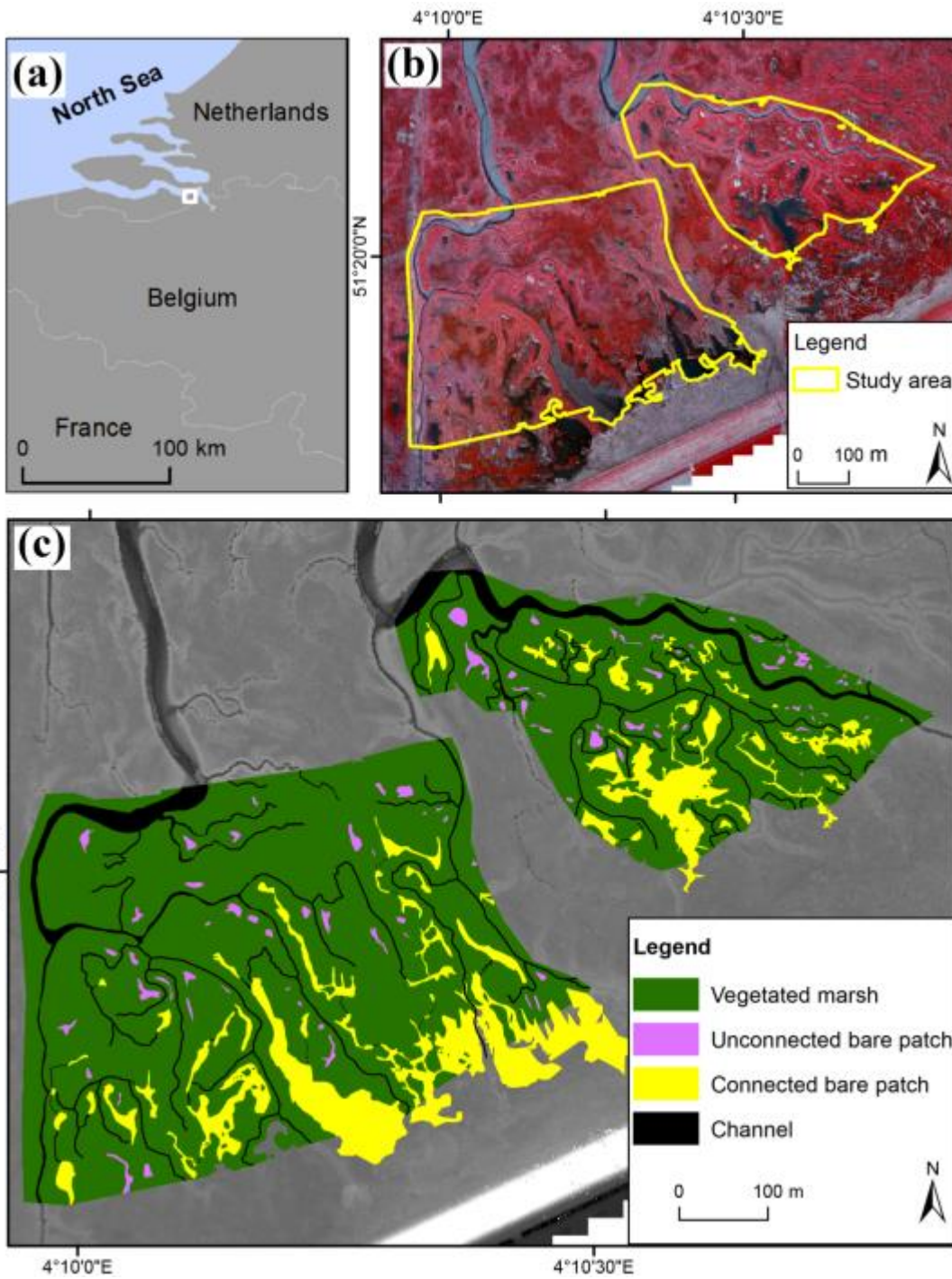
- Roner, M., D'Alpaos, A., Ghinassi, M., Marani, M., Silvestri, S., Franceschinis, E., and Realdon, N.: Spatial variation of salt-marsh organic and inorganic deposition and organic carbon accumulation: inferences from the Venice Lagoon, Italy, *Adv. Water Resour.*, 93, 276–287, <https://doi.org/10.1016/j.advwatres.2015.11.011>, 2016.
- Sarretta, A., Pillon, S., Molinaroli, E., Guerzoni, S., and Fontolan, G.: Sediment budget in the Lagoon of Venice, Italy, *Cont. Shelf Res.*, 30, 934–949, <https://doi.org/10.1016/j.csr.2009.07.002>, 2010.
- Scheffer, M., Carpenter, S., Foley, J. A., Folke, C., and Walker, B.: Catastrophic shifts in ecosystems, *Nature*, 413, 591–596, <https://doi.org/10.1038/35098000>, 2001.
- Scheffer, M. and Carpenter, S. R.: Catastrophic regime shifts in ecosystems: linking theory to observation, *Trends Ecol. Evol.*, 18, 648–656, <https://doi.org/10.1016/j.tree.2003.09.002>, 2003.
- Schepers, L., Kirwan, M., Guntenspergen, G., and Temmerman, S.: Spatio-temporal development of vegetation die-off in a submerging coastal marsh, *Limnol. Oceanogr.*, 62, 137–150, <https://doi.org/10.1002/lno.10381>, 2017.
- Silliman, B. R.: Drought, snails, and large-scale die-off of Southern U.S. salt marshes, *Science*, 310, 1803–1806, <https://doi.org/10.1126/science.1118229>, 2005.
- Silvestri, S., Defina, A., and Marani, M.: Tidal regime, salinity and salt marsh plant zonation, *Estuar. Coast. Shelf Sci.*, 62, 119–130, <https://doi.org/10.1016/j.ecss.2004.08.010>, 2005.
- Stevenson, J. C., Kearney, M. S., and Pendleton, E. C.: Sedimentation and erosion in a Chesapeake Bay brackish marsh system, *Mar. Geol.*, 67, 213–235, [https://doi.org/10.1016/0025-3227\(85\)90093-3](https://doi.org/10.1016/0025-3227(85)90093-3), 1985.
- Temmerman, S., Govers, G., Meire, P., and Wartel, S.: Modelling long-term tidal marsh growth under changing tidal conditions and suspended sediment concentrations, Scheldt estuary, Belgium, *Mar. Geol.*, 193, 151–169, [https://doi.org/10.1016/s0025-3227\(02\)00642-4](https://doi.org/10.1016/s0025-3227(02)00642-4), 2003a.
- Temmerman, S., Govers, G., Meire, P., and Wartel, S.: Simulating the long-term development of levee–basin topography on tidal marshes, *Geomorphology*, 63, 39–55, <https://doi.org/10.1016/j.geomorph.2004.03.004>, 2004.
- Temmerman, S., Govers, G., Wartel, S., and Meire, P.: Spatial and temporal factors controlling short-term sedimentation in a salt and freshwater tidal marsh, Scheldt estuary, Belgium, SW Netherlands, *Earth Surf. Process. Landf.*, 28, 739–755, <https://doi.org/10.1002/esp.495>, 2003b.
- Temmerman, S. and Kirwan, M. L.: Building land with a rising sea, *Science*, 349, 588–589, <https://doi.org/10.1126/science.aac8312>, 2015.
- Temmerman, S., Meire, P., Bouma, T. J., Herman, P. M. J., Ysebaert, T., and De Vriend, H. J.: Ecosystem-based coastal defence in the face of global change, *Nature*, 504, 79–83, <https://doi.org/10.1038/nature12859>, 2013.
- Temmerman, S., Moonen, P., Schoelynck, J., Govers, G., and Bouma, T. J.: Impact of vegetation die-off on spatial flow patterns over a tidal marsh, *Geophys. Res. Lett.*, 39, L03406, <https://doi.org/10.1029/2011gl050502>, 2012.
- Tommasini, L., Carniello, L., Ghinassi, M., Roner, M., and D'Alpaos, A.: Changes in the wind-wave field and related salt-marsh lateral erosion: inferences from the evolution of the Venice Lagoon in the last four centuries, *Earth Surf. Process. Landf.*, 44, 1633–1646, <https://doi.org/10.1002/esp.4599>, 2019.

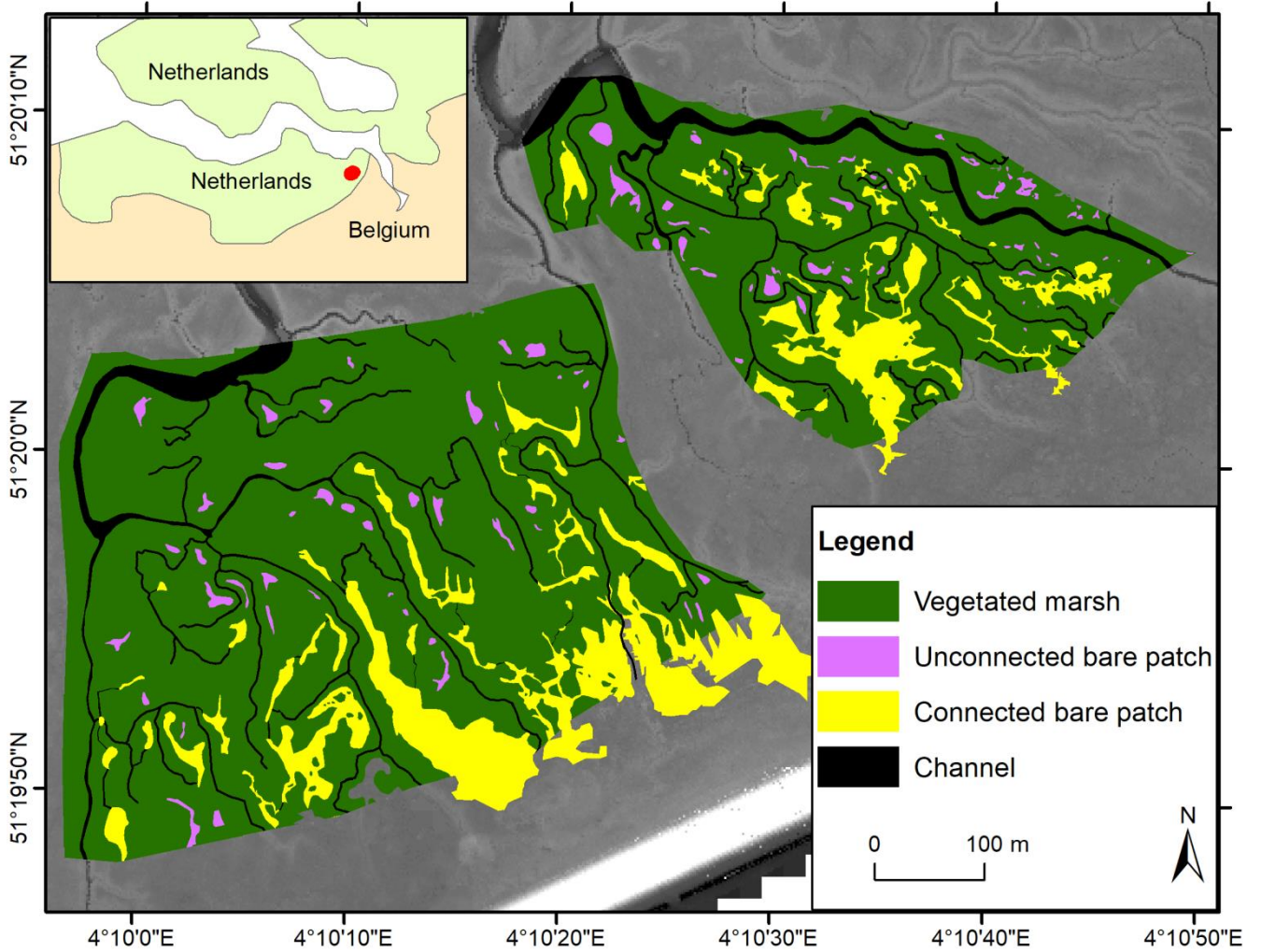
- Turner, R. E. and Rao, Y. S.: Relationships between wetland fragmentation and recent hydrologic changes in a deltaic coast, estuaries, 13, 272–281, <https://doi.org/10.2307/1351918>, 1990.
- 865 Ursino, N., Silvestri, S., and Marani, M.: Subsurface flow and vegetation patterns in tidal environments, *Water Resour. Res.*, 40, 191–201, <https://doi.org/10.1029/2003wr002702>, 2004.
- van Belzen, J., van de Koppel, J., Kirwan, M. L., van der Wal, D., Herman, P. M. J., Dakos, V., Kéfi, S., Scheffer, M., Guntenspergen, G. R., and Bouma, T. J.: Vegetation recovery in tidal marshes reveals critical slowing down under increased inundation, *Nat. Commun.*, 8, 15811, <https://doi.org/10.1038/ncomms15811>, 2017.
- 870 van Damme, S., Struyf, E., Maris, T., Ysebaert, T., Dehairs, F., Tackx, M., Heip, C., and Meire, P.: Spatial and temporal patterns of water quality along the estuarine salinity gradient of the Scheldt estuary (Belgium and The Netherlands): results of an integrated monitoring approach, *Hydrobiologia*, 540, 29–45, <https://doi.org/10.1007/s10750-004-7102-2>, 2005.
- van der Pluijm, A. M. and de Jong, D. J.: Historisch Overzicht Schorareaal in Zuid-West Nederland, Report Werkdocument
- 875 RIKZ/OS-98.860 x, Rijkswaterstaat - Rijksinstituut voor Kust en Zee, Utrecht, NL, 1998.
- van Wesenbeeck, B. K., van de Koppel, J., Herman, P. M. J., Bertness, M. D., van der Wal, D., Bakker, J. P., and Bouma, T. J.: Potential for sudden shifts in transient systems: distinguishing between local and landscape-scale processes, *Ecosystems*, 11, 1133–1141, <https://doi.org/10.1007/s10021-008-9184-6>, 2008.
- Vandenbruwaene, W., Bouma, T. J., Meire, P., and Temmerman, S.: Bio-geomorphic effects on tidal channel evolution: impact of vegetation establishment and tidal prism change, *Earth Surf. Process. Landf.*, 38, 122–132, <https://doi.org/10.1002/esp.3265>, 2013.
- 880
- Vandenbruwaene, W., Temmerman, S., Bouma, T. J., Klaassen, P. C., de Vries, M. B., Callaghan, D. P., van Steeg, P., Dekker, F., van Duren, L. A., Martini, E., Balke, T., Biermans, G., Schoelynck, J., and Meire, P.: Flow interaction with dynamic vegetation patches: implications for biogeomorphic evolution of a tidal landscape, *J. Geophys. Res.: Earth Surf.*, 116, 155–170, <https://doi.org/10.1029/2010jf001788>, 2011.
- 885
- Venier, C., D'Alpaos, A., and Marani, M.: Evaluation of sediment properties using wind and turbidity observations in the shallow tidal areas of the Venice Lagoon, *J. Geophys. Res.: Earth Surf.*, 119, 1604–1616, <https://doi.org/10.1002/2013jf003019>, 2014.
- Wamsley, T. V., Cialone, M. A., Smith, J. M., Atkinson, J. H., and Rosati, J. D.: The potential of wetlands in reducing storm surge, *Ocean Eng.*, 37, 59–68, <https://doi.org/10.1016/j.oceaneng.2009.07.018>, 2010.
- 890
- Wang, C., Menenti, M., Stoll, M. P., Feola, A., Belluco, E., and Marani, M.: Separation of ground and low vegetation signatures in LiDAR measurements of salt-marsh environments, *IEEE Trans. Geosci. Remote Sens.*, 47, 2014–2023, <https://doi.org/10.1109/tgrs.2008.2010490>, 2009.
- Wang, C. and Temmerman, S.: Does biogeomorphic feedback lead to abrupt shifts between alternative landscape states? An empirical study on intertidal flats and marshes, *J. Geophys. Res.: Earth Surf.*, 118, 229–240, <https://doi.org/10.1029/2012jf002474>, 2013.
- 895

- Wilson, C. A., Hughes, Z. J., FitzGerald, D. M., Hopkinson, C. S., Valentine, V., and Kolker, A. S.: Saltmarsh pool and tidal creek morphodynamics: dynamic equilibrium of northern latitude saltmarshes?, *Geomorphology*, 213, 99–115, <https://doi.org/10.1016/j.geomorph.2014.01.002>, 2014.
- 900 Wilson, K. R., Kelley, J. T., Croitoru, A., Dionne, M., Belknap, D. F., and Steneck, R.: Stratigraphic and ecophysical characterizations of salt pools: dynamic landforms of the webhannet salt marsh, Wells, ME, USA, *Estuaries Coasts*, 32, 855–870, <https://doi.org/10.1007/s12237-009-9203-7>, 2009.
- Wilson, K. R., Kelley, J. T., Tanner, B. R., and Belknap, D. F.: Probing the origins and stratigraphic signature of salt pools from north-temperate marshes in Maine, U.S.A, *J. Coast. Res.*, 26, 1007–1026, [https://doi.org/10.2112/jcoastres-d-](https://doi.org/10.2112/jcoastres-d-10-00007.1)
905 10-00007.1, 2010.
- Yang, S. L., Shi, B. W., Bouma, T. J., Ysebaert, T., and Luo, X. X.: Wave attenuation at a salt marsh margin: a case study of an exposed coast on the yangtze estuary, *estuaries coasts*, 35, 169–182, <https://doi.org/10.1007/s12237-011-9424-4>, 2012.
- Zaggia, L. and Ferla, M.: Studies on water and suspended sediment transport at the Venice Lagoon inlets, in: *Proceedings of the Fifth International Symposium WAVES*, Madrid, Spain, 3-7 July 2005, 1–10, 2005.
- 910 Zirino, A., Elwany, H., Neira, C., Maicu, F., Mendoza, G., and Levin, L. A.: Salinity and its variability in the Lagoon of Venice, 2000–2009, *Adv. Oceanogr. Limnol.*, 5, 41–59, <https://doi.org/10.4081/aiol.2014.5350>, 2014.

915 **Table 1. Overview of bare patch number (bare patches smaller than 1 m² were excluded from the analysis), Lidar pixels and GPS measurements of the three field sites.**

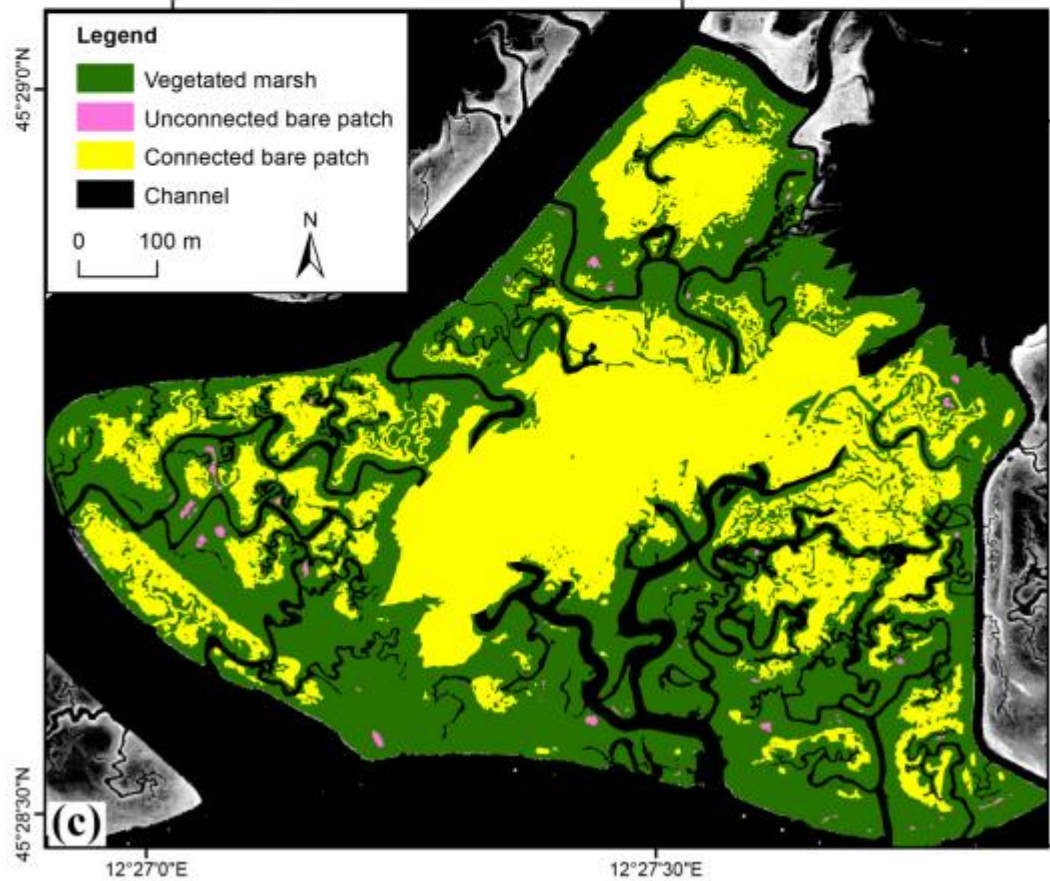
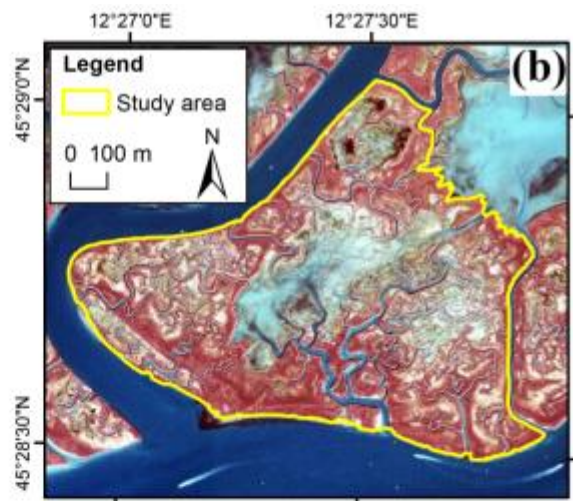
Field site	Type	Number of bare patches	Number of Lidar pixels	Number of GPS measurements
Saefthinghe (NL)	Marsh	-	67729	-
	Unconnected bare patch	97	1722	-
	Connected bare patch	58	12651	-
San Felice (IT)	Marsh	-	361261	340
	Unconnected bare patch	70	2556	-
	Connected bare patch	124	260140	95
Blackwater (USA)	Marsh	-	184871	93
	Unconnected bare patch	255	-	36
	Connected bare patch	227	-	31

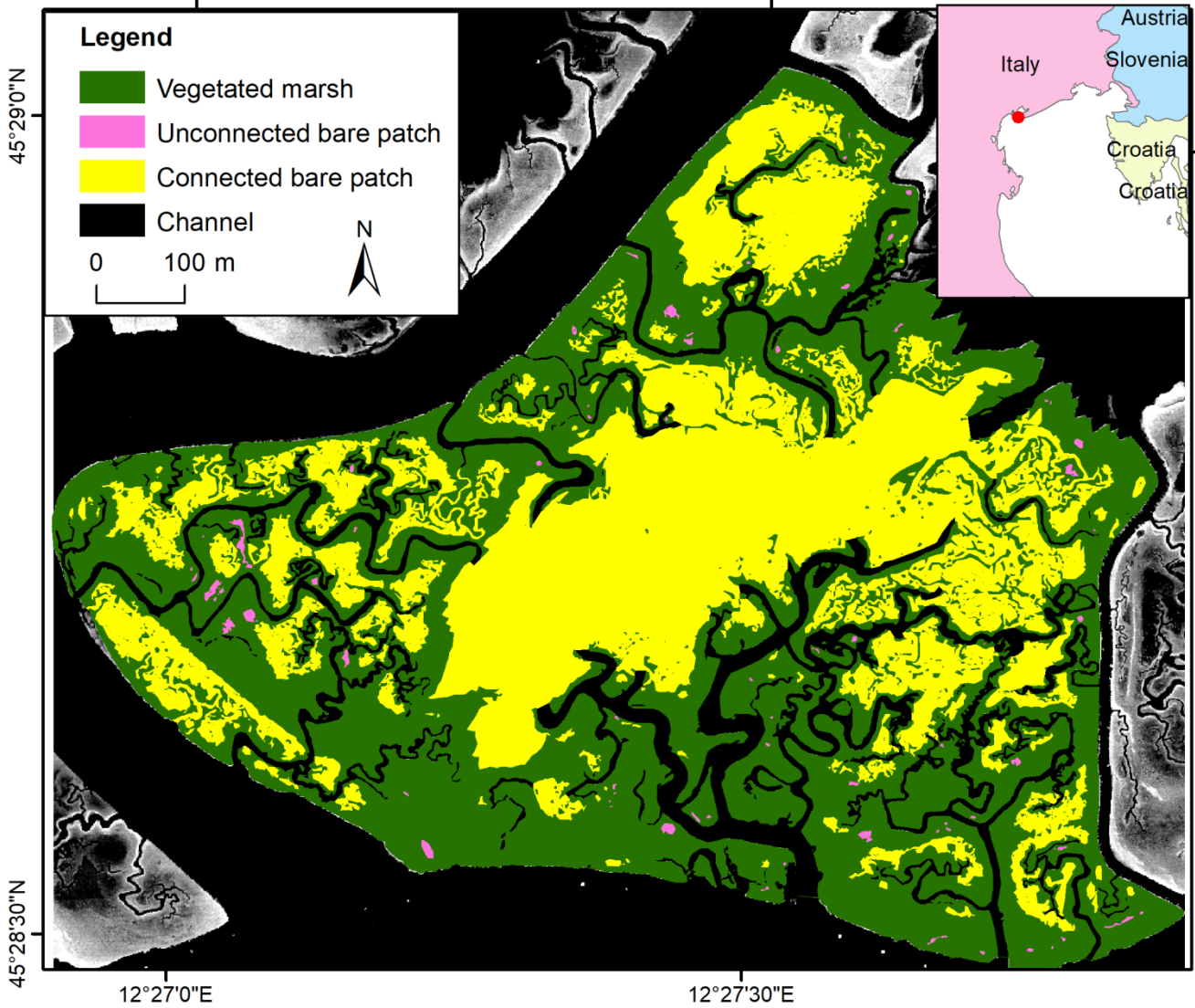




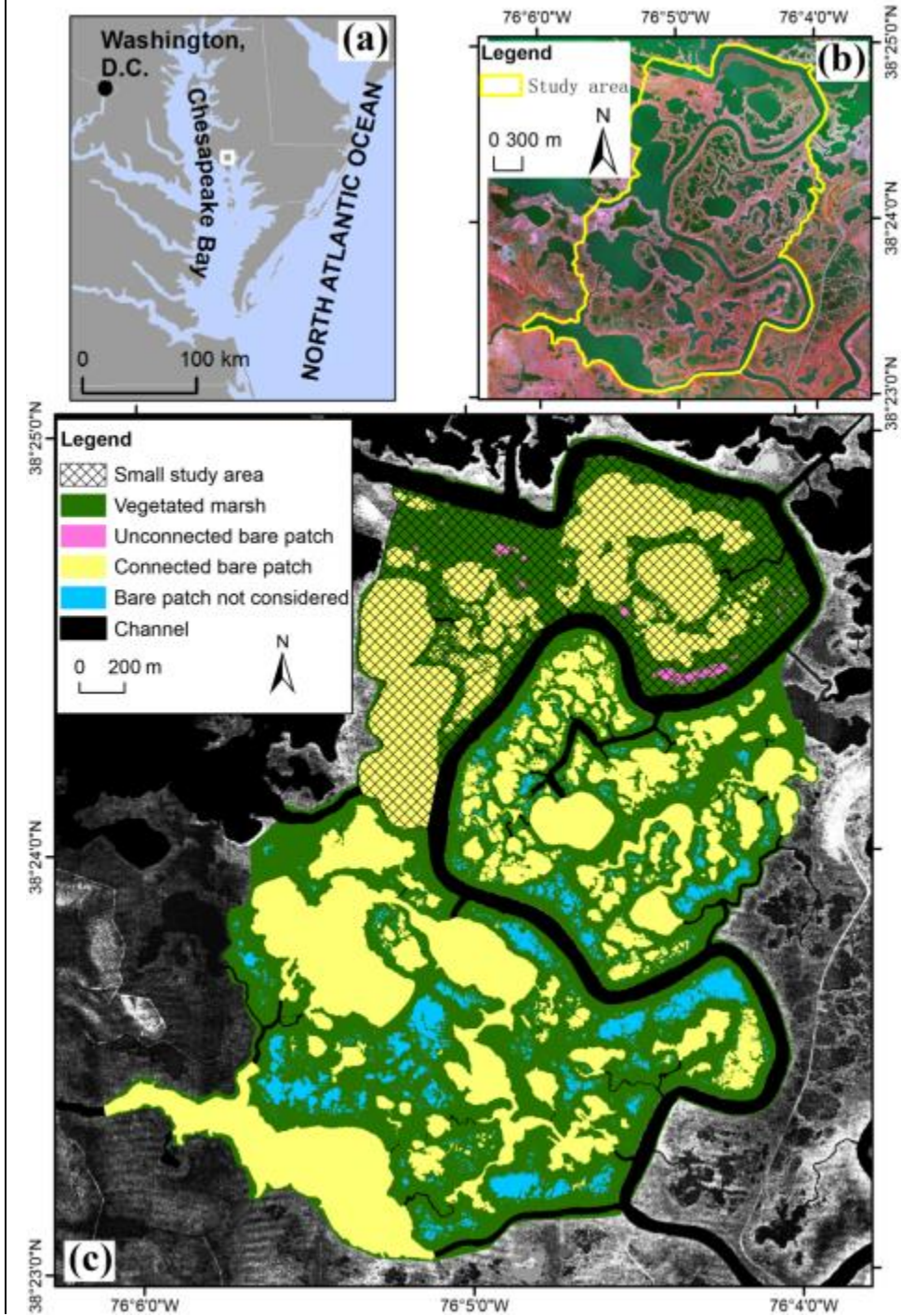
920

Figure 1: Study area in the Saeftinghe marsh. (a) Location within Northwest Europe. (b) False-color aerial photograph of 2004 for the study area. (c) Spatial distribution of vegetated marshes, unconnected bare patches and connected bare patches in 2004 with LIDAR images as background.





925 **Figure 2: Study area in the San Felice marsh. (a) Location within South Europe. (b) IKONOS image of 2006 for the study area. (c)** Spatial distribution of vegetated marshes, unconnected bare patches and connected bare patches in 2002 with LIDAR images as background.



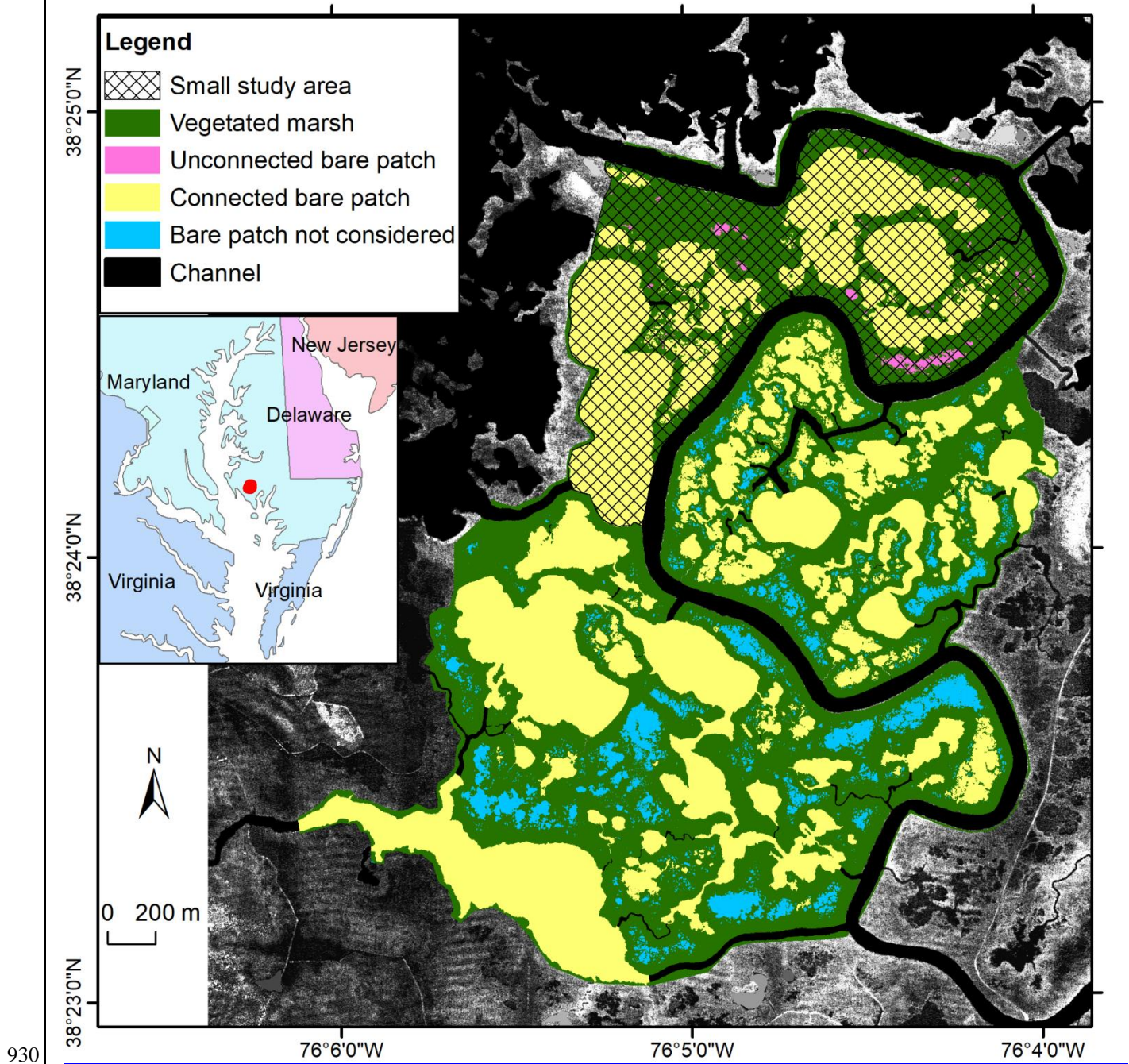
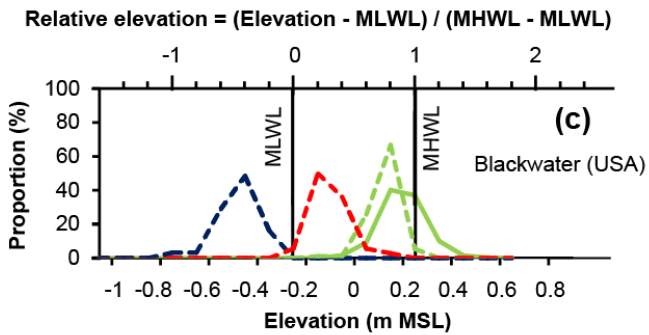
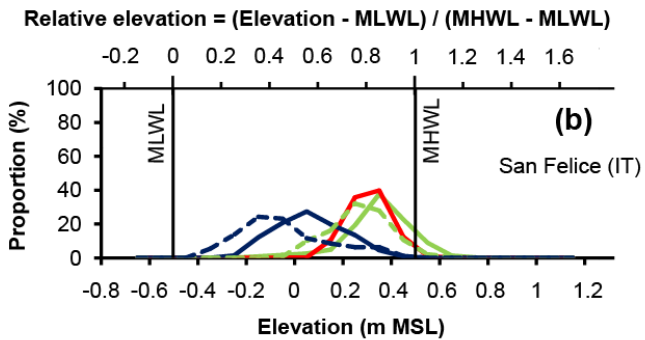
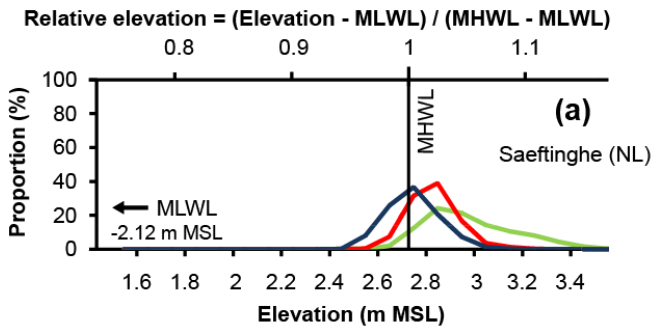
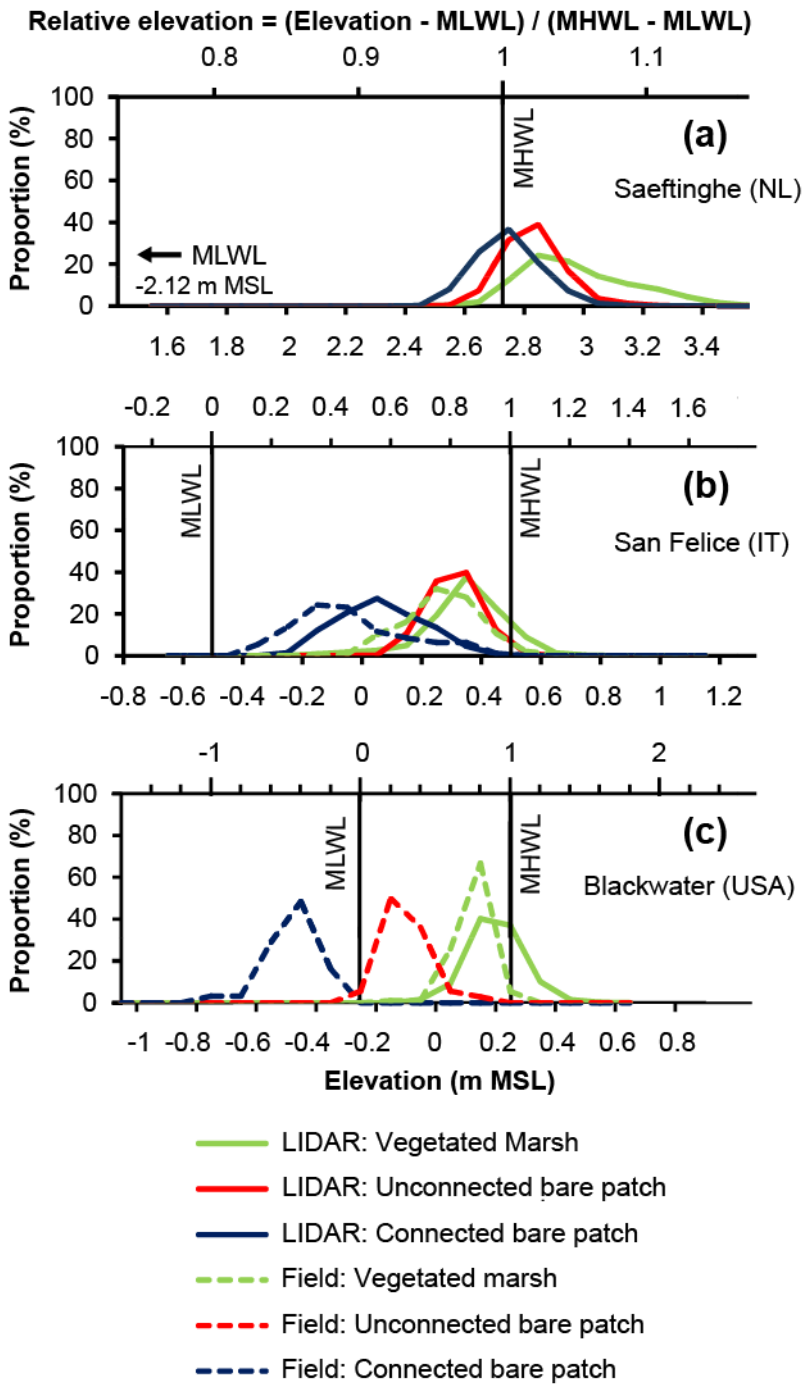


Figure 3: Study area in the Blackwater Marshes. (a) Location within the USA. (b) False color aerial photograph of 2010 for the study area. (c) Spatial distribution of vegetated marshes, unconnected bare patches and connected bare patches in 2010 with LIDAR images as background.



- LIDAR: Vegetated Marsh
- LIDAR: Unconnected bare patch
- LIDAR: Connected bare patch
- - - Field: Vegetated marsh
- - - Field: Unconnected bare patch
- - - Field: Connected bare patch



935

Figure 4: Elevation distribution of vegetated marshes, unconnected bare patches and connected bare patches based on LIDAR surveys and field surveys for (a) Saeftinghe, (b) San Felice and (c) Blackwater. In each figure, the bottom X-axis shows the absolute elevation (in m relative to mean sea level) and the top X-axis shows the relative elevation (dimensionless, as defined in Eq. (1)). The elevation on the bottom main X-axis is relative to the local mean sea level (i.e., m MSL). The relative elevation on the top secondary X-axis is rescaled as a proportion of the local tidal range (see Eq. (1)). The proportion on the Y-axis is calculated based on LIDAR or field

940

measurements as the number of pixels or samples in each elevation class (every 0.1 m) relative to the total number of pixels or samples for each feature. The exact numbers in each category are given in Table 1. The MLWL in Saeftinghe is 2.12 m lower than MSL, which is outside of the range of the main X-axis in Figure 4a. MLWL and MHWL definitions at Blackwater are approximate since water level changes are dominated by meteorological rather than astronomical influences.

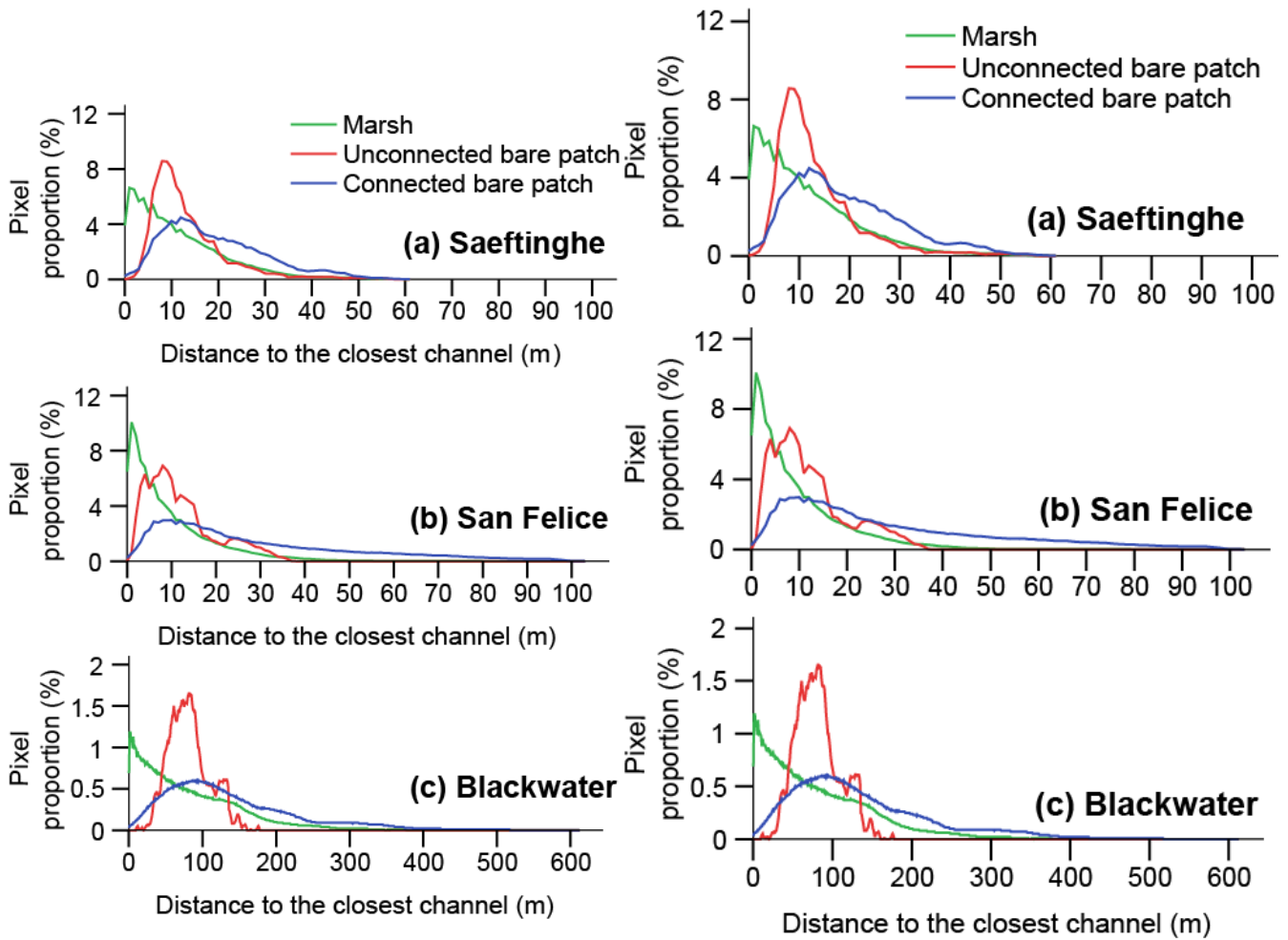
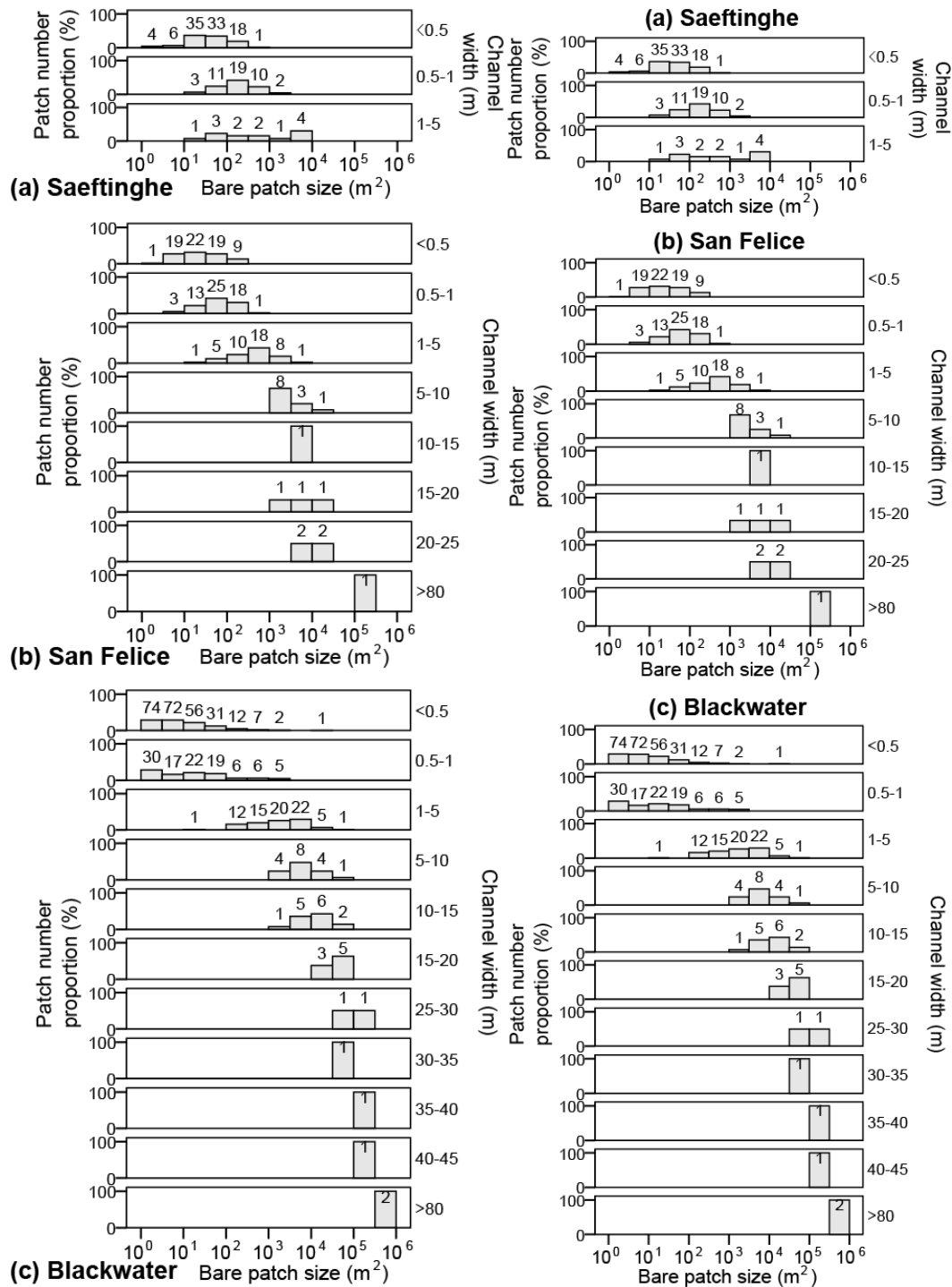


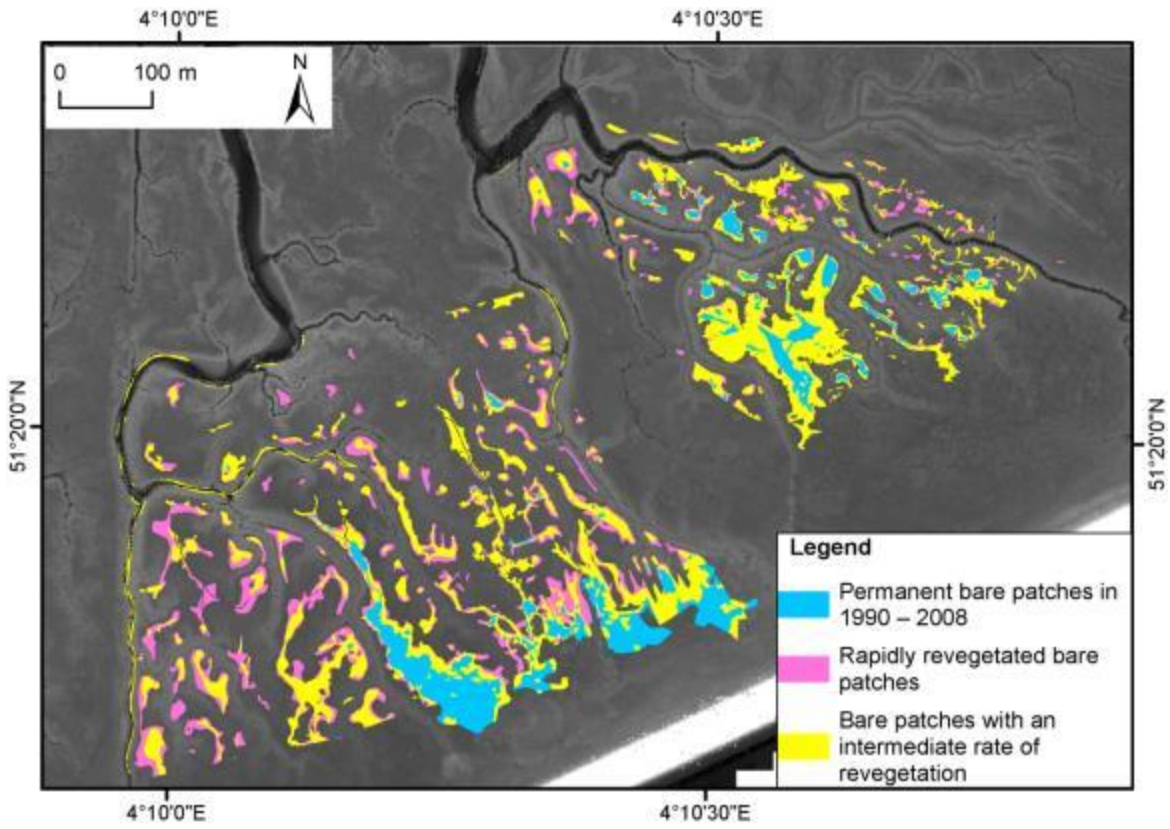
Figure 5: Frequency distribution of distances to the closest channel in (a) Saeftinghe, (b) San Felice, and (c) Blackwater. The proportion is calculated as the number of pixels in each distance class (every 1 m) relative to the total number of pixels for each feature, i.e., vegetated marshes, unconnected bare patches or connected bare patches.

945

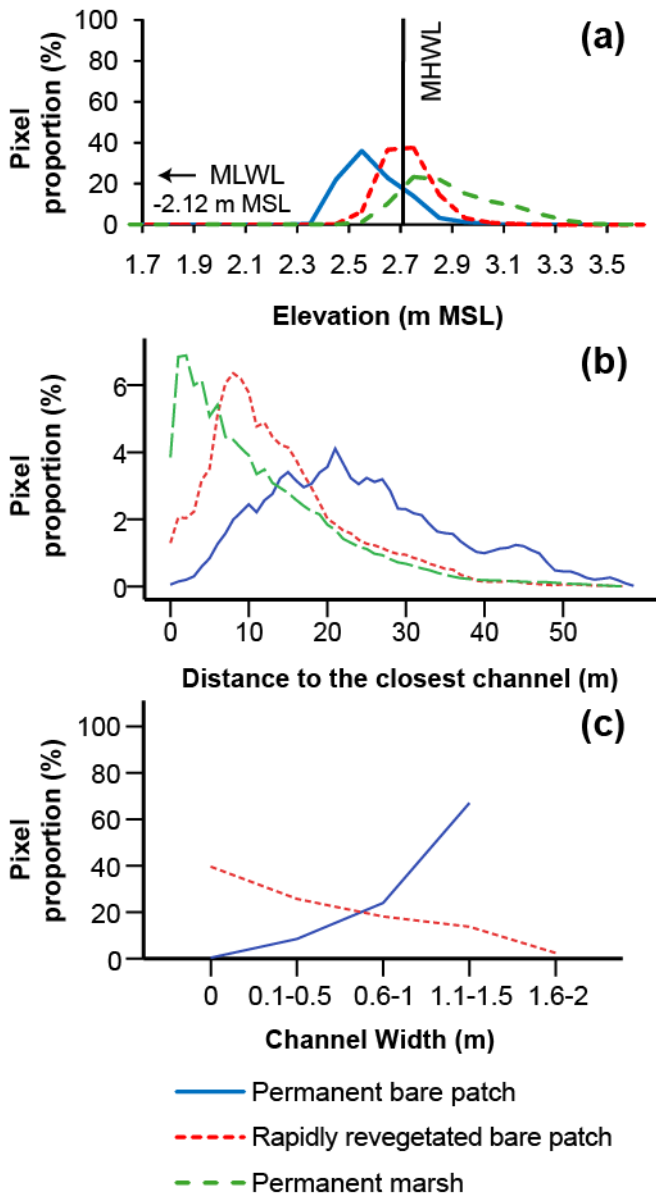


950

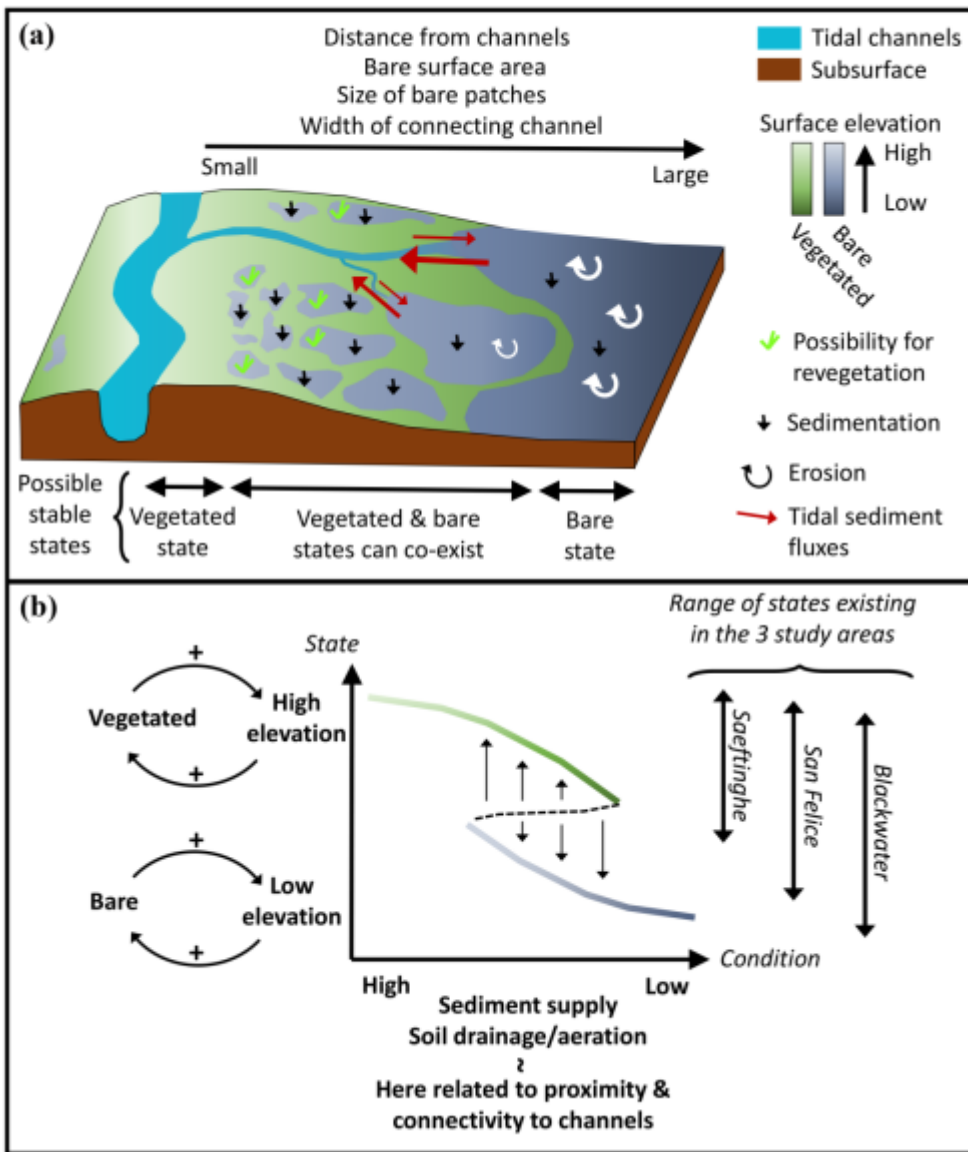
Figure 6: Frequency distribution of bare patch sizes in relation to the connected channel width in (a) Saeftinghe, (b) San Felice, and (c) Blackwater. Note the X-axis is in logarithmic scale. The patch number proportion is calculated as the number of bare patches in each class of bare patch size relative to the total number of bare patches for each category of channel width. The number of bare patches in each size class is labeled at the top of the bars.

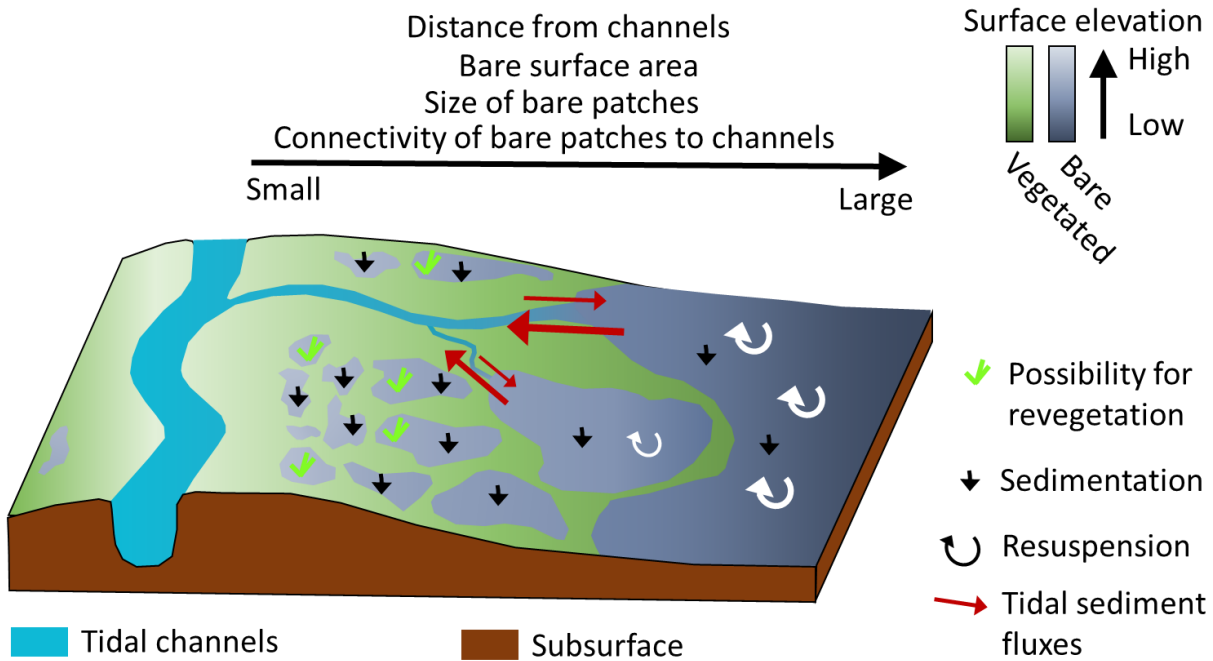


955 [Figure 7: Spatial distribution of bare patches with a different rate of revegetation in the period of 1990–2008 in Saeftinghe.](#)



960 **Figure 87:** Frequency distribution of (a) elevation, (b) distances to the closest channel, and (c) the connected channel width for permanent bare patches, rapidly revegetated bare patches and permanent marsh areas [in Saeftinghe](#). The elevation is relative to mean sea level and binned into 0.1 m intervals. The proportion in panel (a) is calculated as the number of pixels in each elevation class (every 0.1 m) relative to the total number of pixels for each feature. The proportion in panel (b) is calculated as the number of pixels in each distance class (every 1.0 m) relative to the total number of pixels for each feature. The proportion in panel (c) is calculated as the number of pixels in each class of channel width relative to the total number of pixels for each feature.





965

Figure 98: Conceptual model summarizing results and interpretations. In (a) the size of the arrows is indicative for hypothetical magnitudes of sediment accretion (black arrows), erosion (white arrows) and tidal sediment fluxes (red arrows). The hypotheses are that the larger bare patches connected to the channels experience stronger incoming and outgoing tidal currents, and more waves during flooding (longer wind fetch length), favoring erosion and tidal export of sediments via the channels; while the smaller unconnected bare patches experience weak tidal currents (because of obstruction by surrounding vegetation), few or no waves (small wind fetch length), limiting erosion and allowing accretion. In (b) we interpret the results within the framework of the alternative stable state theory, using a hypothetical/conceptual plot of states (Y axis) versus conditions (X axis). State variables are vegetation biomass (high for the vegetated state = the green curve, versus zero for the bare state = the grey curve) and surface elevation (different shades of green and grey). Conditional variables are sediment supply and soil drainage/aeration, which are in our framework spatially varying with distance and connectivity to tidal channels.

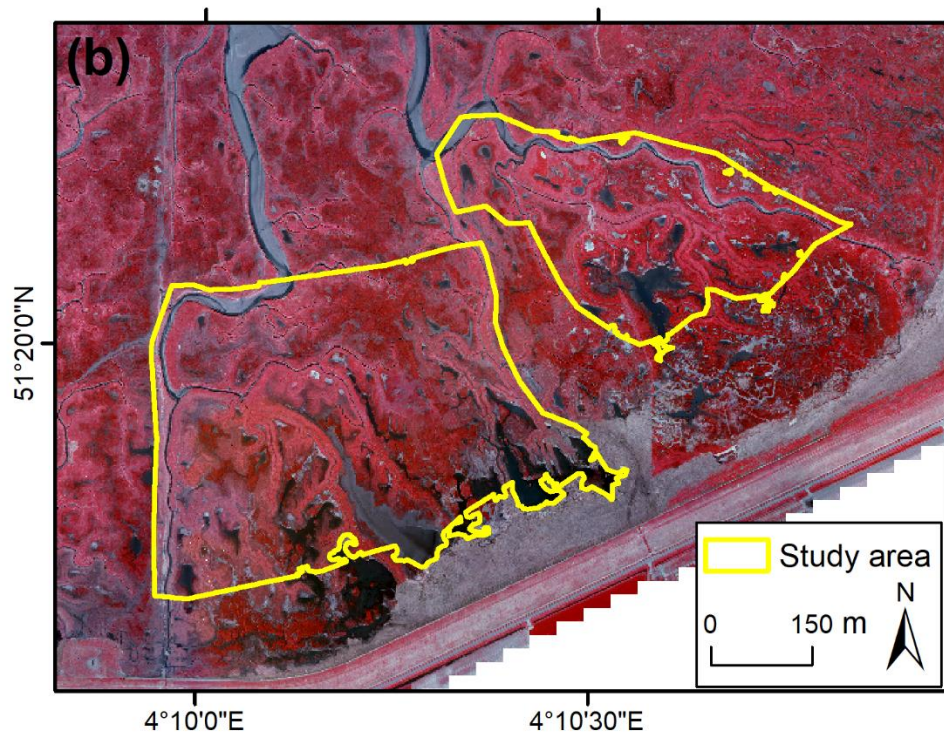
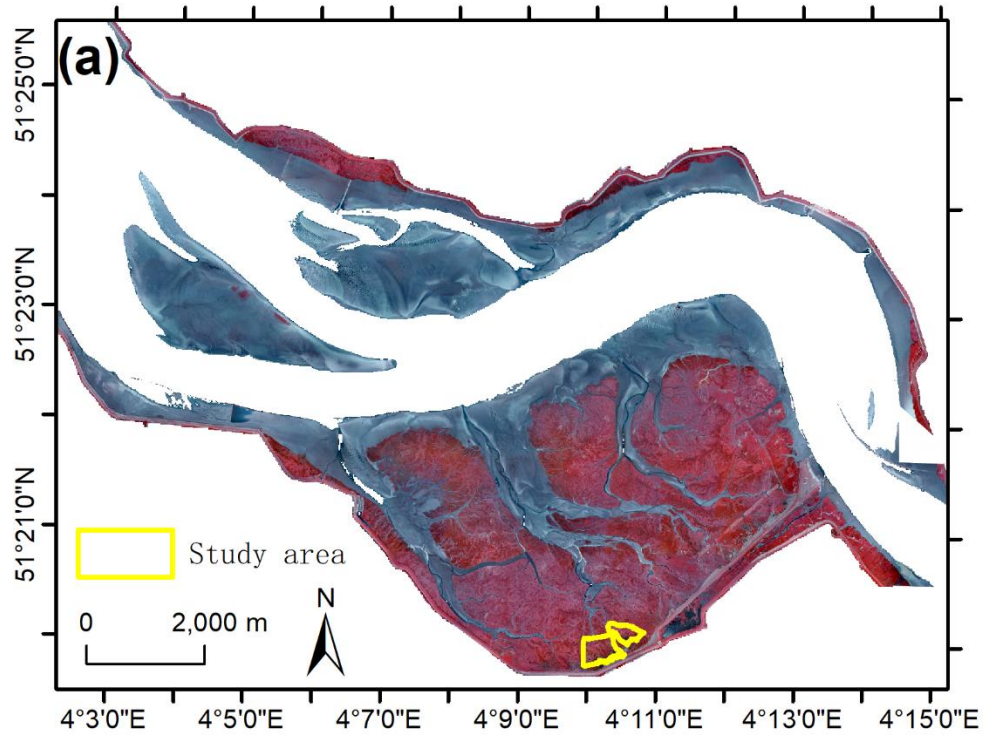
970

975

Supplement of

~~Different Coastal Marsh Sites Reflect Similar Topographic Conditions for Bare Patches and Vegetation Recovery~~
Different coastal marsh sites reflect similar topographic conditions under which bare patches and vegetation recovery occurs

Chen Wang^{1,2}, Lennert Schepers², Matthew L. Kirwan³, Enrica Belluco⁴, Andrea D'Alpaos⁵, Qiao Wang^{6,1}, Shoujing Yin¹, and Stijn Temmerman²



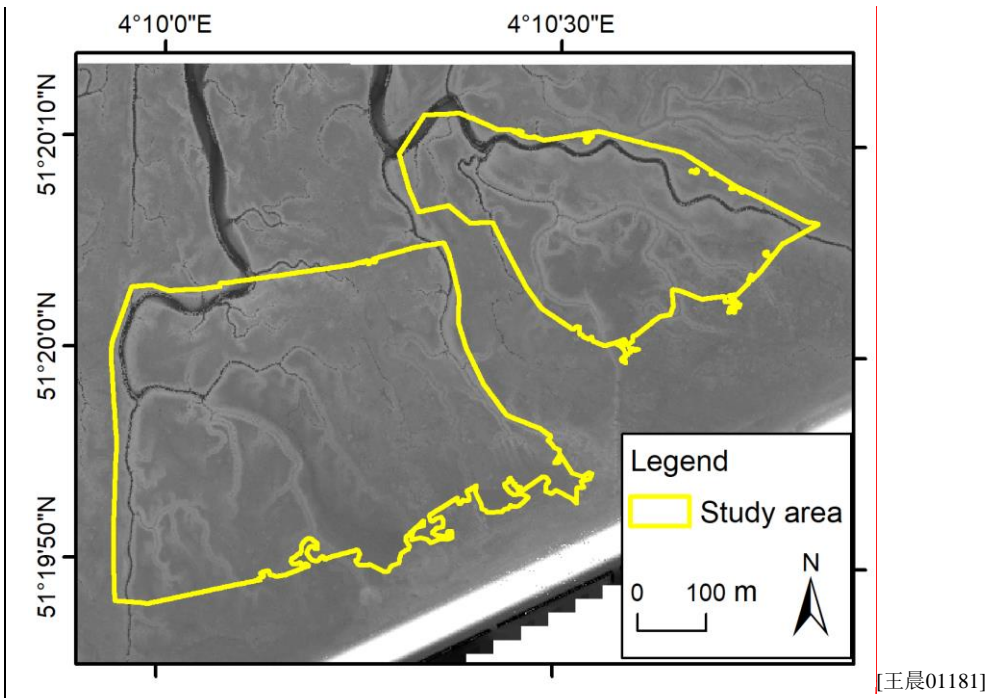
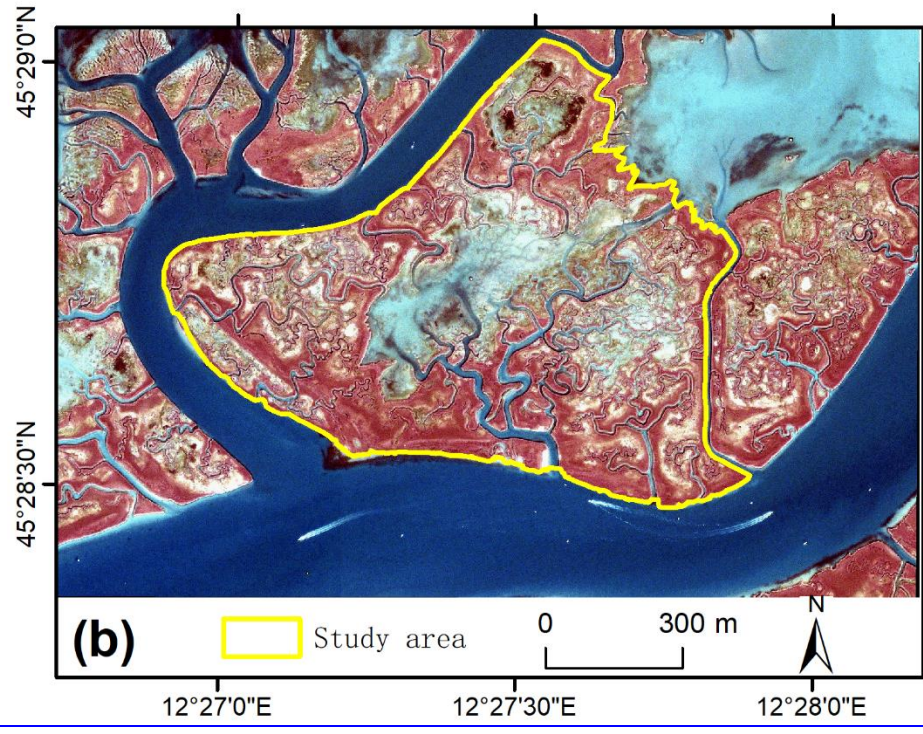
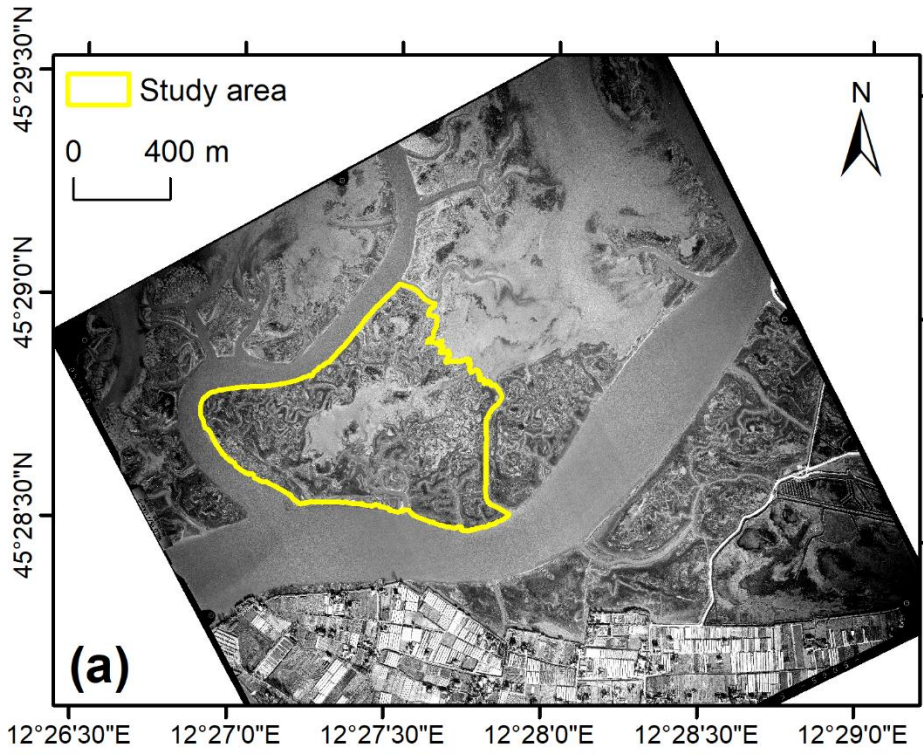


Figure S1. [False-color aerial photograph of 2004 LIDAR data](#) for [the study area in the Saeftinghe marsh](#).



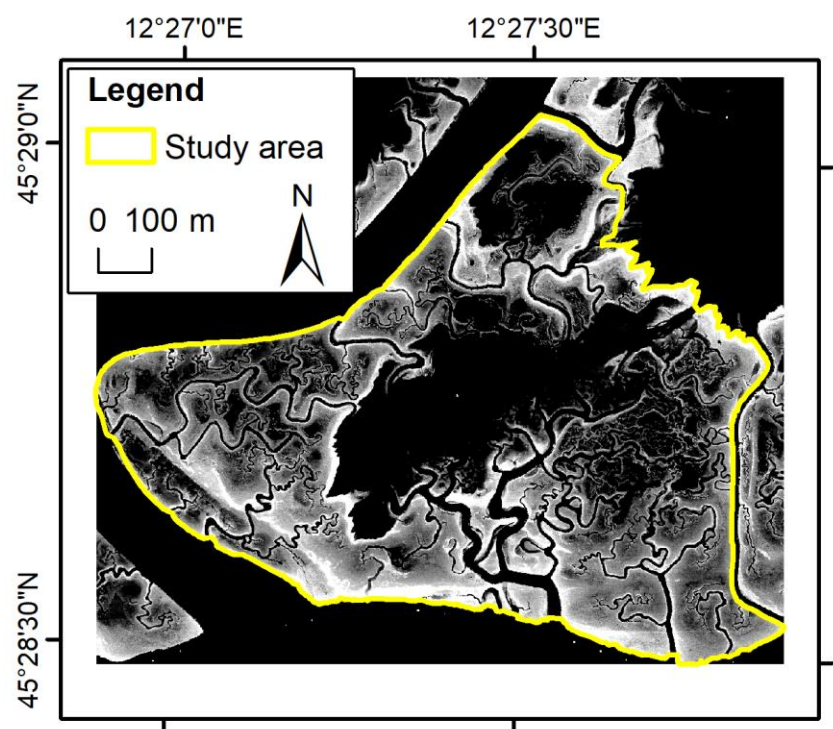
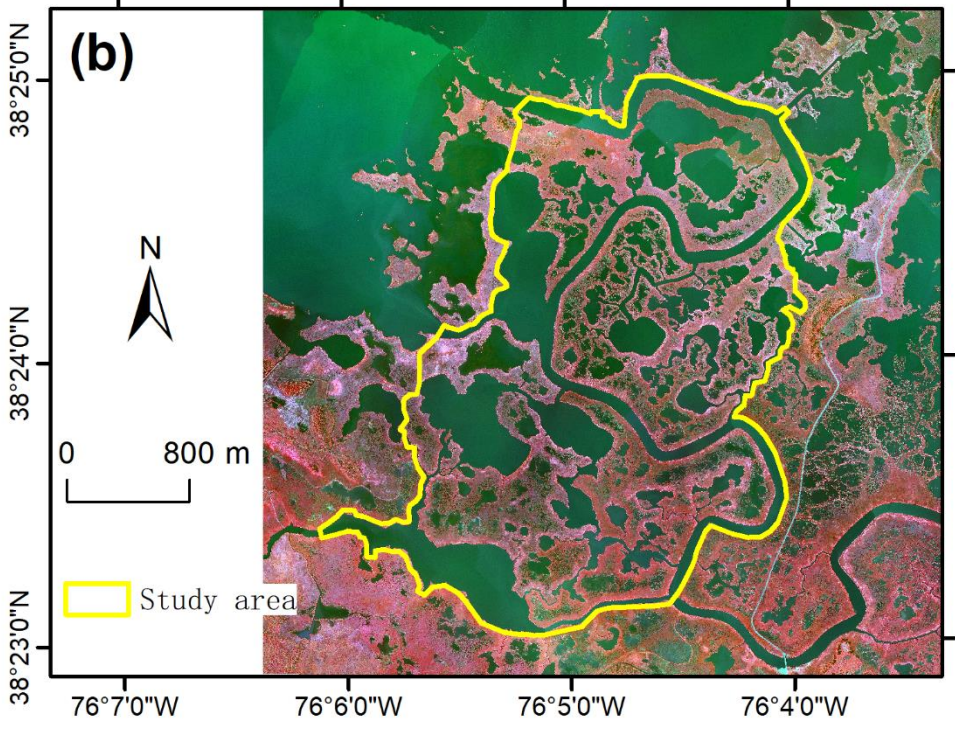
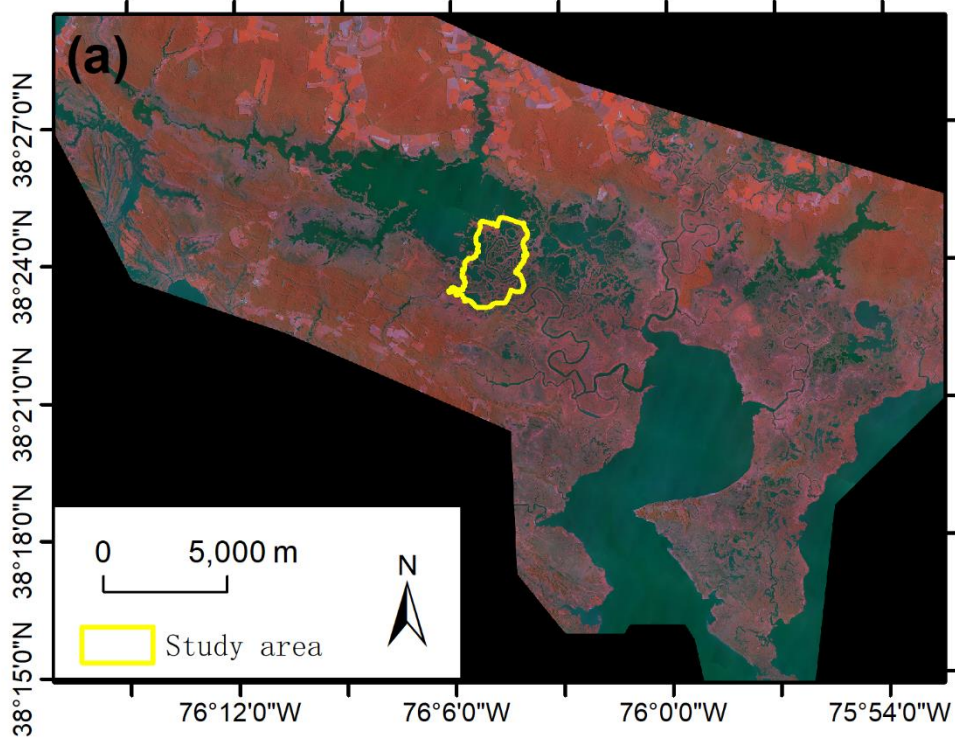


Figure S2. [Aerial photograph acquired in 2000 showing the surrounding landscape \(a\)](#) and [IKONOS image of 2006 \(b\)](#) [LIDAR data](#) for [the study area in the San Felice marsh, San Felice.](#)



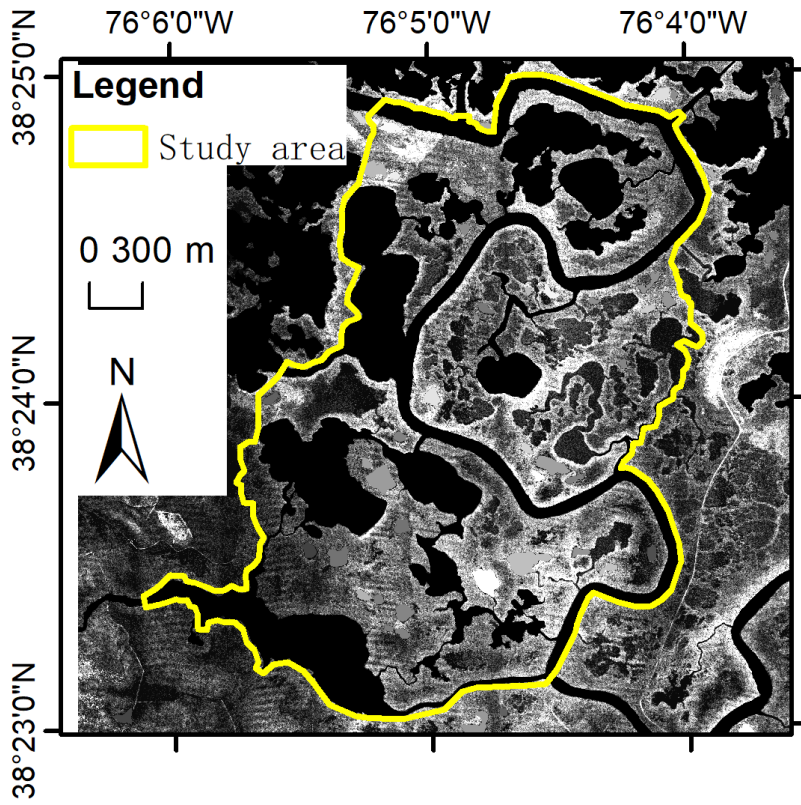


Figure S3. [False-color Aerial photograph of 2010 LIDAR data](#) for [the study area in the Blackwater Marshes, Blackwater.](#)

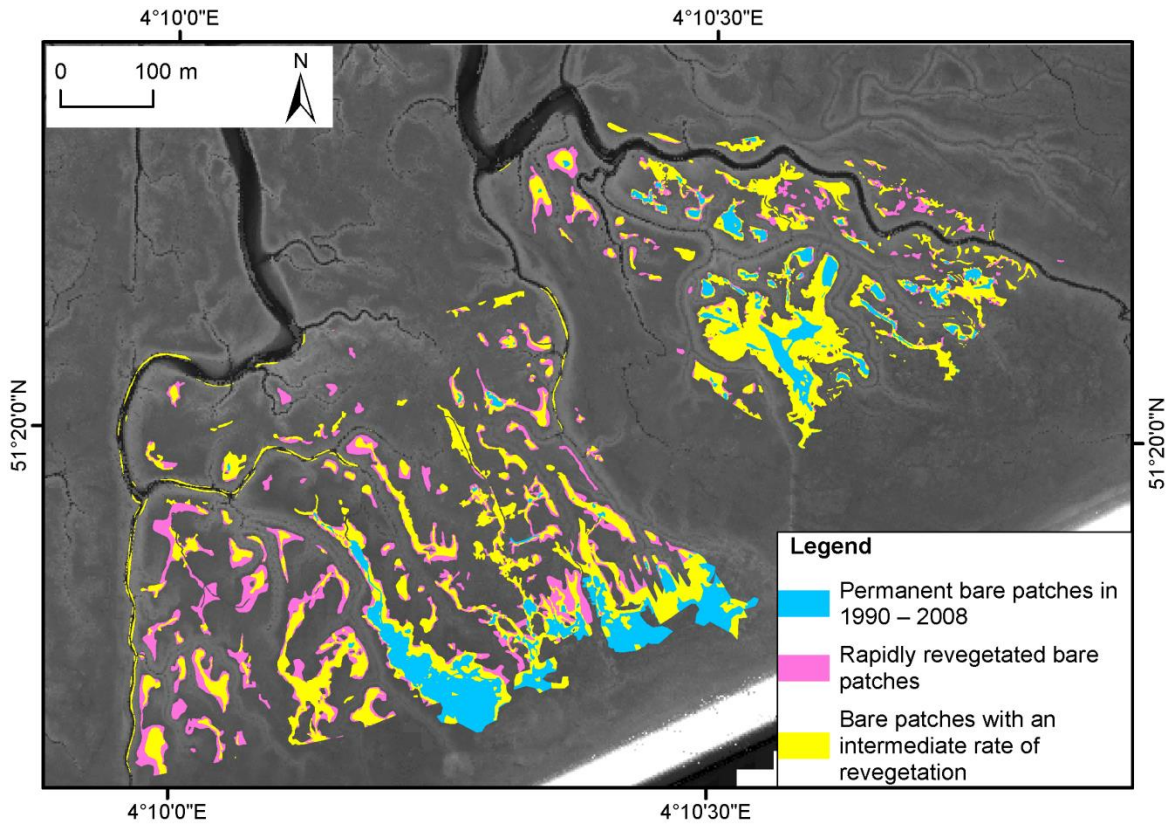


Figure S4: Spatial distribution of bare patches with a different rate of revegetation in the period of 1990–2008 in Saeftinghe.