

1 **Time-Of-Flight monitoring reveals higher sediment redistribution rates related to burrowing animals**
2 **than previously assumed**

3

4 *Paulina Grigusova*¹, *Annegret Larsen*², *Sebastian Achilles*¹, *Roland Brandl*³, *Camilo del Río*^{4,5}, *Nina Farwig*⁶,
5 *Diana Kraus*⁶, *Leandro Paulino*⁷, *Patricio Plissock*^{4,8,9}, *Kirstin Übernickel*¹⁰, *Jörg Bendix*¹

6

7

8 ¹ Laboratory for Climatology and Remote Sensing, Department of Geography, University of Marburg, 35037
9 Marburg, Germany; paulina.grigusova@staff.uni-marburg.de (P.G.); bendix@geo.uni-marburg.de (J.B.)

10 ² Soil Geography and Landscape, Department of Environmental Sciences,
11 Wageningen University & Research, 6700 AA Wageningen, The Netherlands; annegret.larsen@wur.nl

12 ³ Animal Ecology, Department of Biology, University of Marburg, 35032 Marburg, Germany;
13 brandlr@biologie.uni-marburg.de

14 ⁴ Facultad de Historia, Geografía y Ciencia Política, Instituto de Geografía, Pontificia Universidad Católica de
15 Chile, 782-0436 Santiago, Chile; plissock@uc.cl; cdelriol@uc.cl

16 ⁵ Centro UC Desierto de Atacama, Pontificia Universidad Católica de Chile, 782-0436 Santiago, Chile;
17 cdelriol@uc.cl

18 ⁶ Conservation Ecology, Department of Biology, University of Marburg, 35047 Marburg, Germany;
19 diana.kraus@biologie.uni-marburg.de (D.K.); nina.farwig@biologie.uni-marburg.de (N.F.)

20 ⁷ Facultad de Agronomía, Universidad de Concepción, 3780000 Chillán, Chile; lpaulino@udec.cl

21 ⁸ Facultad de Ciencias Biológicas, Departamento de Ecología, Pontificia Universidad Católica de Chile, 8331150
22 Santiago, Chile; plissock@uc.cl

23 ⁹ Center of Applied Ecology and Sustainability (CAPES), Pontificia Universidad Católica de Chile, 8331150
24 Santiago, Chile; plissock@uc.cl

25 ¹⁰ Earth System Dynamics, Department of Geosciences, University of Tübingen, 72076 Tübingen, Germany;
26 kirstin.uebernickel@uni-tuebingen.de

27

28 *Corresponding author:*

29 Paulina Grigusova

30 paulina.grigusova@staff.uni-marburg.de

31

32

33

34

35

36

37

38

39

40

41 **Abstract**

42 Burrowing animals influence surface microtopography and hillslope sediment redistribution, but changes often
43 remain undetected due to a lack of automated high resolution field monitoring techniques. In this study, we
44 present a new approach to quantify microtopographic variations and surface changes caused by burrowing
45 animals and rainfall-driven erosional processes applied to remote field plots in arid and mediterranean Chile.
46 We compared the mass balance of redistributed sediment between burrow and burrow embedding area,
47 quantified the cumulative sediment redistribution caused by animals and rainfall, and upscaled the results to a
48 hillslope scale. The newly developed instrument, a Time-of-Flight camera, showed a very good detection
49 accuracy. The animal-caused cumulative sediment redistribution was $14.6 \text{ cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$ in the Mediterranean
50 and $16.4 \text{ cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$ in the arid climate zone. The rainfall-caused cumulative sediment redistribution within
51 areas affected by burrowing animals was higher ($-10.4 \text{ cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$) in the Mediterranean than the arid
52 climate zone ($-1.4 \text{ cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$). Daily sediment redistribution during rainfall within burrows were up to 350%
53 / 40% higher in the mediterranean / arid zone compared to burrow embedding area, and much higher than
54 previously reported in studies not based on continuous microtopographic monitoring. Furthermore, 38% of the
55 sediment eroding from the burrows accumulated within the burrow entrance while 62% was incorporated into
56 overall hillslope sediment flux. The animals burrowed between on average 1.2 – 2.3 times a month and the
57 burrowing intensity increased after rainfall. Our findings can be implemented into long-term soil erosion models
58 that rely on soil processes but do not yet include animal-induced surface processes on microtopographical
59 scales in their algorithms.

60

61 **Keywords:** Biogeomorphology, bioturbation, sediment transport, burrowing animals, rainfall, Time-of-Flight
62 camera, Chile

63

64

65

66

67

68

69

70

71

72

73

74

75

76

77

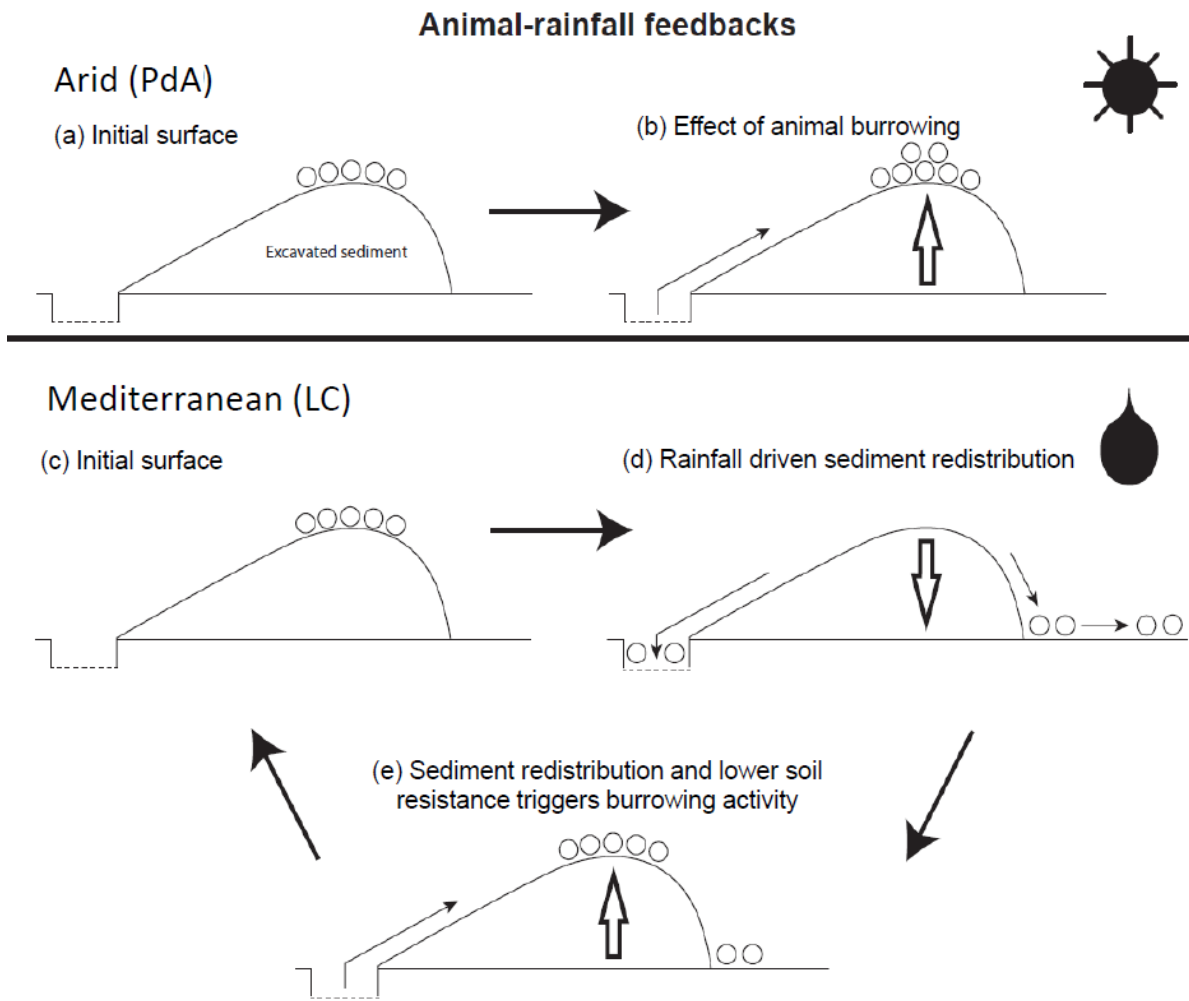
78

79

80

81

82 **Graphical abstract**



83

84

85 1. Introduction

86 Animal burrowing activity affects surface microtopography (Reichman und Seabloom 2002; Kinlaw
87 und Grasmueck 2012), surface roughness (Yair 1995; Jones et al. 2010; Hancock und Lowry 2021) and soil
88 physical properties (Ridd 1996; Yair 1995; Hall et al. 1999; Reichman und Seabloom 2002; Hancock und Lowry
89 2021; Coombes 2016; Larsen et al. 2021; Corenblit et al. 2021). Previous studies estimated both positive as
90 well as negative impacts of burrowing animals on sediment redistribution rates. The results were obtained by
91 applying tests under laboratory conditions using rainfall simulators, conducting several field campaigns weeks
92 to months apart, or by measuring the volume of excavated or eroded sediment in the field using methods such
93 as erosion pins, splash boards, or simple rulers (Imeson und Kwaad 1976; Reichman und Seabloom 2002;
94 Wei et al. 2007; Le Hir et al. 2007; Li et al. 2018; Li et al. 2019b; Li et al. 2019c; Voiculescu et al. 2019; Chen
95 et al. 2021; Übernickel et al. 2021b; Li et al. 2019a). Although burrowing animals are generally seen as
96 ecosystem engineers (Gabet et al. 2003; Wilkinson et al. 2009), their role in soil erosion, in general, and for
97 numerical soil erosion models, in particular, is, to date, limited to predictions of the burrow locations and particle
98 mixing at these locations (Black und Montgomery 1991; Meysman et al. 2003; Yoo et al. 2005; Schiffrers et al.
99 2011). The complex interaction of sediment excavation and accumulation, and erosion processes at the burrow
100 and hillslope scales are not yet included in the modelling, as for this, a suitable method capable of measuring
101 all occurred redistribution processes is needed.

102 The reason for this knowledge gap is that previous studies have not provided data on low magnitude
103 but frequently occurring sediment redistribution due to the specific limitations of their approaches. Field
104 experiments with, for example, rainfall simulators can unveil processes but cannot cover the time-dependant
105 natural dynamics of sediment redistribution. For data samplings that used methods such as erosion pins or
106 splash boards, the sites had to be revisited each time and the data were thus obtained only sporadically
107 (Imeson und Kwaad 1976; Hazelhoff et al. 1981; Richards und Humphreys 2010). Similarly, estimations of the
108 excavated sediment volume are currently limited to one-time measurements or studies conducted several
109 months apart (Black und Montgomery 1991; Hall et al. 1999; Yoo et al. 2005). We expect that non-continuously
110 conducted measurements do not include all frequently occurring excavation and erosion processes. For this,
111 a spatio-temporally high-resolution and continuous monitoring of sediment redistribution is needed.

112 High-resolution, ground-based imaging sensing techniques might overcome such aforementioned
113 problems. Terrestrial laser scanner systems have shown to be a suitable tool for estimation of sediment
114 redistribution and erosion processes (Nasermoaddeli und Pasche 2008; Afana et al. 2010; Eltner et al. 2016a;
115 Eltner et al. 2016b; Longoni et al. 2016). However, they are expensive and labour-intensive. A continuous,
116 automated monitoring of many mound areas in parallel is for this reason not possible. An already applied low-
117 cost (up to 5000 USD) topographic monitoring technique is time-lapse photogrammetry which can be applied
118 at variable observation distances and scales (e.g. (James und Robson 2014; Galland et al. 2016; Eltner et al.
119 2017; MALLALIEU et al. 2017; Kromer et al. 2019; Blanch et al. 2021). For this technique, the surface has to
120 be monitored under various angles for which several devices are needed to be installed in the field.

121 In contrast, The Time-of-Flight (ToF) technology exhibits lower spatial resolution and aerial coverage
122 compared to time-lapse photogrammetry. However, as an active remote sensing tool it can also be used at
123 night. Additionally, the processing is less complex compared to photogrammetry because the distance values
124 are immediately received in a local coordinate system. ToF offers here a new possibility for surface monitoring,
125 as a technique for a cost-effective high-resolution monitoring of sediment redistribution (Eitel et al. 2011;
126 Hänsel et al. 2016) which can be achieved by a simple installation of one device in the field. ToF-based

127 cameras illuminate the targeted object with a light source for a known amount of time and then estimate the
128 distance between the camera and the object by measuring the time needed for the reflected light to reach the
129 camera sensor (Sarbolandi et al. 2018).

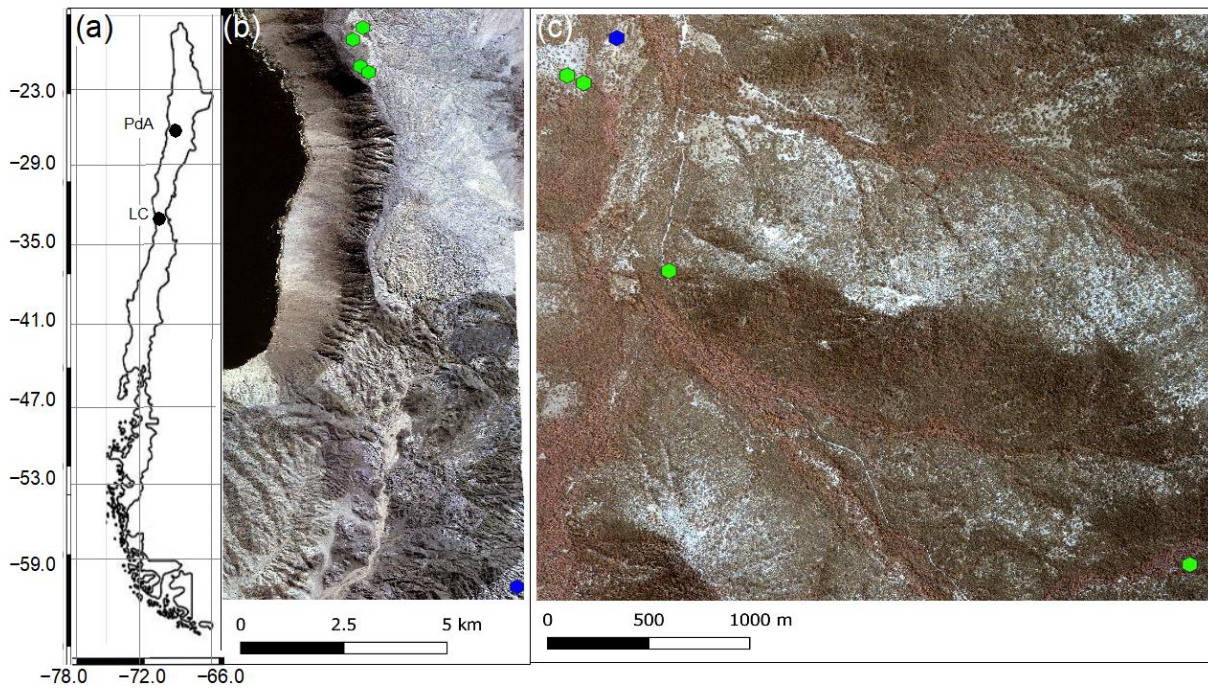
130 In our study we developed, tested and applied a cost-effective Time-of-Flight camera for automated
131 monitoring of the rainfall and animal-driven sediment redistribution burrows and burrow embedding areas by
132 burrowing animals with high temporal (four times a day) and spatial (6 mm) resolution. For this, we equipped
133 several plots in remote study sites in the Chilean arid and mediterranean climate zone. We selected these
134 sites in order to analyse sediment redistribution by burrowing activity of vertebrates under different rainfall
135 regimes and as these sites have been shown to be particularly strongly affected by burrowing activity
136 (Grigusova et al. 2021). We estimated the burrowing intensity and its dependence on rainfall. Then, we
137 quantified the daily sediment redistribution within burrow and burrow embedding area. We analysed the
138 impacts of animal burrowing activity and rainfall on the sediment redistribution and quantified the volume of
139 sediment which is additionally incorporated to the hillslope sediment flux due to the presence of burrows.
140 Finally, we estimated sediment redistribution on a burrow scale and upscaled sediment redistribution rates to
141 the entire hillslopes.

142

143 **2. Study area**

144 Our study sites were located in the Chilean Coastal Cordillera in two climate zones (Fig. 1): in the National
145 Park Pan de Azúcar (further as Pan de Azúcar or PdA) and the National Park La Campana (further as La
146 Campana or LC). The Las Lomas site in PdA is located in the arid climate zone of the Atacama Desert with a
147 precipitation rate of 12 mm year⁻¹, and it has a mean annual temperature of 16.8 °C (Übernicket et al. 2021a).
148 Here, the vegetation cover is below 5%, and it is dominated by small desert shrubs, several species of cacti
149 (*Eulychnia breviflora*, *Copiapoa atacamensis*) and biocrusts (Lehnert et al. 2018). LC is located in the
150 mediterranean climate zone with a precipitation rate of 367 mm year⁻¹ and a mean annual temperature of
151 14.1 °C (Übernicket et al. 2021a). LC is dominated by an evergreen sclerophyllous forest with endemic palm
152 trees, *Jubaea chilensis*. Both research sites have a granitic rock base, and the dominating soil texture is sandy
153 loam (Bernhard et al. 2018). In PdA, the study setup consisted of one north-facing and one south-facing
154 hillslope. The hillslope inclinations were ~20°, and a climate station was located ~15 km from the camera sites.
155 In LC, the setup consisted of two north-facing and one south-facing hillslopes. The hillslope inclinations were
156 ~25°, and a climate station was located ~250 m from the south-facing hillslope (Übernicket et al. 2021a).

157



158
 159 **Figure 1.** Location of the cameras and climate stations on which this study was based. Black points show the
 160 location of the research sites in Chile. The green points represent the camera plots, and the blue points the
 161 climate stations: (a) Location of study sites in Chile: PdA stands for Pan de Azúcar, LC for La Campana; (b)
 162 Study setup in Pan de Azúcar; (c) Study setup in LC. The background images in (b) and (c) are orthophotos
 163 created from WorldView-2 data from 19 July 2019. For exact latitude and longitude see Table A2.

164
 165 **2.1 Local burrowing animals**

166 We estimated the species of the local burrowing animals by conducting a literature review and using motion-
 167 activated wildlife traps. Among the most common vertebrate burrowing animals are in PdA carnivores of the
 168 family Canidae (*Lycalopex culpaeus*, *Lycalopex griseus*); rodents of the family Abrocomidae (*Abrocoma*
 169 *bennetti*), Chnichillidae (*Lagidium viscacia*), Cricetidae (*Abrothrix andinus*, *Phyllotis xanthopygus*, *Phyllotis*
 170 *limatus*, *Phyllotis darwini*) and Octogontidae (Cerquiera 1985, Jimenéz et al. 1992, Übernicket et al. 2021).
 171 The most common burrowing animals in LC are carnivores of the family Canidae and Methitidae, Lagomorpha
 172 of the family Leporidae (*Oryctolagus cuniculus*), and rodents of the families Cricetidae (*Abrothrix longipilis*,
 173 *Abrothrix olivaceus*, *Phyllotis darwini*), Muridae (*Mus musculus*) and Octogontidae (*Octogon degus*,
 174 *Spalacopus cyanus*) (Munoz-Pedreros et al. 2018, Übernicket et al. 2021). The motion-activated traps
 175 recorded several burrowing animals: *Lycalopex culpaeus*, *Oryctolagus cunniculum* and *Abrocoma bennettii*)
 176 (Figure 2). A list of the most common local burrowing animals is shown in Table 1.

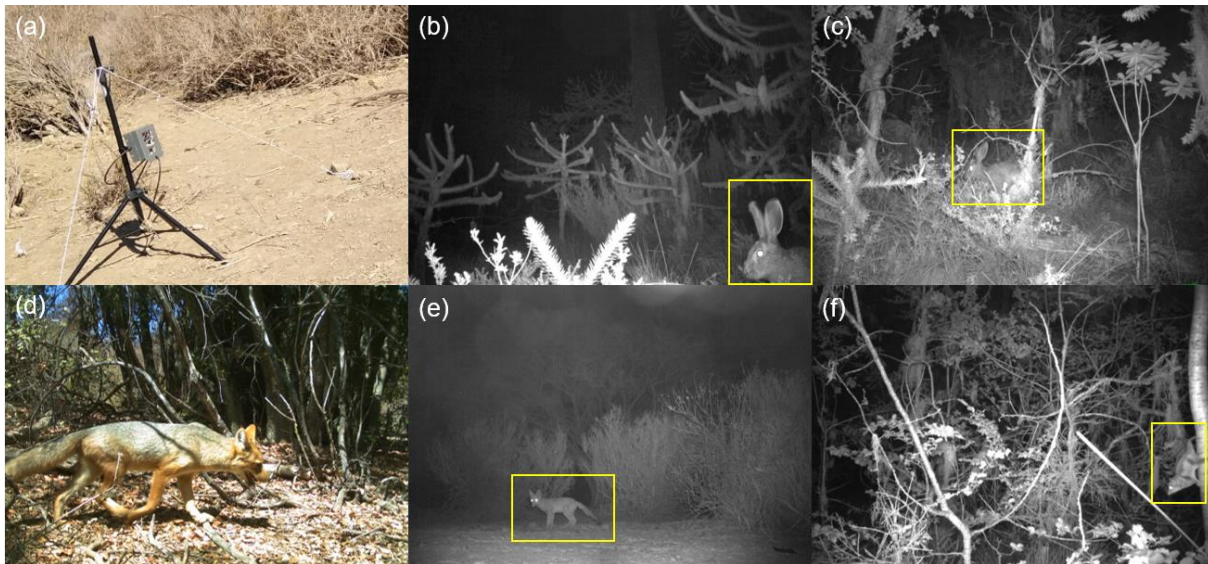
177
 178 **Table 1.** Most common local burrowing animals. The list includes data from our motion-activated wildlife traps,
 179 and reviewe by Übernicket et al. 2021, Cerquiera 1985, Jimenéz et al. 1992, Munoz-Pedreros et al. 2018). "X"
 180 shows at which site the species can be found.

Order	Family	Species	Common name	Site	
				PdA	LC
Carnivora	Canidae	<i>Lycalopex culpaeus</i>	Culpeo	X	X
Carnivora	Canidae	<i>Lycalopex griseus</i>	South-American grey fox	X	X

Carnivora	Methitidae	<i>Conepatus chinga</i>	Molina's Hog noised skunk		X
Lagomorpha	Leporidae	<i>Oryctolagus cuniculus</i>	European rabbit		X
Rodentia	Abrocomidae	<i>Abrocoma bennetti</i>	Bennett's chinchilla rat	X	X
Rodentia	Chinchillidae	<i>Lagidium viscacia</i>	Southern mountain vischacha	X	
Rodentia	Cricetidae	<i>Abrothrix andinus</i>	Andean grass mouse	X	
Rodentia	Cricetidae	<i>Abrothrix longipilis</i>	Long-haired mouse	X	X
Rodentia	Cricetidae	<i>Abrothrix olivaceus</i>	Olive grass mouse	X	X
Rodentia	Cricetidae	<i>Phyllotis darwini</i>	Darwin's leaf-eared mouse	X	X
Rodentia	Cricetidae	<i>Phyllotis xanthopygus</i>	Yellow leaf-eared mouse	X	
Rodentia	Cricetidae	<i>Phyllotis limatus</i>	Lima leaf-eared mouse	X	
Rodentia	Muridae	<i>Mus musculus</i>	Common house mouse	X	X
Rodentia	Octogontidae	<i>Octogon degus</i>	Degu (rat)	X	X
Rodentia	Octogontidae	<i>Spalacopus cyanus</i>	Coruro (rat)	X	X

181

182



183

184 **Figure 2.** Examples of local burrowing animals obtained by motion-activated traps. (a) Motion-activated

185 camera trap. (b) and (c) European rabbit (*Oryctolagus cuniculus*). (d) and (e) Culpeo (*Lycalopex culpaeus*).

186 (f) Bennett's chinchilla rat (*Abrocoma bennettii*). The yellow circle indicates the position of the animal on the

187 photo. Photos by Diana Kraus.

188

189 3. Methodology

190 3.1 Time-of-Flight (ToF) principle

191 A Time-of-Flight-based camera illuminates an object with a light source, usually in a non-visible
 192 spectrum, such as near-infrared, for a precise length of time. ToF cameras rely on the principle of measuring
 193 the phase shift, with different options to modulate the light source to be able to measure the phase shift. The
 194 here employed cameras used pulse-based modulation, meaning the light pulse was first emitted by the
 195 camera, then reflected from the surface, and finally measured by the camera using two temporary windows.
 196 The opening of the first window is synchronized with the pulse emission i.e. the receiver opens the window
 197 with the same Δt as the emitted pulse. Then, the second window is opened, for the same duration Δt , which is

198 synchronised with the closing of the first window. The first temporary window thus measures the incoming
 199 reflected light while the light pulse is also still emitting from the camera. The second temporary window
 200 measures the incoming reflected light when no pulse is emitting from the camera. The captured photon number
 201 (i.e. measured by electrical charge) in both windows can be related according to equation 1 and the distance
 202 from the camera to the object can then be calculated as follows:

$$203 \quad d = \frac{1}{2} * c * t * \left(\frac{g_1}{g_1 + g_2} \right) \quad . \quad (1)$$

204 In Eq. (1), d (m) is the distance from the camera to the object, c (m s⁻¹) is the speed of light (299,792,458 m s⁻¹)
 205 ¹), t (s) is the overall time of the illumination and measurement, g₁ is the ratio of the reflected photons to all
 206 photons accumulated in the first window, and g₂ the ratio of the reflected photons to all photons accumulated
 207 in the second window (Sarbolandi et al. 2018; Li 2014).

208 The sensor in our camera came from Texas Instruments and the data scan contained information on
 209 320 x 240 points. The camera field of view (FOV) and the spatial resolution of the scans depended on the
 210 height of the camera above the surface and camera orientation. The distance was calculated for every point,
 211 and the object was saved in binary format as a collection of 3D points with x-, y- and z-coordinates. The point
 212 clouds taken by the camera were transformed from the binary format to an ASCII format. Each point in the
 213 point cloud was assigned to an x-, y- and z-coordinate. The coordinates were distributed within a three-
 214 dimensional Euclidian space, with the point at the camera nadir (the centre of the camera sensor) being the
 215 point of origin of the 3D Cartesian coordinate system. x- and y-coordinates describe the distance to the point
 216 of origin (m). z-coordinate describes the distance (m) from the object to the camera. The lowest point of the
 217 scanned surface thus has the highest z-coordinate value.

218

219 **3.2 Data processing**

220 The distortion caused by the hillslope and the camera angle was corrected for each point cloud as
 221 follows:

$$222 \quad z_{cor} = z_{uncor} - \tan(\alpha + \beta) * (y_1 - y_i) \quad . \quad (2)$$

223 In Eq. (2), z_{cor} is the corrected distance (m) between the camera and surface (m), z_{uncor} is the uncorrected z-
 224 coordinate (m), α is the tilt angle of the camera (°), β is the surface inclination (°), and y_i (m) is the distance
 225 between each point, and the point with i) an y-coordinate = 0 and ii) the same x-coordinate as the respective
 226 point. The most frequent errors were identified and treated as follows. Due to the ambient light reaching the
 227 camera sensor, the z-coordinate values of some of the points were incorrect (scattering error). To remove this
 228 error, a threshold value was calculated for each point cloud:

$$229 \quad \Omega = mean_{z_{cor}-coordinates} \pm sd_{z_{cor}-coordinates} \quad . \quad (3)$$

230 In Eq. (3), Ω is the threshold value, mean_{z_{cor}-coordinate} is the average value, and sd_{z_{cor}-coordinate} is the standard
 231 deviation of the corrected z-coordinates (m). Then, all points with a z-coordinate above and below this value
 232 were deleted. Point clouds with more than 50% of points above the threshold value Ω were also not considered
 233 for further processing. A drift error occurred when the z-coordinate values of around one-third of the point
 234 clouds decreased by several centimetres from one point cloud to another. Here, the average z-coordinate of
 235 ten point clouds before and after the drift were calculated, and the difference was added to z-coordinates of
 236 the points affected by the drift. The corrected height values were then transformed into a digital surface model
 237 (DSM).

238

239 3.3 Accuracy of the ToF cameras

240 The accuracy of the ToF camera was tested under laboratory conditions by recreating similar surface
241 conditions as in the field (sloping surface, covered by sediment). An artificial mound using sediment extracted
242 from a riverbank in central Germany was used, mimicking a mound created by a burrowing animal. During the
243 test, the camera was installed 100 cm above the surface. The camera FOV was 3 m² and the scan spatial
244 resolution was 6 mm. The surface was scanned twice by the ToF camera. Then 100 – 450 cm³ of sediment
245 was manually extracted from the mound. The volume of the extracted sediment was measured by a measuring
246 cup. After extraction, the surface was again scanned twice by the camera. The experiment was repeated 45
247 times with varying amounts of extracted sediment. The scans were transformed to point clouds in VoxelViewer-
248 0.9.10, and the point clouds were corrected according to Eq. (2) and (3). The z-coordinates of the two point
249 clouds before and two point clouds after the extraction were averaged. The standard deviation of the z-
250 coordinate of the two scans was 0.06 cm. Figure A1 shows the spatially distributed standard deviation. The
251 deviation increases from the centre towards the corners of the scan. The mound was outlined and only the
252 points representing the mound were used in the further analysis. The point clouds were then transformed into
253 DSMs, and the differences between the time steps were calculated. A scan was taken of a smooth surface
254 (linoleum floor) and a point cloud was created from the data. Then, we fitted a plane into the point cloud and
255 calculated the distance between the plane and the camera sensor. The standard variation (0.17 cm) in the
256 distance measurements was saved. Solely, the differences between the DSMs below this variation were
257 considered in the calculation of the detected sediment extraction. The detected extracted sediment volume
258 was then calculated for each experiment as follows:

$$259 Vol_{detected} = \sum_p^1 (DSM_{before} - DSM_{after}) * res^2 \quad , \quad (4)$$

260 In Eq. (4), $Vol_{detected}$ is the volume of the extracted sediment as detected by the camera (cm³), p is the number
261 of pixels, DSM_{before} (cm) is the DSM calculated from the scan taken before the extraction, DSM_{after} (cm) is the
262 DSM calculated from the scan taken after the extraction, res (cm) is the resolution of the scan, which was 0.6
263 cm. To evaluate the camera's accuracy, the measured volume of the extracted sediment was compared to the
264 volume detected by the camera. The camera's accuracy was estimated between the detected volume and
265 measured volume as follows:

$$266 MAE = \sum_1^n \frac{(Vol_{detected} - Vol_{measured})}{area} \quad . \quad (5)$$

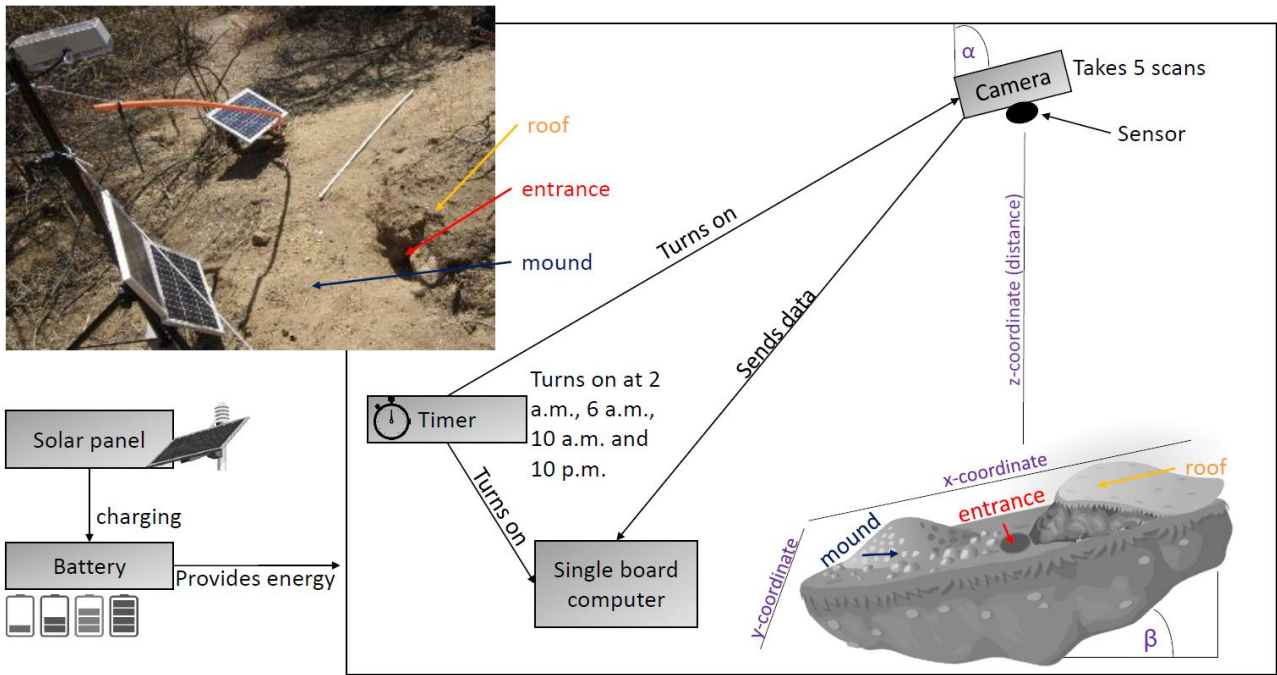
267 In Eq. (5), MAE (cm³/cm²) is the mean absolute error, n is the number of scans, $Vol_{measured}$ (cm³) is the volume
268 of the extracted sediment measured by the measuring cup, and the $area$ is the total surface area monitored
269 by the camera (cm²).

270

271 3.4 Installation of the cameras in the field

272 We installed 8 custom-tailored ToF-based cameras on 4 hillslopes in two climate zones in areas
273 including visible signs of bioturbation activity (burrows) and areas without visible signs of bioturbation (Fig. 3).
274 The cameras were installed in LC on the north-facing upper hillslope (LC-NU), north-facing lower hillslope (LC-
275 NL), south-facing upper hillslope (LC-SU) and the south-facing lower hillslope (LC-SL); in PdA on the north-
276 facing upper hillslope (PdA-NU), north-facing lower hillslope (PdA-NL), south-facing upper hillslope (PdA-SU)
277 and south-facing lower hillslope (PdA-SL). The custom-tailored cameras were installed during a field campaign
278 in March 2019, the monitoring took place for seven months, and the data were collected in October 2019. The
279 construction consisted of a 3D ToF-based sensor from Texas Instruments (Li, 2014), a RaspberryPi single board

280 computer (SBC), a timer, a 12 V 12 Ah battery and three 20 W solar panels for unattended operation (Fig. 2).
 281 Solar panels were located at the camera pole and were recharging the battery via a charge controller. The
 282 camera was located approximately one meter above the surface, facing the surface with a tilt angle of 10
 283 degrees. The timer was set to close the electric circuit 4 times a day: at 1 a.m., 5 a.m., 8 a.m. and 10 p.m. At
 284 these times, the camera and the computer were turned on for 15 minutes. The camera turned on and took five
 285 scans delayed one second from each other and sent them to the SBC. Each camera had its own WiFi (Wireless
 286 Fidelity) and the data could be read from the SBC via Secure Shell (SSH). The cameras collected the data for
 287 the time period of 7 months.
 288



289
 290 **Figure 3.** Scheme and photo example of a Time-of-Flight-based camera installation in the field. The photo
 291 example is from upper north-facing hillslope in La Campana. Black boxes describe single installation parts.
 292 Purple descriptions are the variables needed for the correction of the scans. Roof, entrance and mound
 293 describe parts of the burrow. The x-, y- and z-coordinates are 3D coordinates identifying the position of each
 294 point in space, where the x-coordinate is the length, y-coordinate is the width and the z-coordinate is the
 295 distance between the camera sensor and the surface. α is the inclination of the camera, and β is the surface
 296 inclination.

297
 298 **3.5 Delineation of burrows and burrow embedding areas**

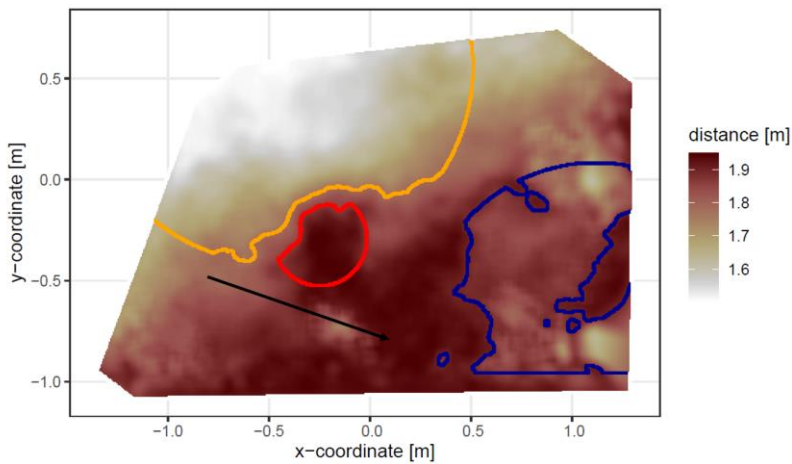
299 The surface area scanned by the cameras was divided by a delineation scheme into burrows (B) and
 300 burrow embedding areas (EM). The burrows included three sub-areas: (i) mound (M), (ii) entrance (E) and (iii)
 301 burrow roof (R). “Mound” describes the sediment excavated by the animal while digging the burrow. “Entrance”
 302 describes the entry to the animal burrow up to the depth possible to obtain via the camera. “Burrow roof”
 303 describes the part of the sediment above and uphill the burrow entrance (BANCROFT et al. 2004). During the
 304 burrow’s creation, sediment was not only excavated but also pushed aside and uphill the entrance, which
 305 created the burrow roof. We assume that this elevated microtopographical feature then forms an obstacle for
 306 sediment transported from uphill, which leads to its accumulation in this area. The remaining surface within

307 the camera's FOV was burrow embedding area. Please note, that this area may still be affected by the
308 burrowing activity of the animal and is not completely unaffected by the animal.

309 For the delineation, we used the DSM calculated from the point cloud, and a slope layer calculated
310 from the DSM (Horn 1981). The DSM had a size of 4 m² a resolution of 0.6 cm. Entrance was assigned to an
311 area determined by a search algorithm starting at the lowest point of the DSM (pixel with the highest z-
312 coordinate value). We increased the circular buffer around the starting point by one pixel until the average
313 depth of the new buffer points was not higher than the height of the camera above the surface, or until the
314 slope of at least 50% of the new buffer points was not 0. Then, we masked all pixels within the buffer with a
315 depth lower than the average depth of the points within the buffer, which had a slope that was 0. The remaining
316 pixels belonged to the entrance area. Then, the surface scan was divided into an uphill and downhill part with
317 regards to the entrance position. Both the uphill and the downhill parts were subdivided into 16 squares, so
318 that each of the four quadrants within the 2D grid (x- and y-axis) contained four squares. The squares had size
319 of 0.5 m².

320 To delineate the mound in the downhill part, we first identified the highest points (pixel with the lowest z-
321 coordinate value) within all 16 squares. We then calculated the distance of these maxima to the entrance, and
322 the pixel located nearest to the entrance was identified as the highest point of the mound (i.e., seed point).
323 Consecutively, we increased the circular buffer around the seed point by one pixel until the average depth of
324 the new buffer points was not lower than the height of the camera above the surface, or until the slope of at
325 least 50% of the new buffer points was not 0. Then, we masked all pixels within the buffer with a depth higher
326 than the average depth of the points within the buffer, which had a slope that was 0. The remaining pixels were
327 classified as mound area. To delineate burrow roof, we used the same approach as for the delineation of
328 mound and applied it on the uphill part of the surface scan. We used the DEM and slope layers for the
329 delineation for several reasons. The distance from the surface to the camera was the most important parameter
330 to derive (i) the deepest point of the entrance and (ii) the highest point of the mound or burrow roof, as this
331 was (mostly) the closest point to the camera. After the angle correction of the z-coordinate according to chapter
332 3.2., the surface inclination of the areas without burrow was 0°, while the angle between the border of the
333 burrow entrance or mound and the burrow embedding surface was above 0°. Because neither the entrance
334 nor the mound have a perfect circular form, we would largely overestimate or underestimate the entrance or
335 mound size. Overestimate by not stopping the search algorithm until the angle between all new points of the
336 buffer to the rest of the buffer was 0°. Underestimate by stopping the algorithm when the angle of one point of
337 the buffer to the nearest point of the buffer was 0°. The value of 50% thus minimized the error. All pixels that
338 were not classified during the entire delineation process were treated as burrow embedding areas.

339 The position and the boundaries of entrance, mound and burrow roof were validated visually (Fig. 4 and A2).



340
 341 **Figure 4.** Corrected digital surface model of the camera on the upper north-facing hillslope in La Campana
 342 with delineated areas. The point of origin of the coordinate system is at the camera nadir. Distance refers to
 343 the distance between surface and camera. The red line delineates the burrow entrance, blue the mound and
 344 orange the burrow roof. The area which was outside of any delineated area was classified as burrow
 345 embedding area. The arrow indicates a downhill direction of the hillslope.

346
 347 In LC, the burrows always consisted of an entrance, mound and burrow roof. In PdA, there was no
 348 burrow roof on the upper hillslopes. Burrows without a burrow roof were located on shallower parts of the
 349 hillslopes (up to an inclination of 5°), and the angle of the burrow entrance to the ground was $\sim 90^\circ$. Burrows
 350 with a burrow roof were located on steeper parts of the hillslopes (with an inclination above 5°), and the angle
 351 of the burrow entrance to the ground was $\sim 45^\circ$.

352
 353 **3.6 Calculation of animal-caused and rainfall-caused sediment redistribution**

354 We pairwise compared the DSMs of each scan with the scan saved before and identified 3 types of
 355 sediment redistribution which occurred in the time period between these images. The 3 types of redistribution
 356 were: a) animal caused; b) rainfall-caused; c) both animal and rainfall caused.

357 The animal-caused sediment redistribution occurred when the animal actively reworked sediment
 358 within its burrow. Following five prerequisites had to be met when the sediment redistribution was caused
 359 solely by the animal: (i) as the animal excavates sediment from the entrance, the depth of the entrance must
 360 increase in the second scan; (ii) as the excavated sediment accumulates on the mound, the height of the
 361 mound must increase in the second scan; (iii) as the burrowing might lead to an expansion or a collapse of the
 362 burrow roof, an increase or decrease of the burrow roof must occur between the scans; (iv) as the animal only
 363 digs within his burrow, no changes must occur between the two scans within the burrow embedding area by
 364 the animal; (v) no rainfall occurred during this period.

365 The rainfall-caused sediment redistribution was calculated as follows: From the data from the climate
 366 stations (Übernicker et al. 2021a), we calculated the daily precipitation in mm. The sediment redistribution
 367 recorded immediately and within five scans before and after a rainfall event is defined to be the result of the
 368 rainfall event. This was necessary as the climate stations are located up to a 15 km distance from the cameras
 369 (Fig. 1). To attribute sediment redistribution to rainfall event, three preconditions had to be met: (i) A rainfall
 370 event occurred; (ii) sediment is eroded from burrow roof, mound and the embedding area; (iii) sediment is
 371 accumulated within the burrow entrance.

372 To attribute sediment redistribution to a combination of animal activity and rainfall, four preconditions
 373 had to be met: (i) A rainfall event occurred; (ii) sediment is eroded from embedding area; (iii) the height of
 374 burrow roof and mound decreased or increased; (iv) the depth of burrow entrance increased.

375 The animal-caused sediment redistribution was calculated as the sediment volume excavated from
 376 the entrance. Animal excavation always increased depth of the burrow entrance. The rainfall-caused sediment
 377 redistribution was calculated as the sediment volume which eroded from the burrow roof and mound. During
 378 a rainfall event, sediment eroding from burrow roof might accumulate within burrow entrances. In this case,
 379 the depth of the burrow entrance decreased. No sediment could erode from the entrance during a rainfall
 380 event. Decreased depth of a burrow entrance always points to sediment redistribution caused by rainfall,
 381 increased depth of burrow entrance always means redistribution by animals. Rainfall-caused redistribution
 382 always occurred before animal-caused redistribution, as without erosion caused by rainfall, the animals did not
 383 need to reconstruct their burrows.

384

385 **3.7 Calculation of daily sediment mass balance budget**

386 The volume of the redistributed sediment was calculated daily and was then cumulated from the first
 387 day of monitoring. For the calculation of the daily sediment redistribution, the change in the surface level
 388 detected by the camera was calculated first. For each day, the scans from the day before and after the
 389 respective day were averaged and subtracted. The average standard deviation of the z-coordinate of these
 390 scans was 0.06 cm. As described in Section 2.2., all values with a difference below and above the threshold
 391 value of 0.2 cm were set to 0. The redistributed sediment volume was then calculated from the surface change
 392 for each pixel as follows:

$$393 \text{Vol}_{\text{redistributed}} = (S_b - S_a) * \text{res}^2 \quad (6)$$

394 In Eq. (6), $\text{Vol}_{\text{redistributed}}$ ($\text{cm}^3 \text{ pixel}^{-1}$) is the volume of the calculated redistributed sediment, S_b (cm) the scan
 395 before, S_a (cm) the scan after the rainfall event and res is the spatial resolution (cm). Using the daily volume
 396 of the redistributed sediment per pixel, we calculated the daily mass balance budget by summing the volume
 397 of sediment eroding or accumulating within each delineated area.

398

399 **3.8 Calculation of the overall volume of redistributed sediment after the period of 7 months**

400 From the camera data, we calculated the average cumulative volume of redistributed sediment for the
 401 period of 7 months within burrows ($\text{Vol}_{\text{burrows}}$ ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$)) and burrow embedding ($\text{Vol}_{\text{embedding}}$ ($\text{cm}^3 \text{ cm}^{-2}$
 402 year^{-1})) areas and the average sediment volume redistributed (excavated) by the animal (Vol_{exc} ($\text{cm}^3 \text{ cm}^{-2}$
 403 year^{-1})), separately for each site. We estimated the volume of sediment that was redistributed during rainfall
 404 events due to the presence of the burrow (Vol_{add} ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$)). Vol_{add} was calculated as the difference in
 405 the redistributed sediment volume between burrows and burrow embedding areas according to Eq. (7).

$$406 \text{Vol}_{\text{add}} = (\text{Vol}_{\text{affected}} - \text{Vol}_{\text{unaffected}}) * 1.71 \quad (7)$$

407 Additionally, we calculated the average volume of the redistributed sediment per burrow ($\text{Vol}_{\text{per burrow}}$ [cm^3
 408 $\text{ burrow}^{-1} \text{ year}^{-1}$]).

$$409 \text{Vol}_{\text{per burrow}} = (\text{Area}_{\text{burrow}} * \text{Vol}) * 1.71 \quad (8)$$

410 In Eq. (8), $\text{Area}_{\text{burrow}}$ (cm^2) is the average size of the burrows that are monitored by the cameras; Vol is $\text{Vol}_{\text{burrow}}$
 411 ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$), Vol_{exc} ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$) or Vol_{add} ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$).

412 We then upscaled the $\text{Vol}_{\text{burrow}}$ ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$), Vol_{exc} ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$) and Vol_{add} ($\text{cm}^3 \text{ cm}^{-2} \text{ year}^{-1}$)
 413 to the hillslope using the following approach. Hillslope-wide upscaling of the results generated in this study

414 was performed by using a previous estimation of vertebrate burrow density (Grigusova et al. 2021). In this
 415 study, the density of burrows was measured in situ within eighty 100 m² plots and then upscaled to the same
 416 hillslopes on which the cameras were located by applying machine-learning methods, using the UAV-data as
 417 predictors. For upscaling, we applied a random forest model with recursive feature elimination. The model was
 418 validated by a repeated Leave-One-Out cross validation. The density of vertebrate burrows was between 6
 419 and 12 100 m² in LC and between 0 and 12 100 m⁻² in Pan de Azúcar. Using the hillslope-wide predicted
 420 vertebrate burrow densities ($Dens_{burrow}$ (number of burrows 100 m⁻²)) from Grigusova et al. 2021, we estimated
 421 the volume of redistributed sediment for each pixel of the raster layers ($Vol_{per\ pixel}$ (cm³ m⁻² year⁻¹)) according
 422 to Eq. (9):

$$423 \quad Vol_{per\ pixel} = Vol_{per\ burrow} * Dens_{burrow} * 1.71 \quad (9)$$

424 The average hillslope-wide volume of redistributed sediment ($Vol_{hillslope-wide}$ (m³ ha⁻¹ year⁻¹)) was then
 425 estimated as follows:

$$426 \quad Vol_{hillslope-wide} = \sum_1^m Vol_{per\ pixel} * 0.001 * 1.71 \quad , \quad (10)$$

427 In Eq (10), m is the number of pixels.

428

429 4. Results

430 4.1 Camera accuracy and data availability

431 The accuracy between the measured extracted sediment volume and sediment volume calculated
 432 from the camera scans was very high (MAE = 0.023 cm³ cm⁻², R² = 0.77, SD = 0.02 cm³ cm⁻², Fig. A3). The
 433 accuracy between the calculated and measured extracted sediment was higher when the two scans taken
 434 before as well as after the extraction of the sediment were averaged and the sediment volume was estimated
 435 using these averaged scans. When calculating the redistributed sediment from solely one scan before and
 436 after extraction, the accuracy slightly decreased (MAE = 0.081 cm³ cm⁻², R² = 0.64). The cameras tended to
 437 overestimate the volume of redistributed sediment. Six out of eight custom-tailored cameras collected data
 438 over the seven-month period (Table A2). One camera collected data for a period of three months and one
 439 camera stopped working a few days after installation. The quantity of usable point clouds taken at 1 a.m., 5
 440 a.m. and 10 p.m. was higher than of point clouds taken at 8 a.m. Approximately 20% of points was removed
 441 from the point clouds before final analysis due to the high scattering at the point cloud corners. After data
 442 filtering (see Section 3.2.), 1326 scans were usable and for 86% of the days, at least one usable scan was
 443 available. The usable scans were distributed continuously within the monitoring period.

444

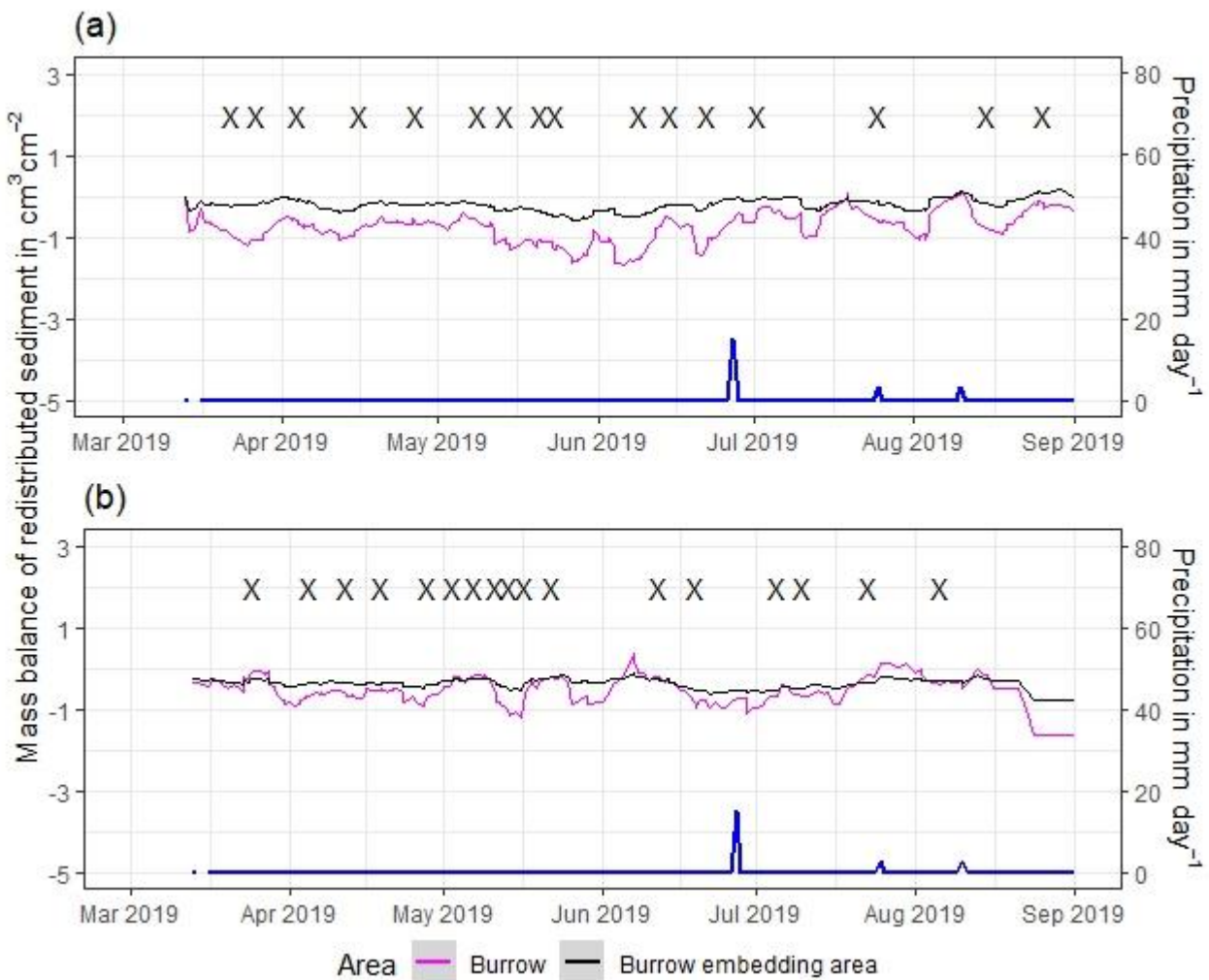
445 4.3 Mass balance of redistributed sediment

446 The cameras detected (i) sediment redistribution directly following rainfall events and (ii) due to the
 447 burrowing activity in times without rainfall (Fig. 5, A4 and A5). In all cases, burrows (entrance, burrow roof and
 448 mound) exhibited higher sediment redistribution rates than burrow embedding areas. In addition, the volume
 449 of redistributed sediment by animal activity was higher after a rainfall event occurred.

450 In the following, the dynamics are exemplary explained for four cameras. Animal burrowing activity
 451 was detected seven times by the camera LC NU (Fig. 5a, A4, A5) during the monitoring period, by an increase
 452 in sediment volume in the area delineated as mound. Simultaneously, the burrow entrance showed signs of
 453 modification and sediment accumulation, but these changes were less clear. Overall, the volume of the
 454 excavated soil varied. From April until June, up to 0.5 cm³ cm⁻² of sediment was excavated by the animal and
 455 accumulated on the mound. From June until September, animal burrowing activity was detected at four time

456 slots (5 June 2019, 9 June 2019, 1 July 2019 and 18 August 2019) and sediment volume of up to $2 \text{ cm}^3 \text{ cm}^{-2}$
 457 accumulated each time on the mound, burrow roof and within the entrance. During the rainfall events of up to
 458 20 mm day^{-1} on 16 June 2019, 27 mm day^{-1} on 29 June 2019 and 7 mm day^{-1} on 13 July 2019, sediment
 459 volume of up to $4 \text{ cm}^3 \text{ cm}^{-2}$ eroded, especially from the burrow roof and the mound while a sediment volume
 460 of up to $1 \text{ cm}^3 \text{ cm}^{-2}$ accumulated within the entrance during each rainfall event. Camera LC-SL (Fig. A4, A5)
 461 showed burrowing activities eight times and sediment volumes of up to $3 \text{ cm}^3 \text{ cm}^{-2}$ accumulated within the
 462 entrance and burrow roof. The camera detected sediment erosion of up to $2 \text{ cm}^3 \text{ cm}^{-2}$ after a rainfall event of
 463 27 mm day^{-1} on 27 July 2019. On the south-upper hillslope, the camera detected animal burrowing activity six
 464 times, with a sediment accumulation of up to $3 \text{ cm}^3 \text{ cm}^{-2}$ (Fig. A2 and A3).

465 In contrast, camera PdA-NU pointed to animal burrowing activity up to 15 times where up to $1 \text{ cm}^3 \text{ cm}^{-2}$
 466 2 of sediment volume was redistributed from the entrance to the mound (Fig. 5b, A4, A5). At the end of June
 467 on 27 June 2019, a rainfall event of 1.5 mm day^{-1} occurred and up to $2 \text{ cm}^3 \text{ cm}^{-2}$ of sediment eroded from the
 468 burrow roof and accumulated within the burrow entrance. We observed increased sediment redistribution by
 469 the animal after the rainfall events. Camera PdA-SL evenly revealed animal burrowing activity up to 15 times
 470 ((Fig. A4, A5)). The burrowing had a strong effect on the sediment redistribution. The rainfall event of 1.5 mm
 471 day^{-1} on 27 June 2019 did not cause any detectable surface change.



472
 473
 474 **Figure 5.** Examples of the mass balance of redistributed sediment for burrows and burrow embedding areas
 475 (a) The record of the camera on the upper north-facing hillslope in La Campana showed that larger rainfall

476 events cause a negative sediment balance (sediment loss), followed by a phase of positive sediment mass
 477 balance after approximately 3 days due to sediment excavation; (b) The record of the camera on the upper
 478 north-facing in Pan de Azúcar hillslope showed a similar pattern to the camera on the upper north-facing
 479 hillslope, but the phase of positive mass balance was delayed in comparison. The blue line is the daily
 480 precipitation in mm day^{-1} , and “X” marks the days at which animal burrowing activity was detected. Mass
 481 balances for all cameras are displayed in Fig. A2 and A3.

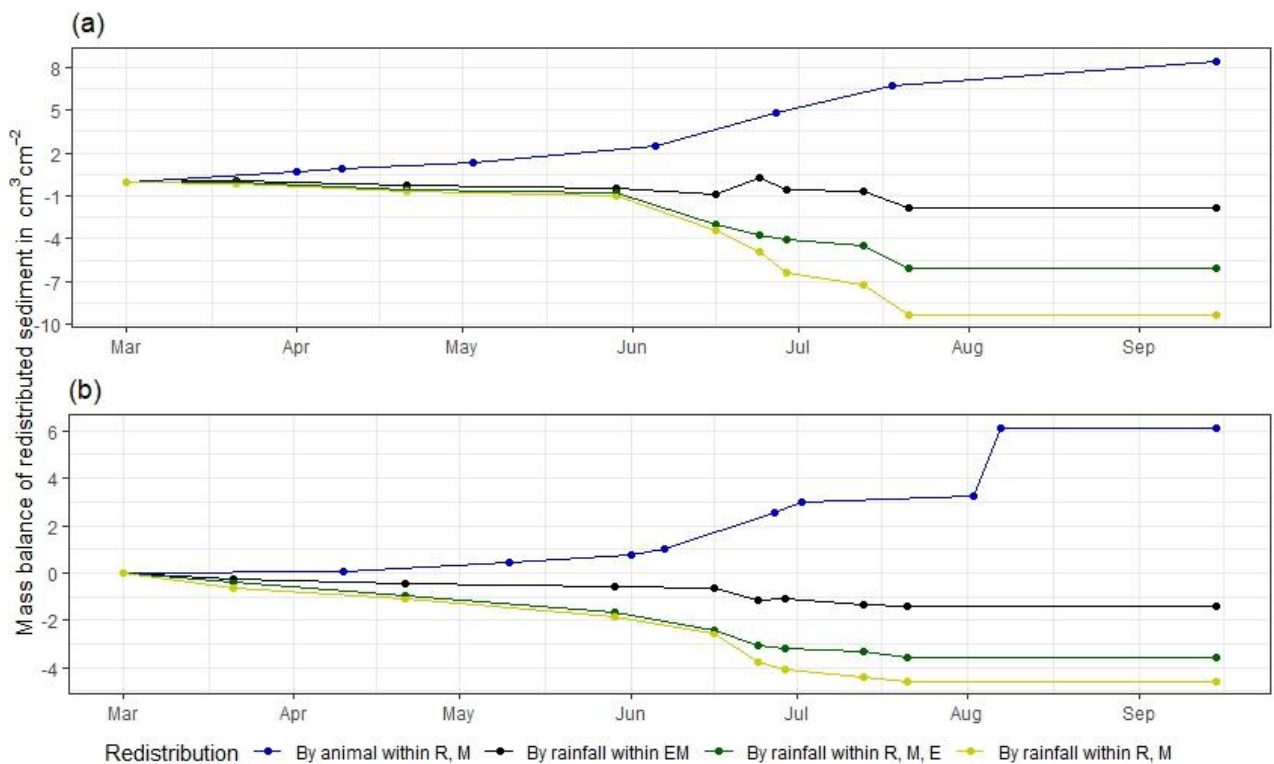
482

483 The analysis of cumulative volume of the redistributed sediment caused by burrowing animal activity
 484 and rainfall over the monitored period of seven months for all eight cameras showed a heterogeneous pattern.

485 In LC, the cumulative volume of the sediment excavated by the animal within the burrow roof and
 486 mound increased continuously (Fig. 6, A7). Especially between the rainfall events from June until August, a
 487 cumulative volume of on average $6.5 \text{ cm}^3 \text{ cm}^{-2}$ was excavated by the animal. We calculated that, on average,
 488 $8.53 \text{ cm}^3 \text{ cm}^{-2}$ cumulatively eroded from the burrow roof and mound; while $2.44 \text{ cm}^3 \text{ cm}^{-2}$ sediment volume
 489 accumulated within the entrance (Fig. 6, A7). These results indicate that 28% of sediment eroding from the
 490 burrow roof accumulated within the entrance, while over 62% of sediment eroded downhill. Averaged over all
 491 camera scans, 338% more sediment was redistributed by rain within burrow compared to the burrow
 492 embedding area (Fig. 7).

493 In PdA, cameras continuously detected animal burrowing activity and excavation of the sediment (Fig.
 494 A7). The volume of the detected excavated sediment increased steadily within all cameras. The cumulative
 495 sediment accumulation surpasses the sediment eroded due to the rainfall. The volume of the sediment eroded
 496 within the burrows was 40% higher than within the burrow embedding areas. The results show that
 497 approximately 50% of the eroded sediment accumulated within the entrance (Fig. 7).

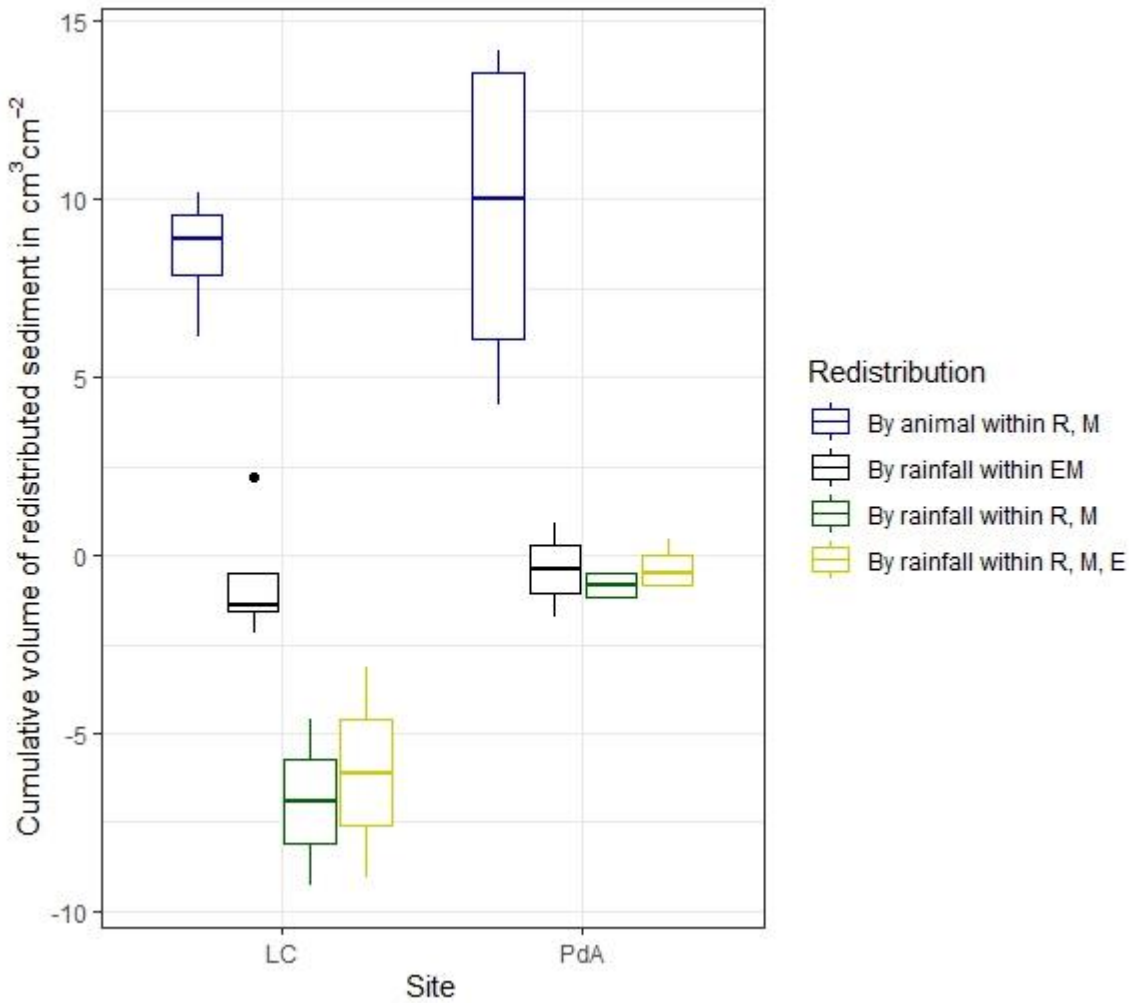
498



499

500 **Figure 6.** Examples of the cumulative volume of redistributed sediment within burrows and burrow embedding
 501 areas caused by animal burrowing activity or rainfall in mediterranean La Campana: (a) Upper north-facing

502 hillslope; (b) Lower south-facing hillslope. Positive values indicate sediment accumulation. Negative values
 503 indicate sediment erosion. E is the burrow entrance; M is the mound; R is burrow roof; EM is the burrow
 504 embedding area. Cumulative volumes for all cameras are in Fig. A7.



505
 506 **Figure 7.** Cumulative volume of the redistributed sediment for all cameras. Positive values indicate sediment
 507 accumulation. Negative values indicate sediment erosion. Whiskers indicate the median of sediment
 508 redistribution. E is the burrow entrance; M the mound; R is the burrow roof; EM is burrow embedding area; LC
 509 stands for National Park La Campana in the mediterranean climate zone; PdA stands for National Park PdA in
 510 the arid climate zone.

511
 512 **4.4 Volume of redistributed sediment**

513 The average size of the burrows was 84.36 cm² (SD = 32.54 cm²) in LC and 91.35 cm² in PdA (SD = 8.53
 514 cm²). The animals burrowed on average 1.2 times month⁻¹ in LC and 2.33 times month⁻¹ in PdA. The volume
 515 of the excavated sediment was 102.22 cm³ month⁻¹ in LC and 124.89 cm³ month⁻¹ in PdA. Each time the
 516 animals burrowed, they excavated 42 cm³ sediment volume in LC and 14.33 cm³ sediment volume in PdA.
 517 The burrowing intensity increased in winter after the rainfall occurrences in LC and stayed constant during the
 518 whole monitoring period in PdA. The burrows deteriorate after rainfall events with a rate of 73.03 cm³ month⁻¹
 519 or 63.90 cm³ event⁻¹ in LC and 10.53 cm³ month or 24.57 cm³ event⁻¹.

520 The overall volume of the sediment excavated by the animal and redistributed during rainfall events
 521 varied between the sites (Table 1). The volume of the sediment redistributed by the animal was lower in LC

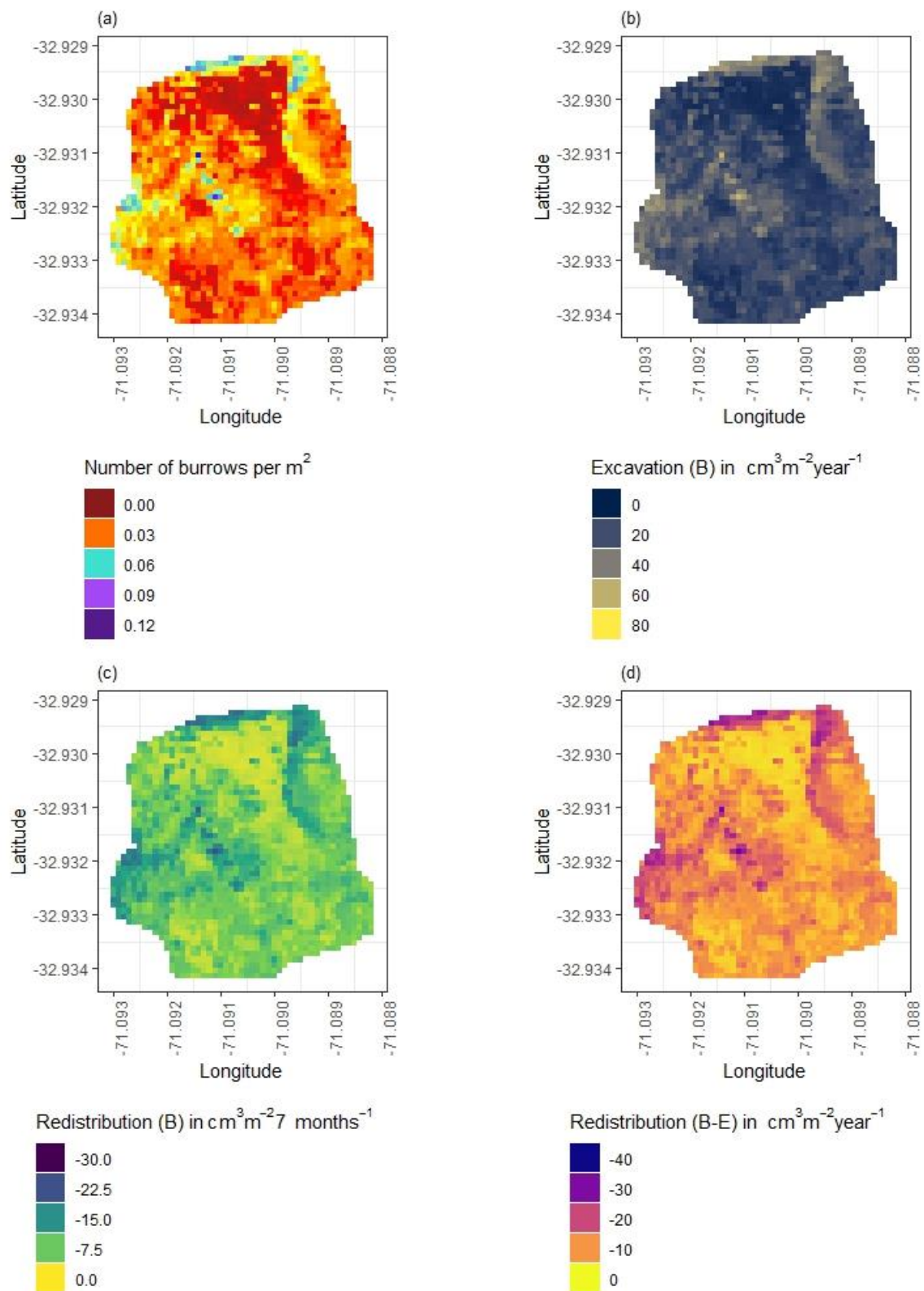
522 than in PdA. However, on the hillslope scale, a higher total area-wide volume of excavation was calculated for
 523 LC compared to PdA, due to the higher burrow density in LC. The volume of the sediment redistributed within
 524 burrows during rainfall events was higher in LC than in PdA. The volume of additionally redistributed sediment
 525 due to the presence of burrows was higher in LC than in PdA (Table 2, Fig. 8).

526

527 **Table 2.** Summary of the volume of redistributed sediment, according to area and disturbance type. Vol_{exc}
 528 describes volume of the sediment excavated by the animals. Vol_{burrow} describes volume of the sediment
 529 redistributed during rainfall events within burrows. Vol_{add} describes the difference in redistributed sediment
 530 volume within burrows and burrow embedding areas during rainfall.

Disturbance	Area	PdA	LC
Vol_{exc}	Affected area	16.41 cm ³ cm ⁻² year ⁻¹	14.62 cm ³ cm ⁻² year ⁻¹
	Per burrow	1498.66 cm ³ burrow ⁻¹ year ⁻¹	1226.61 cm ³ burrow ⁻¹ year ⁻¹
	Hillslope-wide	0.18 m ³ ha ⁻¹ year ⁻¹	0.67 m ³ ha ⁻¹ year ⁻¹
$Vol_{affected}$	Affected area	-1.97 cm ³ cm ⁻² year ⁻¹	-10.44 cm ³ cm ⁻² year ⁻¹
	Per burrow	-126.36 cm ³ burrow ⁻¹ year ⁻¹	-876.38 cm ³ burrow ⁻¹ year ⁻¹
	Hillslope-wide	-0.05 m ³ ha ⁻¹ year ⁻¹	-0.48 m ³ ha ⁻¹ year ⁻¹
Vol_{add}	Affected area	-1.18 cm ³ cm ⁻² year ⁻¹	-7.37 cm ³ cm ⁻² year ⁻¹
	Per burrow	-48.36 cm ³ burrow ⁻¹ year ⁻¹	-619.2 cm ³ burrow ⁻¹ year ⁻¹
	Hillslope-wide	-0.02 m ³ ha ⁻¹ year ⁻¹	-0.34 m ³ ha ⁻¹ year ⁻¹

531



532

533 **Figure 8.** Example of the hillslope-wide volume of redistributed sediment on the south-facing hillslope in La
 534 Campana: (a) Density of burrows as estimated by Grigusova et al. (2021); (b) Volume of the sediment
 535 excavated by the animals; (c) Volume of the sediment redistributed during rainfall events within burrows; (d)
 536 Volume of additionally redistributed sediment during rainfall events due to the presence of the burrows. The
 537 values were calculated per burrow as stated in Section 3.7. by subtracting the sediment volume redistributed
 538 within burrows from the sediment volume redistributed within burrow embedding area and then upscaled. The
 539 letters in brackets indicate if the upscaling was conducted using data from burrows or burrow embedding areas.

540 “B” stands for burrow. By “EM-B”, the redistribution calculated within burrow embedding areas was subtracted
541 from the redistribution calculated within burrows to obtain the additional volume of redistributed sediment due
542 to the burrows’ presence.

543

544

545 **5. Discussion**

546 Our results showed that the custom-made ToF device is a suitable tool for high-resolution, automated
547 monitoring of surface changes, applicable also in remote areas. The ability of a continuous observation of
548 sediment redistribution over a longer time during our study provided new insights into the importance of
549 burrowing animals for sediment redistribution. Our research reveals that the presence of vertebrate burrows
550 increases hillslope sediment redistribution rates much more than previously assumed (up to 208%). We
551 showed that the quantity of animal-related sediment redistribution, however, varied with rainfall occurrence,
552 with an increase in sediment redistribution between 40% in the arid research area and 338% percent in the
553 mediterranean research area.

554

555 **5.1 Suitability of the ToF method for surface monitoring**

556 The here proposed monitoring technique enables an automatic monitoring of surface changes on a
557 microtopographic scale, and its measurement continuity allows for the analysis of ongoing
558 biogeomorphological processes in high temporal resolution.

559 With regard to the costs, measurement frequency and sampling autonomy, the custom-made ToF
560 device stands in contrast to earlier studies that used laser scanning technology to monitor microtopographic
561 changes (Table A5). Previous studies mainly applied expensive laser scanning for the estimation of sediment
562 redistribution, and the research sites had to be personally revisited for each of the measurements
563 (Nasermoaddeli und Pasche 2008; Eltner et al. 2016a; Eltner et al. 2016b; Hänsel et al. 2016). The estimated
564 costs in studies using time-lapse photogrammetry were similar to our study (up to 5000 USD) (James und
565 Robson 2014; Galland et al. 2016; MALLALIEU et al. 2017; Eltner et al. 2017; Kromer et al. 2019; Blanch et
566 al. 2021). However, for time-lapse monitoring, several devices needing different viewing angles increases
567 installation efforts significantly.

568 In terms of data quality, our ToF device is more precise or comparable to those employed in other
569 studies. The accuracy of the camera ($R^2 = 0.77$) was in the range of previous studies ($R^2 = 0.26$ – 0.83 (Eitel et
570 al. 2011), Table A5). The horizontal point spacing of our cameras was 0.32 cm, and the maximum number of
571 points per cm^2 was 8.5. These values are similar to previous studies in which the used devices had a horizontal
572 point spacing in the range of 0.25–0.57 cm (Kaiser et al. 2014; Nasermoaddeli und Pasche 2008)) (Table A5),
573 and the maximum number of points per cm^2 in a range of 1 point–25 points cm^{-2} (Eitel et al. 2011; Longoni et
574 al. 2016) (Table A5).

575 Our cameras tended to slightly overestimate or underestimate the volume of redistributed sediment.
576 This error occurs when the pulse reflects from several vertical objects such as walls or, in our case, branches
577 or stones and then enters the camera sensor. This phenomenon was also observed in previous studies
578 applying laser scanners and is inevitable if the goal is to study surface changes under natural field conditions
579 (Kukko und Hyypä 2009; Ashcroft et al. 2014). During operation of the cameras, we learnt that our newly
580 developed instruments are particularly capable of delivering usable scans at night. This is likely due to the

581 strong scattered sunlight reaching the camera sensor during the day, blurring the data (Li 2014). Thus, in future
582 studies, we recommend focusing on nocturnal operation to prevent light contamination from the surroundings.

583 We could thus prove that ToF cameras are a suitable and cost-effective method for a continuous
584 monitoring of sediment redistribution at a microtopographic scale without the need of time, labour and cost
585 intensive laser scanning/time-lapse photogrammetry campaigns.

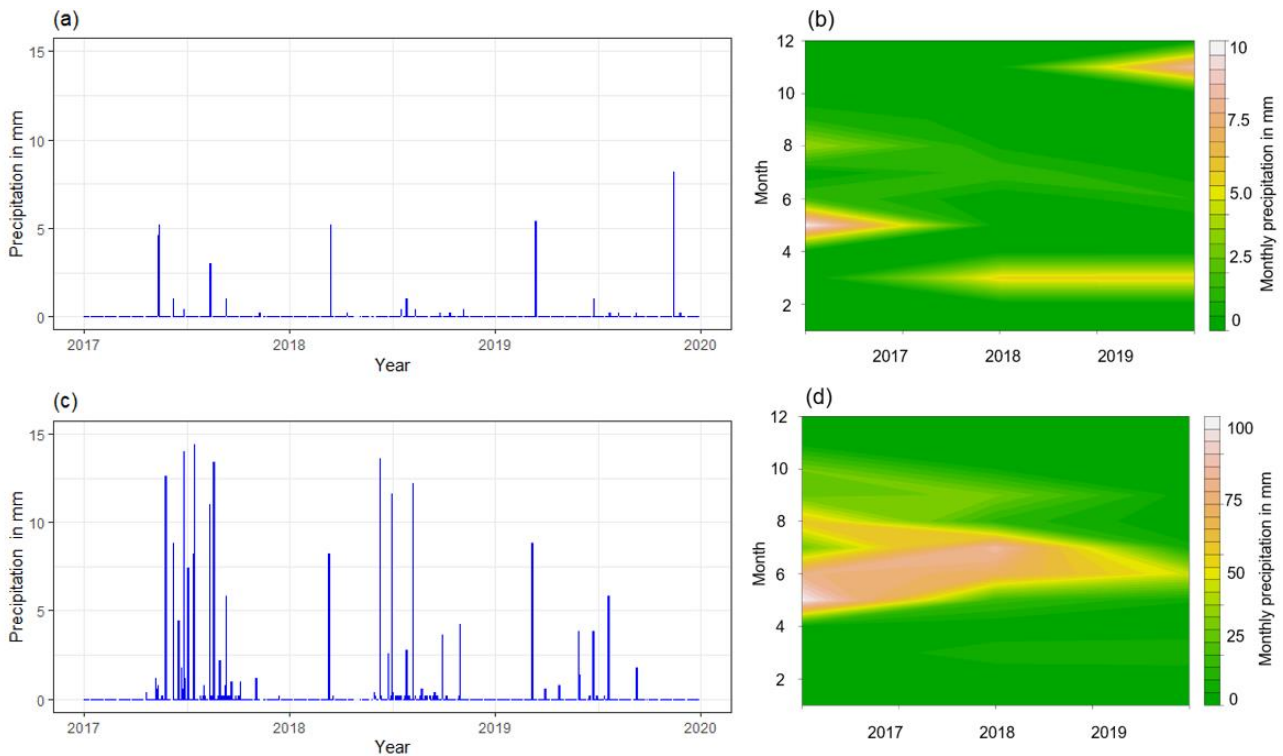
586

587 **5.2 Burrowing cycles and climate variability**

588 Our monitoring covered a time period from March 2019 until October 2019. To use our temporally
589 limited data on an annual basis, it is important to consider the annual course of burrowing. The burrowing
590 activity of the animals varies between the families. The most common burrowing families within our study area
591 are Canidae, Cricetidae and Octogontidae (Tab 1). Highest burrowing activity of the family Canidae was
592 observed during breeding time between August until October (Monteverde und Piudo 2011). The peak of the
593 burrowing activity of the family Cricetidae is in austral winter from June to August (Vargas Cademartori 2003).
594 The burrowing activity of the rodents of the family Octogontidae depends largely on the habitat productivity.
595 They tend to be most active from March to July (Malizia 1998). The time span, in which the most common
596 burrowing families within our study area are active, March until October, was fully covered by our dataset. We
597 therefore didn't miss any burrowing activity within the current year.

598 Our monitoring period lasts from March until October. In 2019, March and April were dry season; May
599 until August were rain season; and September i again dry season. Our monitoring time period is thus
600 approximately evenly covered by dry and rain season. Furthermore, the annual variability and the variability
601 within our monitoring time period are similar. In LC, the variance in year 2019 was 3.2 mm with a standard
602 deviation (SD) of 1.8 mm. The variance between March 2019 and October 2019 (e.g. our monitored time
603 period) was 3.3 mm and the SD was as well 1.8 mm. The annual variance in PdA was 0.3 mm with an SD of
604 0.5 mm. The variance between March 2019 and October 2019 was 0.1 mm and the SD was as well 0.4 mm.
605 We can therefore upscale our values to the time period of one year.

606 As for the precipitation, in the years 2017, 2018 and 2019, strong inter-annual climate variability and
607 strongly declining precipitation rate was detected. In LC, the precipitation amounted 351 mm in 2017, 245 mm
608 in 2018 and 72 mm in 2019. The inter – annual variance was 16.5 mm with an SD of 4.1 mm. In PdA, the
609 precipitation amounted 16.2 mm in 2017, 8 mm in 2018 and 15 mm in 2019. The inter – annual variance was
610 0.2 mm with an SD of 0.4 mm. Compared with previous years, our results might on average underestimate
611 rainfall-induced erosion rates. However, the relative erosion rate compared between burrow area and burrow
612 embedding area would stay the same.



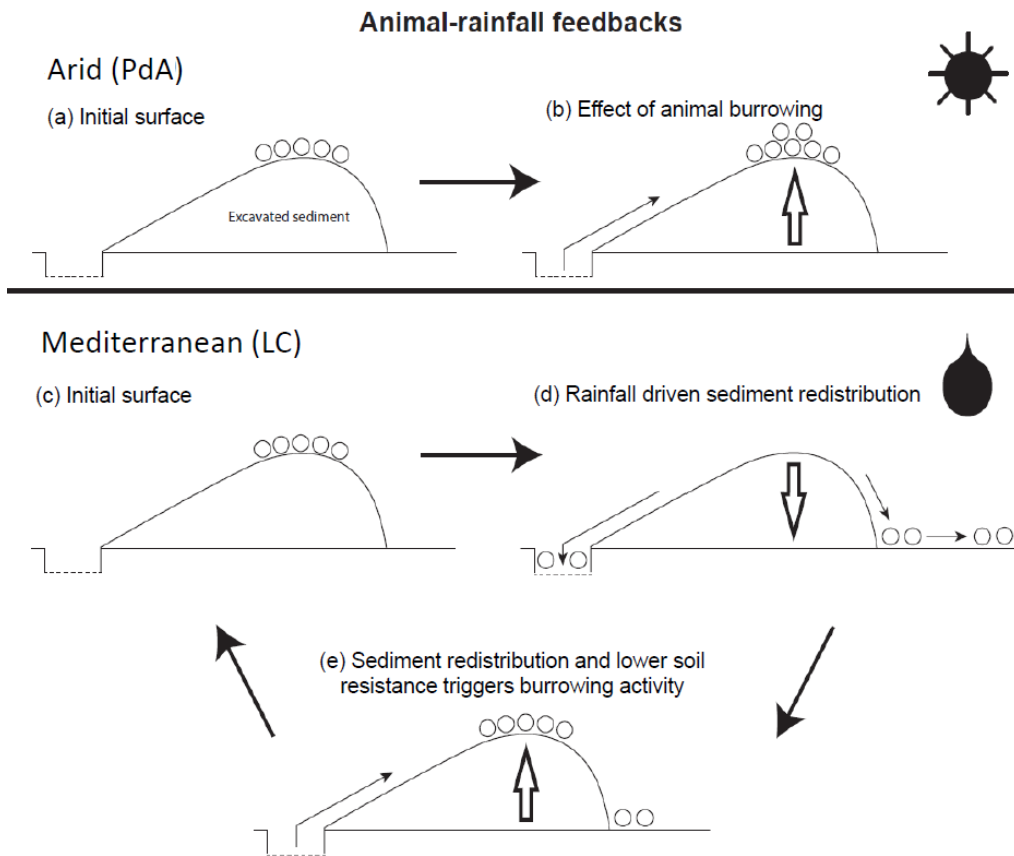
613
 614 **Figure 9.** Climate variability. (a) Daily precipitation in PdA. (b) Monthly precipitation in PdA. (c) Daily
 615 precipitation in LC. (d) Monthly precipitation in LC.

616
 617 **5.3. Sediment Redistribution**

618 Our research reveals that the presence of vertebrate burrows generally increases hillslope sediment
 619 redistribution. We show, however, that the ratio between the sediment redistribution caused by rainfall within
 620 burrow and burrow embedding areas varies between climate zones. Sediment redistribution within burrow
 621 areas was 40% higher at the arid research site, and at the mediterranean research site, it was 338% higher
 622 when compared to burrow embedding area (Table A6).

623 By monitoring microtopographical changes in a high spatio-temporal resolution, we found that the
 624 occurrence of larger rainfall events played a two-fold, accelerating role in influencing sediment redistribution
 625 (Fig. 10). Firstly, rainfall-runoff eroded burrow material caused increased sediment loss. This was followed by
 626 animal burrowing activity after the rainfall. This means that rainfall triggered animal burrowing activity which
 627 was very likely related to a lower burrowing resistance of the soil due to the increased soil moisture (Rutin
 628 1996; Romañach et al. 2005; Herbst und Bennett 2006). This double feedback led to frequently occurring but
 629 small redistribution rates. However, cumulatively, the mechanism increased downhill sediment fluxes. Previous
 630 studies most likely missed this low magnitude but frequent surface processes due to a lower monitoring
 631 duration and frequency, or artificial laboratory conditions, and thus, did not quantify the full volume of
 632 redistributed sediment associated with burrowing activity. To quantify all occurred sediment redistribution
 633 processes, a continuous surface monitoring, like the here presented, is needed.

634



635

636 **Figure 10.** Scheme of animal-driven and rainfall-driven sediment redistribution processes in both investigated
 637 climate zones: (a) Describes the initial surface of the burrow before the start of a sediment redistribution
 638 process, and (b) the animal excavation process in the arid climate zone. Here, due to rarely occurring rainfall
 639 events, sediment redistribution is mostly controlled by the animal burrowing activity; (c) describes the initial
 640 burrow surface in the mediterranean climate zone, (d) the process of sediment redistribution during a rainfall
 641 event and (e) the subsequent animal burrowing activity. Burrowing is triggered by decreased soil resistance
 642 due to the increased soil moisture after rainfall as well as by sediment accumulation within the burrow's
 643 entrance. Burrowing activity leads to a new supply of sediment being excavated to the surface. In the
 644 mediterranean climate zone, sediment redistribution is controlled by both animal burrowing activity and rainfall.
 645 The alternating excavation and erosion process ultimately lead to an increase in redistribution rates.

646

647 Our results indicate an up to 338% increase in the sediment volume redistributed during rainfall events
 648 measured within burrows when compared to burrow embedding areas. In contrast to our result, the maximum
 649 increase estimated in previous studies was 208% (Table A6, (Imeson und Kwaad 1976). The two climate
 650 zones also show different patterns: In the mediterranean climate, the contribution of animals' (vertebrates')
 651 burrowing activity appear larger than previously observed by using field methods such as erosion pins or
 652 splash traps (from -3% until - 208%, Table A6, (Imeson und Kwaad 1976; Hazelhoff et al. 1981; Black und
 653 Montgomery 1991). In contrast, in arid PdA, our study found a much smaller increase (40%, Table A6) in the
 654 sediment volume redistributed during rainfall events measured within burrows when compared to burrow
 655 embedding areas. This is lower than previously estimated (125%, Table A6, (Black und Montgomery 1991).
 656 However, solely one rainfall event above 0.2 mm day^{-1} occurred during our monitoring period. Hence, we
 657 conclude that the contribution of burrowing activity of animals to hillslope sediment transport is much larger in

658 areas with frequent rainfall events than previously thought, while it has been realistically estimated by previous
659 studies for areas with rare rainfall events (Table A6).

660 Magnitudes of sediment volume redistributed within burrows similar to our results were previously
661 obtained solely in studies applying rainfall simulators. These studies estimated an increase in the volume of
662 sediment redistributed during rainfall events, measured within burrows when compared to burrow embedding
663 areas, to be between 205% and 473% (Table A6, (Li et al. 2018; Chen et al. 2021). However, a rainfall simulator
664 can only provide data on surface processes within a plot of a few m² in size and under ideal laboratory
665 conditions while ignoring the uphill microtopography, vegetation cover and distribution (Iserloh et al. 2013),
666 which were shown to reduce erosion rates. More importantly, the rainfall intensity on hillslopes decreases with
667 (i) the angle of incidence of the rain, (ii) the inclination of the surface and (iii) the relative orientation of the
668 sloping surface to the rain vector (Sharon 1980). When simulating a rainfall event with the same rainfall volume
669 as in the field, the rain is induced directly over the treated surface and has thus a higher velocity which leads
670 to an increased splash erosion than under natural conditions (Iserloh et al. 2013). We thus propose that the
671 rainfall experiments overestimate the erosion rate while the correct erosion rate can be measured solely under
672 field conditions.

673 Cumulative sediment redistribution within burrow roof, mound and entrance was, on average, 28%
674 lower than cumulative sediment redistribution only within the mound and the burrow roof (Figure A7). These
675 results suggest that 28% of the eroded sediment from animal mounds and burrow roofs is re-accumulated
676 within the burrow entrance during rainfall-runoff events, and the remaining 62% is incorporated into overall
677 hillslope sediment flux. Our numbers contrast with previous studies, which quantified that about 58% of the
678 sediment excavated by animals will accumulate back in the burrow entrance and only 42% is incorporated to
679 downhill sediment flux (Andersen 1987; Reichman und Seabloom 2002). Hence, our results indicate not only
680 higher redistribution rates within burrows by burrowing animals but also point to much higher supply of
681 sediment to the downhill sediment flux than previously thought.

682 Our cost-effective ToF device provides data on surface changes in a high spatio-temporal resolution.
683 The high temporal resolution was able to unravel ongoing low magnitude but frequent animal excavation and
684 erosion processes. The high spatial resolution enabled us to estimate the exact volume of sediment fluxes
685 from the burrows downhill. The here presented results indicate that the contribution of burrowing animals on
686 the burrow as well as on the hillslope scale was much higher than previously assumed. Our results can be
687 integrated into long-term soil erosion models that rely on soil processes and improve their accuracy by
688 including animal-induced surface processes on microtopographical scales in their algorithms.

689

690

691 **Funding:** This study was funded by the German Research Foundation, DFG [grant numbers
692 BE1780/52-1, LA3521/1-1, FA 925/12-1, BR 1293-18-1], and is part of the DFG Priority Programme
693 SPP 1803: EarthShape: Earth Surface Shaping by Biota, sub-project “Effects of bioturbation on rates
694 of vertical and horizontal sediment and nutrient fluxes”.

695 **Institutional Review Board Statement:** Not applicable.

696 **Informed Consent Statement:** Not applicable.

697 **Acknowledgments:** We thank CONAF for the kind support provided during our field campaign.

698 **Competing interests:** There is no conflict of interest.

699 **Author contribution:** JB, AL and SA planned the campaign; PG and SA performed the measurements; PG
700 analysed the data and wrote the manuscript draft; AL, JB, NF, RB, KÜ, LP, CR, DK and PP reviewed and edited
701 the manuscript.

702 **Code/Data availability:** Code and all raw data can be provided by the corresponding author upon request.

703

704 **Appendices**

705 **Table A1.** List of abbreviations

α [°]	Tilt angle of the camera
b [°]	Surface inclination
Ω	Threshold value for the scan scattering error
B	Burrow
Area _{burrow}	mean in the field measured size of the burrows which are monitored
Area	total surface area monitored by the camera
BD	Bulk density
c [m/s]	Speed of light
D	Distance from the camera to the object
Dens _{burrow}	Burrow density
DSM	Digital surface model
DSM _{after}	DSM calculated from the scan taken after the extraction
DSM _{before}	DSM calculated from the scan taken before the extraction
EM	Burrow embedding area
Entrance	entrance to the animal burrow
g [-]	ratio [-] of the reflected photons to all photons
LC	National Park LC
LC-NL	Camera in LC on the lower north-facing hillslope
LC-NU	Camera in LC on the upper north-facing hillslope
LC-SL	Camera in LC on the lower south-facing hillslope
LC-SU	Camera in LC on the upper south-facing hillslope
MAE	Mean absolute error
MAP [°]	Mean annual precipitation
m.a.s.l.	Meters above sea level
MAT	Mean annual temperature
mClay [%]	Mean content of clay
mean _{z-coordinate}	Mean value of the z-coordinates
Mound	the sediment excavated by the animal while digging the burrow
mSand [%]	Mean content of sand
mSilt [%]	Mean content of silt
n	Number of scans
PdA	National Park Pan de Azúcar
PdA-NL	Camera in PdA on the lower north-facing hillslope

PdA-NU	Camera in PdA on the upper north-facing hillslope
PdA-SL	Camera in PdA on the lower south-facing hillslope
PdA-SU	Camera in PdA on the upper south-facing hillslope
Res	Resolution
Roof	sediment pushed aside and uphill the entrance during burrow creation
S_a	scan after the rainfall event
S_b	scan before the rainfall event
SBC	Single board computer
sd_{z-coordinate}	standard deviation of the z-coordinates
SSH	Secure shell
t [s]	Overall time of camera illumination
TOC [%]	Total organic carbon
ToF	Time-of-Flight
Vol_{burrow}	volume of redistributed sediment within burrow
Vol_{detected}	volume of the extracted sediment as detected by the camera
Vol_{add}	difference in redistributed sediment volume between burrows and burrow embedding areas
Vol_{exc}	Volume of the sediment excavated by the animal
Vol_{hillslope-wide}	Hillslope-wide volume of redistributed sediment
Vol_{measured}	volume of the extracted sediment measured by the measuring cup
Vol_{per burrow}	Volume of redistributed sediment per burrow
Vol_{per pixel}	Volume of redistributed sediment per pixel
Vol_{redistributed}	volume of the calculated redistributed sediment
Vol_{embedding}	volume of redistributed sediment within burrow embedding area
y_i	distance of the point to the point of origin at the camera nadir
Z_{cor}	Corrected z-coordinate
Z_{uncor}	Uncorrected z-coordinate

706

707 **Table A2.** Number of usable scans for each camera

Camera	Latitude	Longitude	Number of scans	Percentage of usable scans taken at 1am / 5am / 8am / 10pm	Time period
PdA-NU	-25.98131	-70.6166	238	29 / 27 / 20 / 24	18.3.-18.9.
PdA-NL	-25.98277	-70.61278	52	24 / 0 / 40 / 36	27.3.-31.5
PdA-SU	-25.97477	-70.61641	351	30 / 26 / 32 / 11	16.3.-19.9.
PdA-SL	-25.97177	-70.61409	167	48 / 38 / 7 / 8	16.3.-19.9.
LC-NU	-32.95230	-71.06231	215	37 / 20 / 8 / 33	9.3.-9.9.
LC-NL	-32.93928	-71.08613	3	-	6.3.-12.9
LC-SU	-32.93078	-71.09066	160	22 / 28 / 26 / 25	28.3.-22.5
LC-SL	-32.93110	-71.08987	167	27 / 25 / 22 / 26	16.3.-19.9.

708

709 **Table A3.** Summary of the volume of redistributed sediment, according to area and disturbance type. Vol_{exc}
710 describes volume of the sediment excavated by the animals. Vol_{burrow} describes volume of the sediment
711 redistributed during rainfall events within burrows. Vol_{add} describes the difference in redistributed sediment
712 volume within burrows and burrow embedding area during rainfall.

Disturbance	Area	PdA	LC
Vol_{exc}	Burrow	16.41 cm ³ cm ⁻² year ⁻¹	14.62 cm ³ cm ⁻² year ⁻¹
	Per burrow	1498.66 cm ³ burrow ⁻¹ year ⁻¹	1226.61 cm ³ burrow ⁻¹ year ⁻¹
	Hillslope-wide	0.18 m ³ ha ⁻¹ year ⁻¹	0.67 m ³ ha ⁻¹ year ⁻¹
Vol_{burrow}	Burrow	-1.97 cm ³ cm ⁻² year ⁻¹	-10.44 cm ³ cm ⁻² year ⁻¹
	Per burrow	-126.36 cm ³ burrow ⁻¹ year ⁻¹	-876.38 cm ³ burrow ⁻¹ year ⁻¹
	Hillslope-wide	-0.05 m ³ ha ⁻¹ year ⁻¹	-0.48 m ³ ha ⁻¹ year ⁻¹
Vol_{add}	Burrow	-1.18 cm ³ cm ⁻² year ⁻¹	-7.37 cm ³ cm ⁻² year ⁻¹
	Per burrow	-48.36 cm ³ burrow ⁻¹ year ⁻¹	-619.2 cm ³ burrow ⁻¹ year ⁻¹
	Hillslope-wide	-0.02 m ³ ha ⁻¹ year ⁻¹	-0.34 m ³ ha ⁻¹ year ⁻¹

713

714

715 **Table A4.** Summary of the volume of redistributed sediment, according to area and disturbance type. Vol_{exc}
716 describes volume of the sediment excavated by the animals. Vol_{burrow} describes volume of the sediment
717 redistributed during rainfall events within burrows. Vol_{add} describes the difference in redistributed sediment
718 volume within burrows and burrow embedding areas during rainfall.

Disturbance	Area	PdA	LC
Vol_{exc}	Burrow	9.57 cm ³ cm ⁻² 7 months ⁻¹	8.53 cm ³ cm ⁻² 7 months ⁻¹
	Per burrow	874.22 cm ³ burrow ⁻¹ 7 months ⁻¹	715.52 cm ³ burrow ⁻¹ 7 months ⁻¹
	Hillslope-wide	0.11 m ³ ha ⁻¹ 7 months ⁻¹	0.39 m ³ ha ⁻¹ 7 months ⁻¹
Vol_{burrow}	Burrow	-1.15 cm ³ cm ⁻² 7 months ⁻¹	-6.09 cm ³ cm ⁻² 7 months ⁻¹
	Per burrow	-73.71 cm ³ burrow ⁻¹ 7 months ⁻¹	-511.22 cm ³ burrow ⁻¹ 7 months ⁻¹
	Hillslope-wide	-0.03 m ³ ha ⁻¹ 7 months ⁻¹	-0.28 m ³ ha ⁻¹ 7 months ⁻¹
Vol_{add}	Burrow	-0.69 cm ³ cm ⁻² 7 months ⁻¹	-4.30 cm ³ cm ⁻² 7 months ⁻¹
	Per burrow	-28.21 cm ³ burrow ⁻¹ 7 months ⁻¹	-361.20 cm ³ burrow ⁻¹ 7 months ⁻¹
	Hillslope-wide	-0.01 m ³ ha ⁻¹ 7 months ⁻¹	-0.2 m ³ ha ⁻¹ 7 months ⁻¹

719

720 **Table A5.** Review of studies which used laser scanners for the estimation of surface processes.

Reference	R²	Error	Horizontal point spacing	Points per Model cm²	Price
------------------	----------------------	--------------	-------------------------------------	--	--------------

Our results	0.77	0.15 cm	0.32 cm	8.5	Texas Instruments OPT3101	900 USD
(Eitel et al. 2011)	0.23- 0.86	0.07 cm	NA	25	Leica ScanStation 2	102 375 USD
(Eltner et al. 2013)	NA	0.4 cm	NA	6.4	Riegl LMS- Z420i	16 795 USD
(Kaiser et al. 2014)	NA	NA	0.57 cm	NA	Riegl LMS- Z420i	16 795 USD
(Longoni et al. 2016)	NA	NA	NA	1	Riegl LMS- Z420i	16 795 USD
(Morris et al. 2011)	NA	0.5 cm	NA	NA	Maptek I-Site 4400LR	240 000 USD
(Nasermoaddeli und Pasche 2008)	NA	0.2 cm	0.25 cm	NA	Leica Cyrax HDS 2500	4500 USD
(Thomsen et al. 2015)	NA	NA	0.4 cm	NA	Leica ScanStation 2	102 375 USD

721

722

723

Table A6. Review of studies which estimated the sediment redistribution within burrows and burrow embedding areas and the proposed impact.

Reference	Climate	Animals	Method	Monitoring period	Frequency	Burrows	Burrow embedding area	Impact
(Imeson und Kwaad 1976)	continental	rodents	erosion pins	15 months	monthly	20 mm		NA
(Imeson und Kwaad 1976)	continental	rodents	splash boards	15 months	monthly	91.75g 24.49 cm ⁻² = 3.75 cm ³ cm ⁻²	94g	-3%
(Imeson und Kwaad 1976)	continental	rodents	rainfall simulation (7.5 cm / hour intensity)	One-time measurement	NA	0.2 g – 0.73 g	0.009 g – 0.23 g	+208 %
(Imeson 1977)	continental	vertebrates	rainfall simulation	One-time measurement	NA	0.18-0.3 100 J ⁻¹ m ⁻² rain	0.146 100 J ⁻¹ m ⁻² rain	+123 %

(Hazelhoff et al. 1981)	continen tal	earthwor ms	splash traps	12 months	monthly	NA	NA	+180 %
(Black und Montgomery 1991)	arid	pocket gopher	erosion pins	10 months	2 months	NA	NA	+125 %
(Hakonson 1999)	tempera te	pocket gophers	rainfall simulato r (60 mm / hour)	2 years	2 – 3 weeks	2.4 – 8.7 mg ha ⁻¹	4.4 – 15 mg ha ⁻¹	-43%
(Li et al. 2018)	tempera te	mole crickets	rainfall simulati on (36 mm / hour)	One time measureme nt	15 measure ments	22.1 g 115 cm ⁻² = 5.2 cm ³ cm ⁻²	5 g 123 cm ⁻² = 1.09 cm ³ cm ⁻²	+473 %
(Li et al. 2018)	tempera te	mole crickets	rainfall simulati on (36 mm / hour)	One time measureme nt	15 measure ments	35.3 g 220.5 cm ⁻² = 6.24 cm ³ cm ⁻²	5 g 123 cm ⁻² = 1.09 cm ³ cm ⁻²	+473 %
(Chen et al. 2021)	lab	chinese zocor	rainfall simulati on (80 mm / hour)	One-time measureme nt	3 measure ments	2,69 g cm ⁻² = 2.69 cm ³ cm ⁻²	0,88 g cm ⁻² = 0.88 cm ³ cm ⁻²	+205 %

724

725

726 **Table A7.** Review of studies which estimated the sediment redistribution within burrows, average burrow
727 density as found in the literature and area-wide yearly contribution of burrowing animals to sediment
728 redistribution.

Climate	Animals	Burrows	Average burrow density	Average burrow size	Area-wide redistribution
Continental	rodents	91.75g 24.49 cm ⁻² = 3.75 cm ³ cm ⁻² (Imeson und Kwaad 1976)	14 625 m ⁻² = 0.02 m ⁻² (Pang und Guo 2017)	24.49 cm ² (Imeson und Kwaad 1976)	0.183 m ³ ha ⁻¹ year ⁻¹
Temperate	mole crickets	22.1 g 115 cm ⁻² = 5.2 cm ³ cm ⁻² (Li et al. 2018)	405 ha ⁻¹ (Castner und Fowler 1984)	115 cm ² (Li et al. 2018)	0.24 m ³ ha ⁻¹ year ⁻¹

Temperate	mole crickets	35.3 g cm ⁻² = 220.5 cm ³ cm ⁻² (Li et al. 2018)	405	ha ⁻¹	220.5 cm ² (Li et al. 2018)	0.56 m ³ year ⁻¹	ha ⁻¹
Lab	chinese zocor	2,69 g cm ⁻² = 2.69 cm ³ cm ⁻² (Chen et al. 2021)	94.69	2500m ⁻²	1256 cm ²	1.35 m ³ year ⁻¹	ha ⁻¹

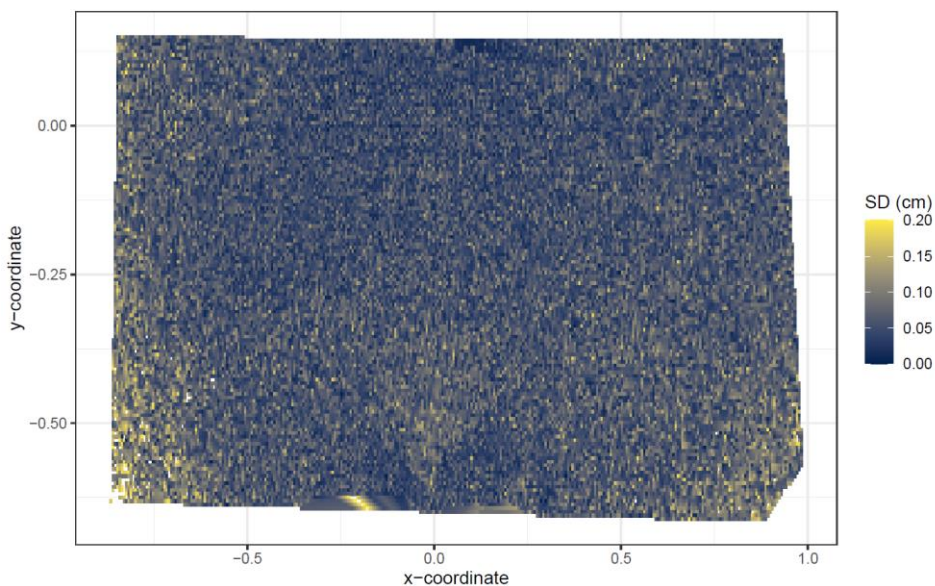
729

730 **Table A8.** Review of studies which estimated the volume of sediment excavated by burrowing animals.

	Climate	Animals	Method	Monitoring period	Frequency	volume of the excavated sediment
(Black und Montgomery 1991)	arid	porcupines	mound volume	3 years	yearly	0.2 m ³ ha ⁻¹ year ⁻¹
(Black und Montgomery 1991)	arid	isopods	mound volume	3 years	yearly	0.11 m ³ ha ⁻¹ year ⁻¹
(Black und Montgomery 1991)	arid	pocket gopher	mound volume	2 years	3 model runs	0.05 – 0.11 m ³ ha ⁻¹ year ⁻¹
(Rutin 1996)	subtropical	scorpions	mound volume	6 months	2-29 days	0.42 m ³ ha ⁻¹ year ⁻¹
(Hall et al. 1999)	alpine	rodents	mound volume	1 year	yearly	0.02 m ³ ha ⁻¹ year ⁻¹
(Hall et al. 1999)	alpine	bears	mound volume	1 year	yearly	0.49 m ³ ha ⁻¹ year ⁻¹
(Yoo et al. 2005)	arid	pocket gopher	mound volume	1 year	One model run	0.1-0.2 m ³ ha ⁻¹ year ⁻¹

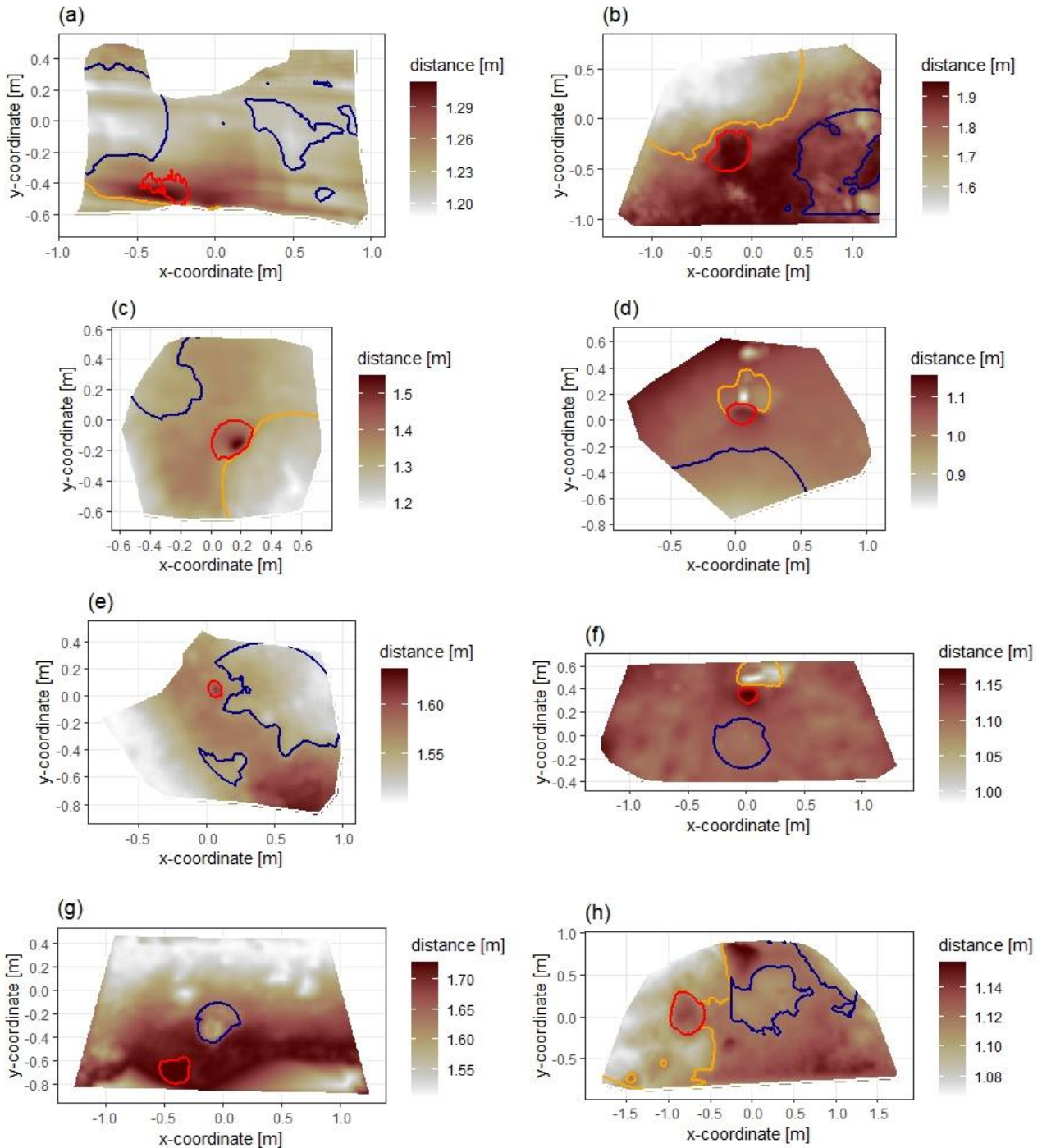
731

732

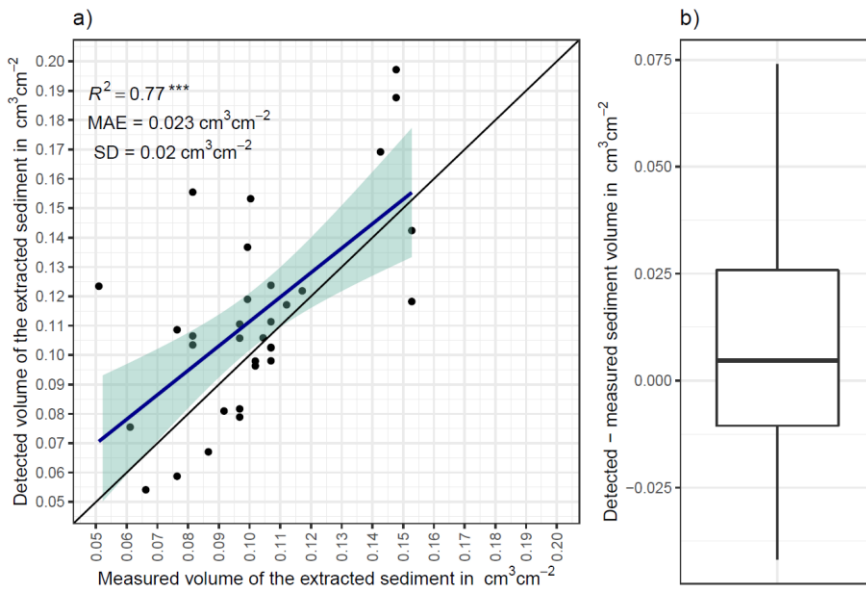


733

734 **Figure A1.** Standard deviation of the z-coordinate of unprocessed five scans showed exemplary for the camera
 735 on the upper north-facing hillside. SD is standard deviation. The error increases with distance from the camera
 736 nadir point. The standard deviation was here calculated from scans before any corrections.
 737
 738



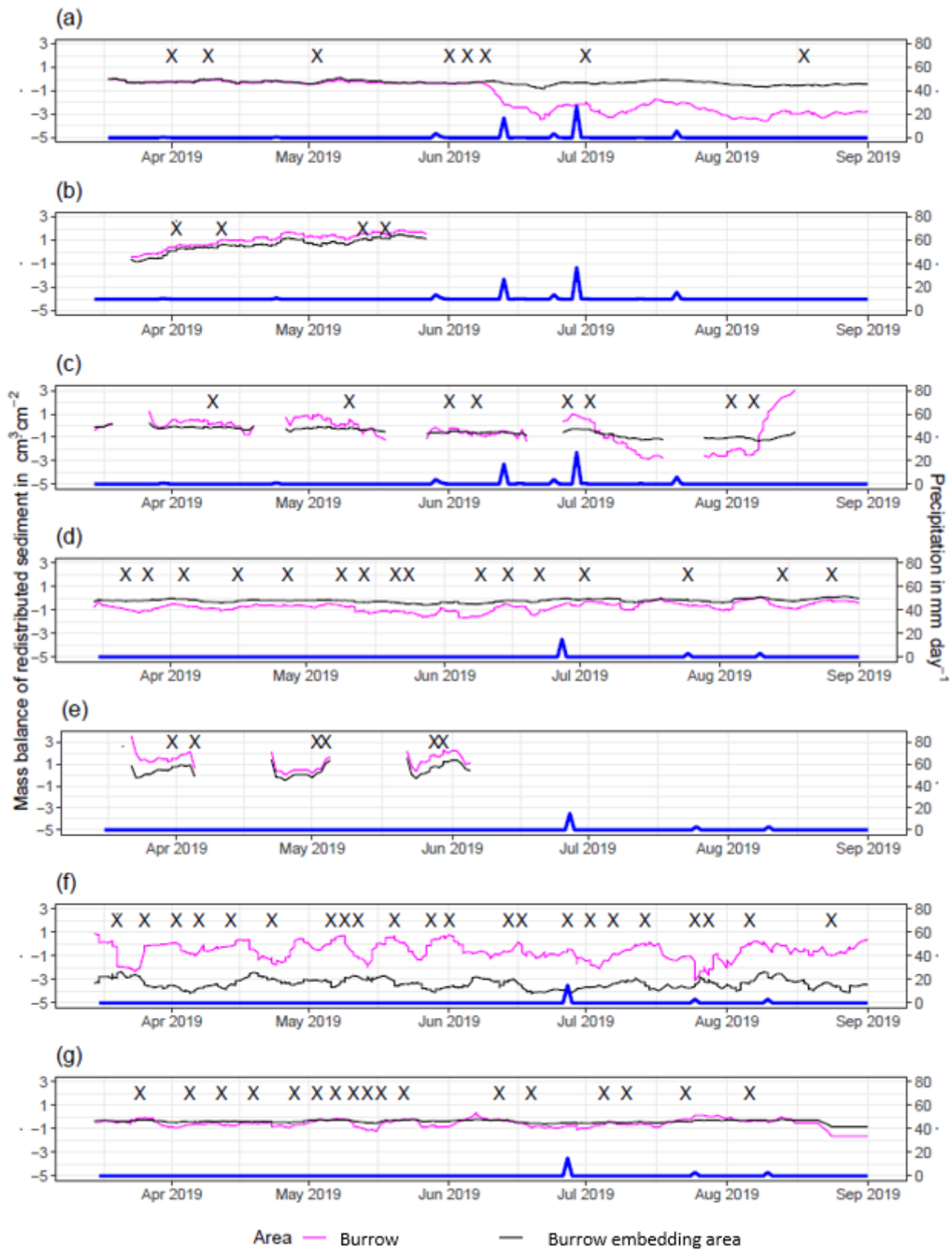
739
 740 **Figure A2.** Delineation of the areas. The point of origin of the coordinate system is at the camera nadir. Depth
 741 is the distance between the surface and the camera. Red is the outline of the burrow entrance. Green is the
 742 outline of mound. Orange is the outline of burrow roof. Area which is not outlined is burrow embedding area.
 743 Arrow indicates downhill direction of the hillslope. (a) LC-NU. (b) LC-NL (c) LC-SU. (d) LC-SL. (e) PdA-NU. (f)
 744 PdA-NL. (g) PdA-SU. (h) PdA-SL.



745

746 **Figure A3.** a) Estimation of Time-of-Flight camera accuracy based on averaging two surface scans before and
 747 after the sediment extraction under controlled conditions. The x-axis shows the exact sediment volume
 748 measured with a cup. The y-axis represents the volume of the sediment calculated from the camera scans
 749 (according to Equation (4)). The blue line is the linear regression calculated from the measured and detected
 750 volume. The green shadow shows the confidence interval of 95% for the linear regression slope. $^{***}p \leq 0.001$.
 751 MAE is the mean absolute error, SD is standard deviation and R^2 the coefficient of determination. b) Measured
 752 sediment volume subtracted from the detected sediment volume for all measurements.

753

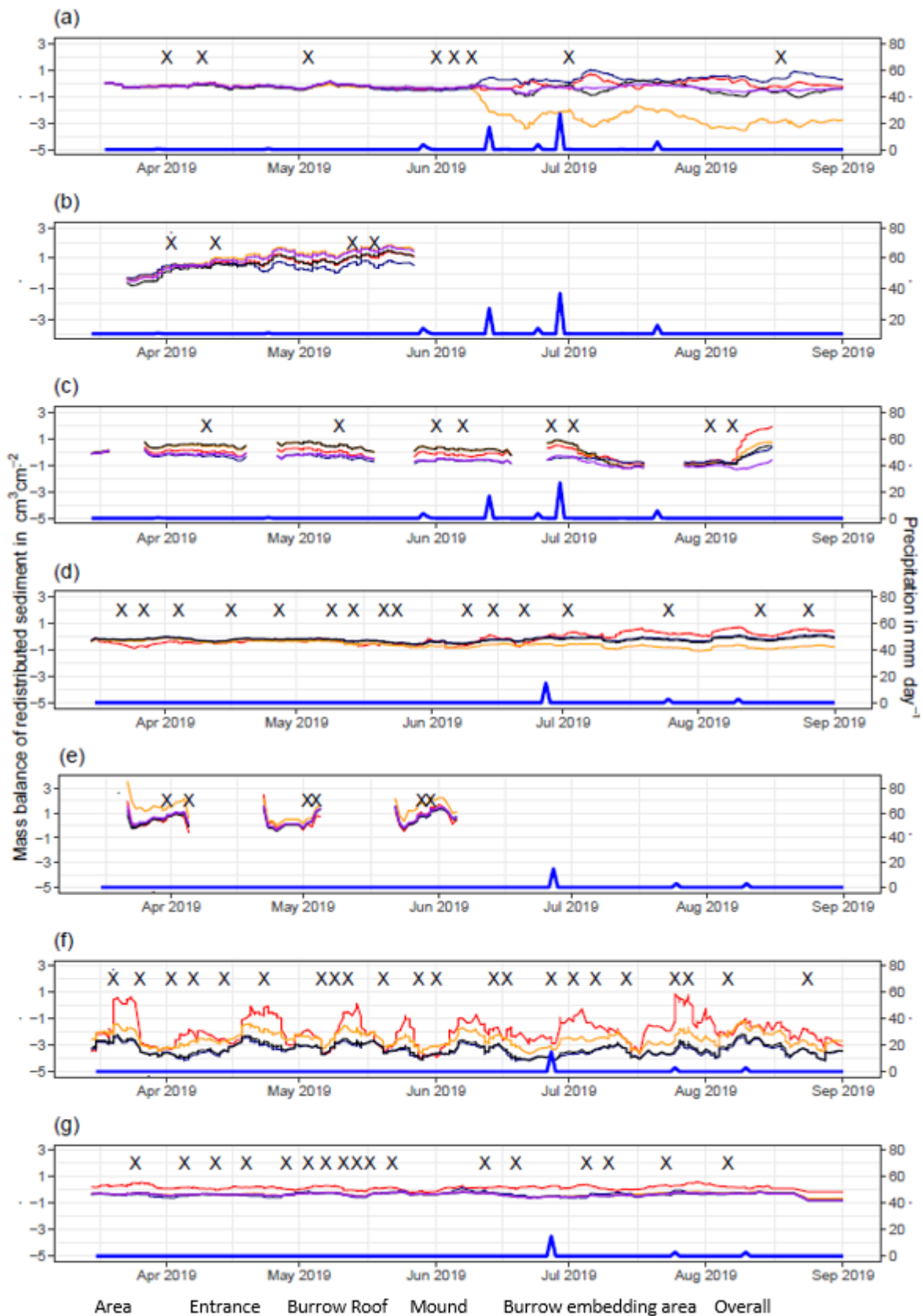


754

755 **Figure A4.** Sediment mass balance for the period of 7 months separately for burrows and burrow embedding

756 areas as measured by the cameras. (a) LC-NU. (b) LC-SU. (c) LC-SL. (d) PdA-NU. (e) PdA-NL.

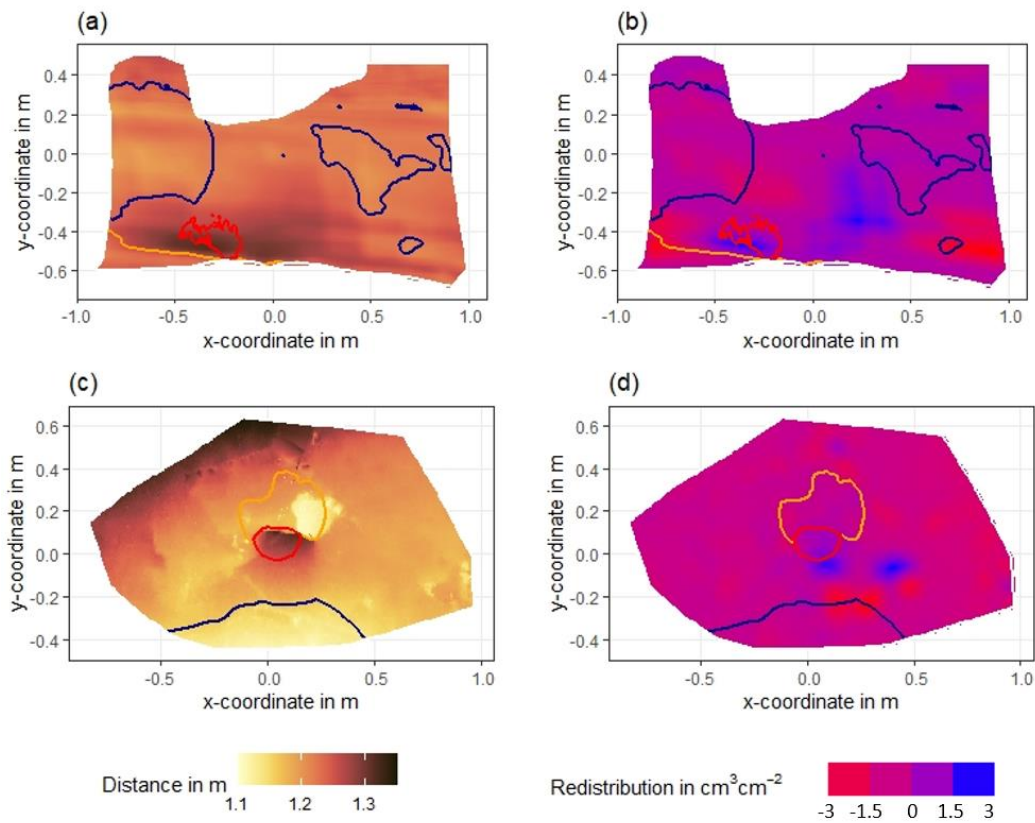
757 (f) PdA-SU. (g) PdA-SL. For abbreviations see Table A1.



758

759 **Figure A5.** Sediment mass balance for the period of 7 months separately for all delineated areas as measured
 760 by the cameras. (a) LC-NU. (b) LC-SU. (c) LC-SL. (d) PdA-NU. (e) PdA-NL. (f) PdA-SU. (g) PdA-SL. For
 761 abbreviations see Table A1.

762

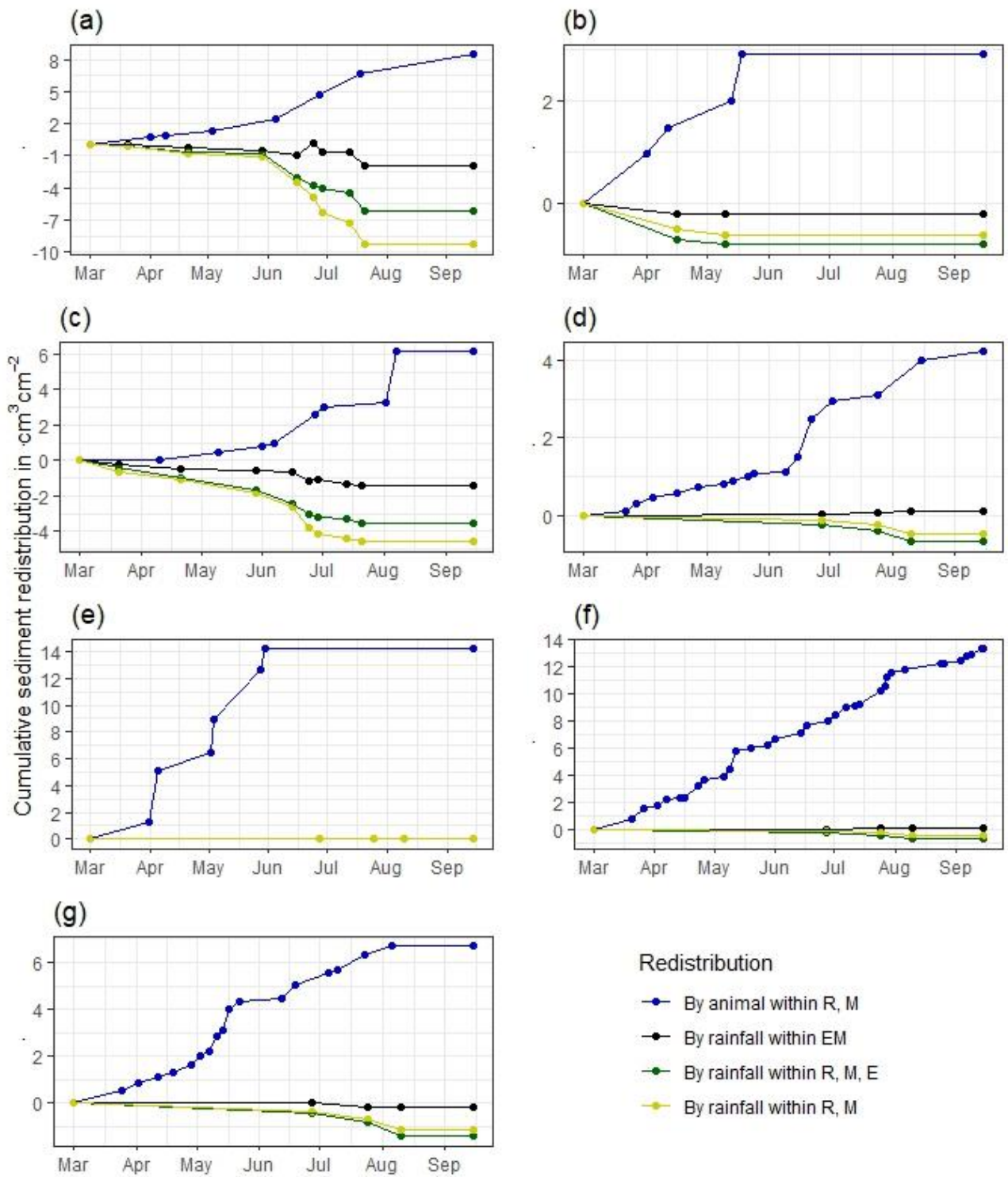


763

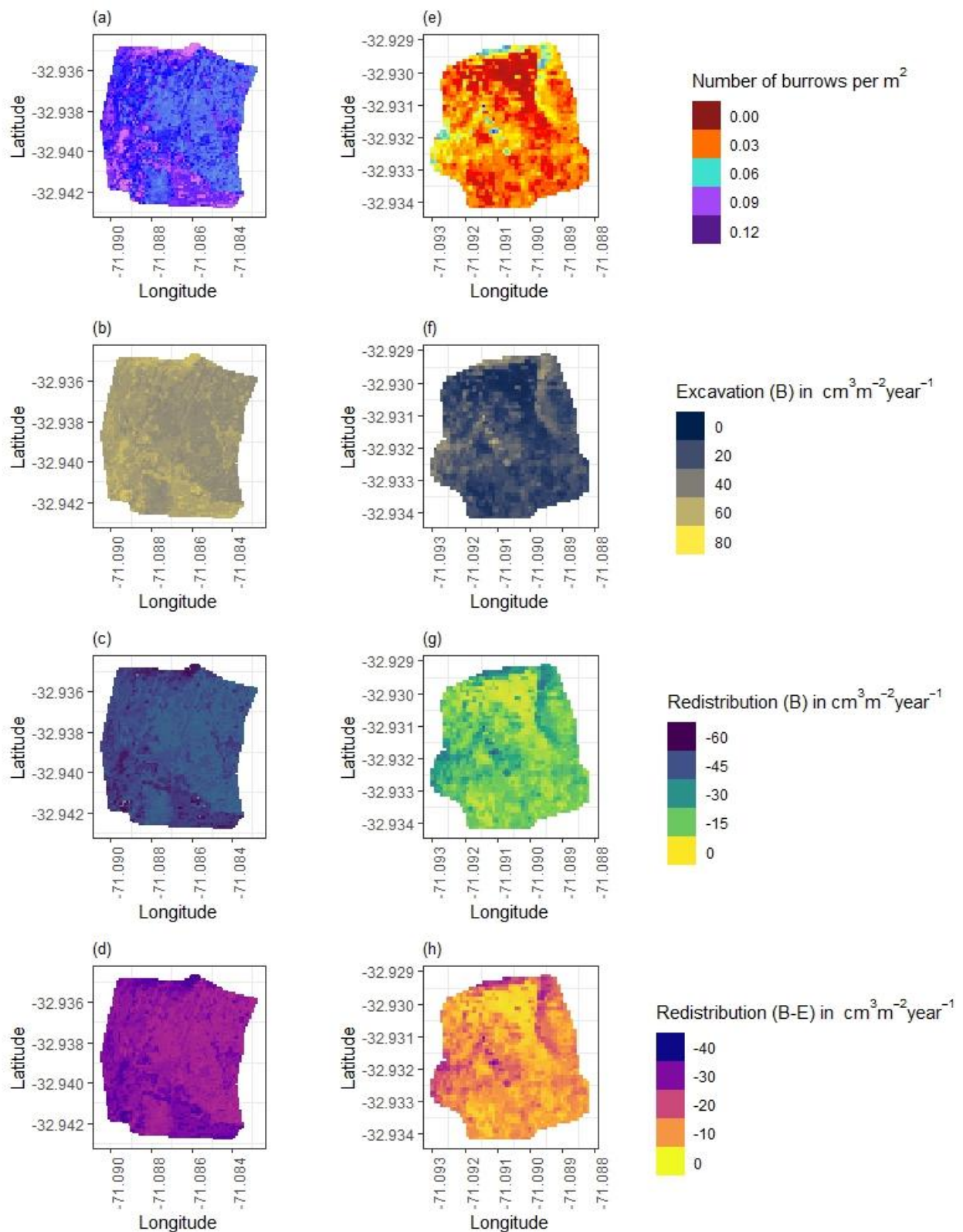
764 **Figure A6.** Examples of surface scans showing the digital surface model (DSM) before a rainfall event (a, c)
 765 at two camera locations in La Campana, and the calculated volume of redistributed sediment (b, d) after the
 766 rainfall event: (a) DSM of a scan from the camera on the upper north-facing hillslope in La Campana; (b)
 767 Detected sediment redistribution ($\text{cm}^3 \text{cm}^{-2}$) on the upper north-facing hillslope in La Campana after a rainfall
 768 event of 17.2 mm day^{-1} ; (c) DSM of a scan from the camera on the upper south-facing hillslope in La Campana;
 769 (d) Detected sediment redistribution ($\text{cm}^3 \text{cm}^{-2}$) on the upper south-facing hillslope after a rainfall event of
 770 17.2 mm day^{-1} . Red is the outline of the burrow entrance. Green is the outline of mound. Orange is the outline of
 771 the burrow roof. The area which is not outlined is burrow embedding area. Redistribution is the volume of the
 772 redistributed sediment, either accumulated (positive value) or eroded (negative value) per $\text{cm}^3 \text{cm}^{-2}$. After the
 773 rainfall events, sediment mostly accumulated within the burrow entrance or near mounds and eroded from
 774 burrow roofs and mounds.

775

776

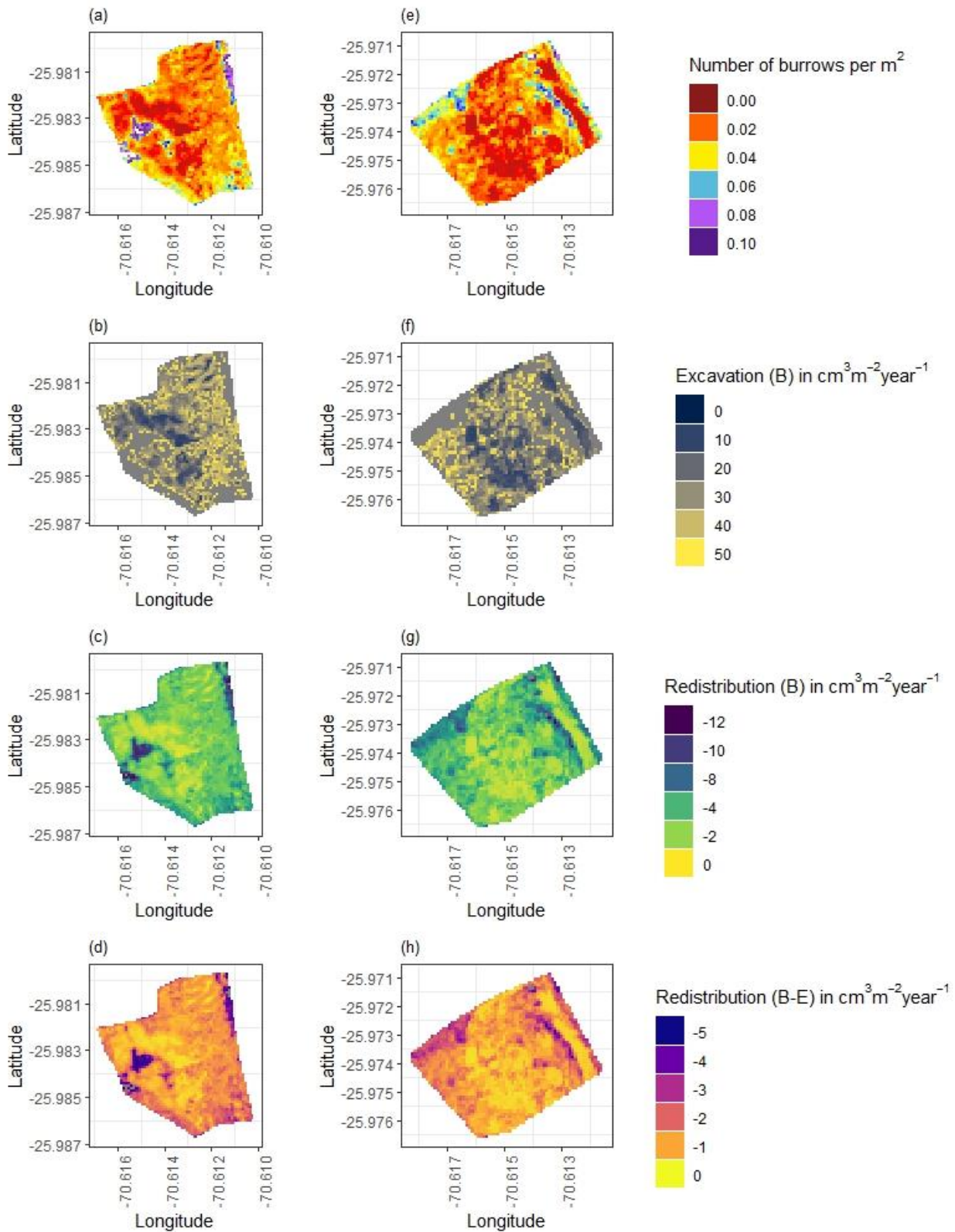


777
 778 **Figure A7.** Cumulative volume of redistributed sediment for all cameras. Positive values indicate sediment
 779 accumulation. Negative values indicate sediment erosion. Whiskers are the median sediment redistribution. E
 780 is the burrow entrance. M is the mound. R is burrow roof. EM is burrow embedding area. LC is mediterranean
 781 climate zone. PdA is arid climate zone. (a) LC-NU. (b) LC-SU. (c) LC-SL. (d) PdA-NU. (e) PdA-NL. (f) PdA-
 782 SU. (g) PdA-SL. For abbreviations see Table A1.
 783



784
 785 **Figure A8.** Hillslope-wide volume of redistributed sediment for a time period of one year in LC. (a-d) North-
 786 facing hillslope. (e-h) South-facing hillslope. (a) and (e) Density of burrows as estimated by Grigusova et al.
 787 2021. (b) and (f) Volume of the sediment excavated by the animals. (c) and (g) Volume of the sediment
 788 redistributed during rainfall events within burrows. (d) and (h) Volume of additionally redistributed sediment
 789 during rainfall events due to presence of the burrows. The values were calculated per burrow as stated in
 790 section 3.7 by subtracting the sediment volume redistributed within burrows from the sediment volume

791 redistributed within burrow embedding area and then upscaled. B stays for burrow, EM stays for burrow
 792 embedding area.
 793



794
 795 **Figure A9.** Hillslope-wide volume of redistributed sediment for a time period of one year in Pan de Azúcar. (a-
 796 d) North-facing hillslope. (e-h) South-facing hillslope. (a) and (e) Density of burrows as estimated by Grigusova
 797 et al. 2021. (b) and (f) Volume of the sediment excavated by the animals. (c) and (g) Volume of the sediment
 798 redistributed during rainfall events within burrows. (d) and (h) Volume of additionally redistributed sediment

799 during rainfall events due to presence of the burrows. The values were calculated per burrow as stated in
800 section 3.7 by subtracting the sediment volume redistributed within burrow from the sediment volume
801 redistributed within burrow embedding area and then upscaled. B stays for burrow, EM stays for burrow
802 embedding area by the burrowing animal.

803

804 Literaturverzeichnis

- 805 Afana, A.; Solé-Benet, A.; Pérez, J. L. (2010): Determination of Soil Erosion Using Laser Scanners.
806 Experimental Station of Arid Zone, CSIC, Almeria, Spain, zuletzt geprüft am 22.12.2021.
- 807 Andersen, Douglas C. (1987): *Geomys Bursarius* Burrowing Patterns: Influence of Season and
808 Food Patch Structure. In: *Ecology* 68 (5), S. 1306–1318. DOI: 10.2307/1939215.
- 809 Ashcroft, Michael B.; Gollan, John R.; Ramp, Daniel (2014): Creating vegetation density profiles
810 for a diverse range of ecological habitats using terrestrial laser scanning. In: *Methods Ecol Evol*
811 5 (3), S. 263–272. DOI: 10.1111/2041-210X.12157.
- 812 BANCROFT, W. J.; HILL, D.; ROBERTS, J. D. (2004): A new method for calculating volume of
813 excavated burrows: the geomorphic impact of Wedge-Tailed Shearwater burrows on Rottne
814 Island. In: *Funct Ecology* 18 (5), S. 752–759. DOI: 10.1111/j.0269-8463.2004.00898.x.
- 815 Bernhard, Nadine; Moskwa, Lisa-Marie; Schmidt, Karsten; Oeser, Ralf A.; Aburto, Felipe; Bader,
816 Maaïke Y. et al. (2018): Pedogenic and microbial interrelations to regional climate and local
817 topography: New insights from a climate gradient (arid to humid) along the Coastal Cordillera
818 of Chile. In: *CATENA* 170 (4), S. 335–355. DOI: 10.1016/j.catena.2018.06.018.
- 819 Black, Thomas A.; Montgomery, David R. (1991): Sediment transport by burrowing mammals,
820 Marin County, California. In: *Earth Surf. Process. Landforms* 16 (2), S. 163–172. DOI:
821 10.1002/esp.3290160207.
- 822 Blanch, Xabier; Eltner, Anette; Guinau, Marta; Abellan, Antonio (2021): Multi-Epoch and Multi-
823 Imagery (MEMI) Photogrammetric Workflow for Enhanced Change Detection Using Time-
824 Lapse Cameras. In: *Remote Sensing* 13 (8), S. 1460. DOI: 10.3390/rs13081460.
- 825 Castner, James L.; Fowler, H. G. (1984): Distribution of Mole Crickets (Orthoptera: Gryllotalpidae:
826 *Scapteriscus*) and the Mole Cricket Parasitoid *Larra bicolor* (Hymenoptera: Sphecidae) in Puerto
827 Rico. In: *The Florida Entomologist* 67 (3), S. 481. DOI: 10.2307/3494730.
- 828 Chen, Mingyu; Ma, Li; Shao, Ming'an; Wei, Xiaorong; Jia, Yuhua; Sun, Shuchen et al. (2021):
829 Chinese zokor (*Myospalax fontanierii*) excavating activities lessen runoff but facilitate soil
830 erosion – A simulation experiment. In: *CATENA* 202 (8), S. 105248. DOI:
831 10.1016/j.catena.2021.105248.
- 832 Coombes, Martin A. (2016): Biogeomorphology: diverse, integrative and useful. In: *Earth Surf.*
833 *Process. Landforms* 41 (15), S. 2296–2300. DOI: 10.1002/esp.4055.
- 834 Corenblit, Dov; Corbara, Bruno; Steiger, Johannes (2021): Biogeomorphological eco-evolutionary
835 feedback between life and geomorphology: a theoretical framework using fossorial mammals.
836 In: *Die Naturwissenschaften* 108 (6), S. 55. DOI: 10.1007/s00114-021-01760-y.
- 837 Eitel, Jan U.H.; Williams, C. Jason; Vierling, Lee A.; Al-Hamdan, Osama Z.; Pierson, Frederick B.
838 (2011): Suitability of terrestrial laser scanning for studying surface roughness effects on
839 concentrated flow erosion processes in rangelands. In: *CATENA* 87 (3), S. 398–407. DOI:
840 10.1016/j.catena.2011.07.009.
- 841 Eltner, A.; Mulrow, C.; Maas, H.-G. (2013): QUANTITATIVE MEASUREMENT OF SOIL
842 EROSION FROM TLS AND UAV DATA. In: *Int. Arch. Photogramm. Remote Sens. Spatial Inf.*
843 *Sci.* XL-1/W2, S. 119–124. DOI: 10.5194/isprsarchives-XL-1-W2-119-2013.
- 844 Eltner, A.; Schneider, D.; Maas, H.-G. (2016a): INTEGRATED PROCESSING OF HIGH
845 RESOLUTION TOPOGRAPHIC DATA FOR SOIL EROSION ASSESSMENT
846 CONSIDERING DATA ACQUISITION SCHEMES AND SURFACE PROPERTIES. In: *Int.*
847 *Arch. Photogramm. Remote Sens. Spatial Inf. Sci.* XLI-B5, S. 813–819. DOI:

848 10.5194/isprsarchives-XLI-B5-813-2016.

849 Eltner, Anette; Kaiser, Andreas; Abellan, Antonio; Schindewolf, Marcus (2017): Time lapse
850 structure-from-motion photogrammetry for continuous geomorphic monitoring. In: *Earth Surf.*
851 *Process. Landforms* 42 (14), S. 2240–2253. DOI: 10.1002/esp.4178.

852 Eltner, Anette; Kaiser, Andreas; Castillo, Carlos; Rock, Gilles; Neugirg, Fabian; Abellán, Antonio
853 (2016b): Image-based surface reconstruction in geomorphometry – merits, limits and
854 developments. In: *Earth Surf. Dynam.* 4 (2), S. 359–389. DOI: 10.5194/esurf-4-359-2016.

855 Gabet, Emmanuel J.; Reichman, O. J.; Seabloom, Eric W. (2003): The Effects of Bioturbation on
856 Soil Processes and Sediment Transport. In: *Annu. Rev. Earth Planet. Sci.* 31 (1), S. 249–273.
857 DOI: 10.1146/annurev.earth.31.100901.141314.

858 Galland, Olivier; Bertelsen, Håvard S.; Guldstrand, Frank; Girod, Luc; Johannessen, Rikke F.;
859 Bjugger, Fanny et al. (2016): Application of open-source photogrammetric software MicMac for
860 monitoring surface deformation in laboratory models. In: *J. Geophys. Res. Solid Earth* 121 (4),
861 S. 2852–2872. DOI: 10.1002/2015JB012564.

862 Grigusova, Paulina; Larsen, Annegret; Achilles, Sebastian; Klug, Alexander; Fischer, Robin; Kraus,
863 Diana et al. (2021): Area-Wide Prediction of Vertebrate and Invertebrate Hole Density and
864 Depth across a Climate Gradient in Chile Based on UAV and Machine Learning. In: *Drones* 5
865 (3), S. 86. DOI: 10.3390/drones5030086.

866 Hakonson, T. E. (1999): The Effects of Pocket Gopher Burrowing on Water Balance and Erosion
867 from Landfill Covers. In: *J. environ. qual.* 28 (2), S. 659–665. DOI:
868 10.2134/jeq1999.00472425002800020033x.

869 Hall, Kevin; Boelhouwers, Jan; Driscoll, Kevin (1999): Animals as Erosion Agents in the Alpine
870 Zone: Some Data and Observations from Canada, Lesotho, and Tibet. In: *Arctic, Antarctic, and*
871 *Alpine Research* 31 (4), S. 436–446. DOI: 10.1080/15230430.1999.12003328.

872 Hancock, Greg; Lowry, John (2021): Quantifying the influence of rainfall, vegetation and animals
873 on soil erosion and hillslope connectivity in the monsoonal tropics of northern Australia. In:
874 *Earth Surf. Process. Landforms* 46 (10), S. 2110–2123. DOI: 10.1002/esp.5147.

875 Hänsel, Phoebe; Schindewolf, Marcus; Eltner, Anette; Kaiser, Andreas; Schmidt, Jürgen (2016):
876 Feasibility of High-Resolution Soil Erosion Measurements by Means of Rainfall Simulations
877 and SfM Photogrammetry. In: *Hydrology* 3 (4), S. 38. DOI: 10.3390/hydrology3040038.

878 Hazelhoff, L.; van Hoof, P.; Imeson, A. C.; Kwaad, F. J. P. M. (1981): The exposure of forest soil to
879 erosion by earthworms. In: *Earth Surf. Process. Landforms* 6 (3-4), S. 235–250. DOI:
880 10.1002/esp.3290060305.

881 Herbst, M.; Bennett, N. C. (2006): Burrow architecture and burrowing dynamics of the endangered
882 Namaqua dune mole rat (*Bathergus janetta*) (Rodentia: Bathyergidae). In: *Journal of Zoology*
883 270 (3), S. 420–428. DOI: 10.1111/j.1469-7998.2006.00151.x.

884 Horn, B.K.P. (1981): Hill shading and the reflectance map. In: *Proc. IEEE* 69 (1), S. 14–47. DOI:
885 10.1109/PROC.1981.11918.

886 Imeson, A. C. (1977): Splash erosion, animal activity and sediment supply in a small forested
887 Luxembourg catchment. In: *Earth Surf. Process. Landforms* 2 (2-3), S. 153–160. DOI:
888 10.1002/esp.3290020207.

889 Imeson, A. C.; Kwaad, F. J. P. M. (1976): Some Effects of Burrowing Animals on Slope Processes
890 in the Luxembourg Ardennes. In: *Geografiska Annaler: Series A, Physical Geography* 58 (4), S.
891 317–328. DOI: 10.1080/04353676.1976.11879941.

892 Iserloh, T.; Ries, J. B.; Arnáez, J.; Boix-Fayos, C.; Butzen, V.; Cerdà, A. et al. (2013): European
893 small portable rainfall simulators: A comparison of rainfall characteristics. In: *CATENA* 110 (2),
894 S. 100–112. DOI: 10.1016/j.catena.2013.05.013.

895 James, M. R.; Robson, S. (2014): Sequential digital elevation models of active lava flows from
896 ground-based stereo time-lapse imagery. In: *ISPRS Journal of Photogrammetry and Remote*
897 *Sensing* 97 (3), S. 160–170. DOI: 10.1016/j.isprsjprs.2014.08.011.

898 Jones, Clive G.; Gutiérrez, Jorge L.; Byers, James E.; Crooks, Jeffrey A.; Lambrinos, John G.;
899 Talley, Theresa S. (2010): A framework for understanding physical ecosystem engineering by

- 900 organisms. In: *Oikos* 119 (12), S. 1862–1869. DOI: 10.1111/j.1600-0706.2010.18782.x.
- 901 Kaiser, Andreas; Neugirg, Fabian; Rock, Gilles; Müller, Christoph; Haas, Florian; Ries, Johannes;
902 Schmidt, Jürgen (2014): Small-Scale Surface Reconstruction and Volume Calculation of Soil
903 Erosion in Complex Moroccan Gully Morphology Using Structure from Motion. In: *Remote*
904 *Sensing* 6 (8), S. 7050–7080. DOI: 10.3390/rs6087050.
- 905 Kinlaw, A.; Grasmueck, M. (2012): Evidence for and geomorphologic consequences of a reptilian
906 ecosystem engineer: The burrowing cascade initiated by the Gopher Tortoise. In:
907 *Geomorphology* 157-158 (4), S. 108–121. DOI: 10.1016/j.geomorph.2011.06.030.
- 908 Kromer, Ryan; Walton, Gabe; Gray, Brian; Lato, Matt; Group, Robert (2019): Development and
909 Optimization of an Automated Fixed-Location Time Lapse Photogrammetric Rock Slope
910 Monitoring System. In: *Remote Sensing* 11 (16), S. 1890. DOI: 10.3390/rs11161890.
- 911 Kukko, Antero; Hyyppä, Juha (2009): Small-footprint Laser Scanning Simulator for System
912 Validation, Error Assessment, and Algorithm Development. In: *photogramm eng remote sensing*
913 75 (10), S. 1177–1189. DOI: 10.14358/PERS.75.10.1177.
- 914 Larsen, A.; Nardin, W.; Lageweg, W. I.; Bätz, N. (2021): Biogeomorphology, quo vadis? On
915 processes, time, and space in biogeomorphology. In: *Earth Surf. Process. Landforms* 46 (1), S.
916 12–23. DOI: 10.1002/esp.5016.
- 917 Le Hir, P.; Monbet, Y.; Orvain, F. (2007): Sediment erodability in sediment transport modelling:
918 Can we account for biota effects? In: *Continental Shelf Research* 27 (8), S. 1116–1142. DOI:
919 10.1016/j.csr.2005.11.016.
- 920 Lehnert, Lukas W.; Thies, Boris; Trachte, Katja; Achilles, Sebastian; Osses, Pablo; Baumann, Karen
921 et al. (2018): A Case Study on Fog/Low Stratus Occurrence at Las Lomitas, Atacama Desert
922 (Chile) as a Water Source for Biological Soil Crusts. In: *Aerosol Air Qual. Res.* 18 (1), S. 254–
923 269. DOI: 10.4209/aaqr.2017.01.0021.
- 924 Li, Guorong; Li, Xilai; Li, Jinfang; Chen, Wenting; Zhu, Haili; Zhao, Jianyun; Hu, Xiasong
925 (2019a): Influences of Plateau Zokor Burrowing on Soil Erosion and Nutrient Loss in Alpine
926 Meadows in the Yellow River Source Zone of West China. In: *Water* 11 (11), S. 2258. DOI:
927 10.3390/w11112258.
- 928 Li, Larry (2014): Time-of-Flight Camera – An Introduction. Technical White Paper. Hg. v. Texas
929 Instruments. Online verfügbar unter <https://www.ti.com/lit/wp/sloa190b/sloa190b.pdf>, zuletzt
930 geprüft am 22.12.2021.
- 931 Li, T. C.; Shao, M. A.; Jia, Y. H.; Jia, X. X.; Huang, L. M.; Gan, M. (2019b): Small - scale
932 observation on the effects of burrowing activities of ants on soil hydraulic processes. In: *Eur J*
933 *Soil Sci* 70 (2), S. 236–244. DOI: 10.1111/ejss.12748.
- 934 Li, Tongchuan; Jia, Yuhua; Shao, Ming'an; Shen, Nan (2019c): *Camponotus japonicus* burrowing
935 activities exacerbate soil erosion on bare slopes. In: *Geoderma* 348 (4), S. 158–167. DOI:
936 10.1016/j.geoderma.2019.04.035.
- 937 Li, Tongchuan; Shao, Ming'an; Jia, Yuhua; Jia, Xiaoxu; Huang, Laiming (2018): Small-scale
938 observation on the effects of the burrowing activities of mole crickets on soil erosion and
939 hydrologic processes. In: *Agriculture, Ecosystems & Environment* 261 (4), S. 136–143. DOI:
940 10.1016/j.agee.2018.04.010.
- 941 Longoni, Laura; Papini, Monica; Brambilla, Davide; Barazzetti, Luigi; Roncoroni, Fabio; Scaioni,
942 Marco; Ivanov, Vladislav (2016): Monitoring Riverbank Erosion in Mountain Catchments Using
943 Terrestrial Laser Scanning. In: *Remote Sensing* 8 (3), S. 241. DOI: 10.3390/rs8030241.
- 944 Malizia, Ana I. (1998): Population dynamics of the fossorial rodent *Ctenomys talarum* (Rodentia:
945 Octodontidae). In: *Journal of Zoology* 244 (4), S. 545–551. DOI: 10.1111/j.1469-
946 7998.1998.tb00059.x.
- 947 MALLALIEU, JOSEPH; CARRIVICK, JONATHAN L.; QUINCEY, DUNCAN J.; SMITH,
948 MARK W.; JAMES, WILLIAM H.M. (2017): An integrated Structure-from-Motion and time-
949 lapse technique for quantifying ice-margin dynamics. In: *J. Glaciol.* 63 (242), S. 937–949. DOI:
950 10.1017/jog.2017.48.
- 951 Meysman, Filip J. R.; Boudreau, Bernard P.; Middelburg, Jack J. (2003): Relations between local,

- 952 nonlocal, discrete and continuous models of bioturbation. In: *Journal of Marine Research* 61
 953 (3), S. 391–410. DOI: 10.1357/002224003322201241.
- 954 Monteverde, Martín J.; Piudo, Luciana (2011): Activity Patterns of the Culpeo Fox (*Lycalopex*
 955 *Culpaeus Magellanica*) in a Non-Hunting Area of Northwestern Patagonia, Argentina. In:
 956 *Mammal Study* 36 (3), S. 119–125. DOI: 10.3106/041.036.0301.
- 957 Morris, Rowena H.; Buckman, Solomon; Connelly, Paul; Dragovich, Deirdre; Ostendorf, Bertram;
 958 and Bradstock, Ross A. (2011): The dirt on assessing post-fire erosion in the Mount Lofty
 959 Ranges: comparing methods.
- 960 Nasermoaddeli, M. B.; Pasche, E. (2008): Application of terrestrial 3D scanner in quantification of
 961 the riverbank erosion and deposition. Institute of river and coastal engineering, Technical
 962 university Hamburg-Harburg, Hamburg,. Online verfügbar unter
 963 [https://www.tuhh.de/t3resources/wb/Publikationen/MA-](https://www.tuhh.de/t3resources/wb/Publikationen/MA-Veroeffentlichungen/nasermoaddelli/riverflow2008.pdf)
 964 [Veroeffentlichungen/nasermoaddelli/riverflow2008.pdf](https://www.tuhh.de/t3resources/wb/Publikationen/MA-Veroeffentlichungen/nasermoaddelli/riverflow2008.pdf), zuletzt geprüft am 22.12.2021.
- 965 Pang, Xiao Pan; Guo, Zheng Gang (2017): Plateau pika disturbances alter plant productivity and
 966 soil nutrients in alpine meadows of the Qinghai-Tibetan Plateau, China. In: *Rangel. J.* 39 (2), S.
 967 133. DOI: 10.1071/RJ16093.
- 968 Reichman, O. J.; Seabloom, Eric W. (2002): The role of pocket gophers as subterranean ecosystem
 969 engineers. In: *Trends in Ecology & Evolution* 17 (1), S. 44–49. DOI: 10.1016/S0169-
 970 5347(01)02329-1.
- 971 Richards, Paul J.; Humphreys, Geoff S. (2010): Burial and turbulent transport by bioturbation: a 27-
 972 year experiment in southeast Australia. In: *Earth Surf. Process. Landforms* 21 (2), n/a-n/a. DOI:
 973 10.1002/esp.2007.
- 974 Ridd, Peter V. (1996): Flow Through Animal Burrows in Mangrove Creeks. In: *Estuarine, Coastal*
 975 *and Shelf Science* 43 (5), S. 617–625. DOI: 10.1006/ecss.1996.0091.
- 976 Romañach, Stephanie S.; Reichman, O. J.; Seabloom, E. W. (2005): Seasonal influences on
 977 burrowing activity of a subterranean rodent, *Thomomys bottae*. In: *Journal of Zoology* 266 (3),
 978 S. 319–325. DOI: 10.1017/S0952836905006941.
- 979 Rutin, J. (1996): The burrowing activity of scorpions (*Scorpio maurus palmatus*) and their potential
 980 contribution to the erosion of Hamra soils in Karkur, central Israel. In: *Geomorphology* 15 (2),
 981 S. 159–168. DOI: 10.1016/0169-555X(95)00120-T.
- 982 Sarbolandi, Hamed; Plack, Markus; Kolb, Andreas (2018): Pulse Based Time-of-Flight Range
 983 Sensing. In: *Sensors (Basel, Switzerland)* 18 (6). DOI: 10.3390/s18061679.
- 984 Schiffers, Katja; Teal, Lorna Rachel; Travis, Justin Mark John; Solan, Martin (2011): An open
 985 source simulation model for soil and sediment bioturbation. In: *PloS one* 6 (12), e28028. DOI:
 986 10.1371/journal.pone.0028028.
- 987 Sharon, David (1980): The distribution of hydrologically effective rainfall incident on sloping
 988 ground. In: *Journal of Hydrology* 46 (1-2), S. 165–188. DOI: 10.1016/0022-1694(80)90041-4.
- 989 Thomsen, L. M.; Baartman, J. E. M.; Barneveld, R. J.; Starkloff, T.; Stolte, J. (2015): Soil surface
 990 roughness: comparing old and new measuring methods and application in a soil erosion model.
 991 In: *SOIL* 1 (1), S. 399–410. DOI: 10.5194/soil-1-399-2015.
- 992 Übernichel, Kirstin; Ehlers, Todd A.; Paulino, Leandro; Fuentes Espoz, Juan-Pablo (2021a): Time
 993 series of meteorological stations on an elevational gradient in National Park La Campana, Chile.
 994 Unter Mitarbeit von Kirstin Übernichel, Todd A. Ehlers, Leandro Paulino, Juan-Pablo Fuentes
 995 Espoz, Ramiro Bernales-Noguera, Willi Kappler et al.
- 996 Übernichel, Kirstin; Pizarro-Araya, Jaime; Bhagavathula, Susila; Paulino, Leandro; Ehlers, Todd A.
 997 (2021b): Reviews and syntheses: Composition and characteristics of burrowing animals along a
 998 climate and ecological gradient, Chile. In: *Biogeosciences* 18 (20), S. 5573–5594. DOI:
 999 10.5194/bg-18-5573-2021.
- 1000 Voiculescu, Mircea; Ianăș, Ana-Neli; Germain, Daniel (2019): Exploring the impact of snow vole
 1001 (*Chionomys nivalis*) burrowing activity in the Făgăraș Mountains, Southern Carpathians
 1002 (Romania): Geomorphic characteristics and sediment budget. In: *CATENA* 181 (1), S. 104070.
 1003 DOI: 10.1016/j.catena.2019.05.016.

- 1004 Wei, Xinghu; Li, Sen; Yang, Ping; Cheng, Huaishun (2007): Soil erosion and vegetation succession
1005 in alpine Kobresia steppe meadow caused by plateau pika—A case study of Nagqu County,
1006 Tibet. In: *Chin. Geograph.Sc.* 17 (1), S. 75–81. DOI: 10.1007/s11769-007-0075-0.
- 1007 Wilkinson, Marshall T.; Richards, Paul J.; Humphreys, Geoff S. (2009): Breaking ground:
1008 Pedological, geological, and ecological implications of soil bioturbation. In: *Earth-Science*
1009 *Reviews* 97 (1-4), S. 257–272. DOI: 10.1016/j.earscirev.2009.09.005.
- 1010 Yair, A. (1995): Short and long term effects of bioturbation on soil erosion, water resources and soil
1011 development in an arid environment. In: *Geomorphology* 13 (1-4), S. 87–99. DOI:
1012 10.1016/0169-555X(95)00025-Z.
- 1013 Yoo, Kyungsoo; Amundson, Ronald; Heimsath, Arjun M.; Dietrich, William E. (2005): Process-
1014 based model linking pocket gopher (*Thomomys bottae*) activity to sediment transport and soil
1015 thickness. In: *Earth Surf. Process. Landforms* 33 (11), S. 917. DOI: 10.1130/G21831.1.
- 1016
- 1017